Multipoint Thermospheric Wind Observations: Networks of Fabry-Perot Interferometers and Plans for ICON

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Special thanks to:

- Daniel Fisher
- Rafael Mesquita
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- ICON Team
- NATION Team





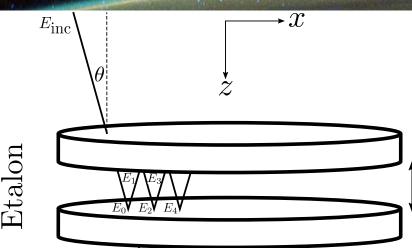
Outline

- Networks of ground-based wind instruments
- Midnight temperature maximum
 - Direct evidence of converging wind field
 - Propagation to midlatitudes
- ICON wind measurements





Thermosphere





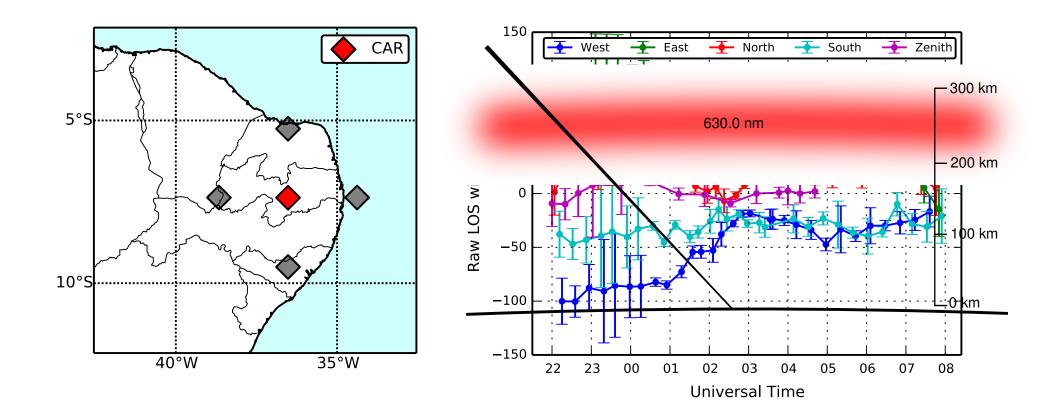
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Doppler shift \rightarrow Wind Doppler width \rightarrow Temperature



630.0 nm

A single Fabry-Perot interferometer (FPI)

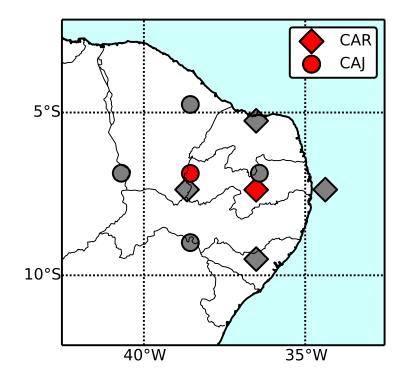


Estimate temporal evolution of the wind over a small (~500km) region.



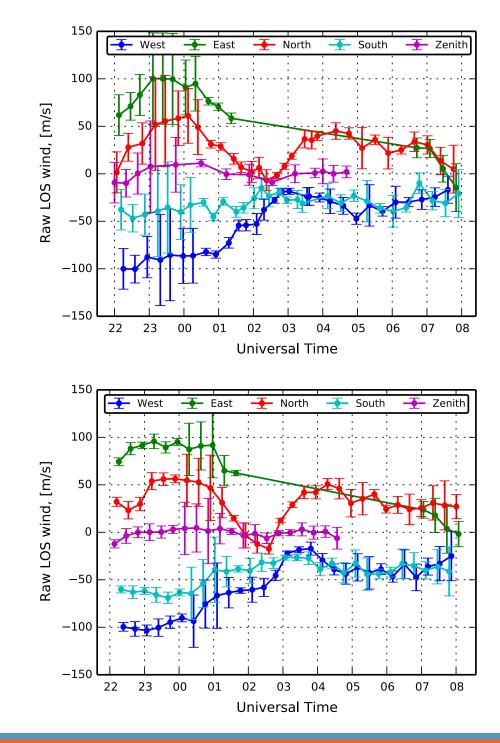


A network of FPIs

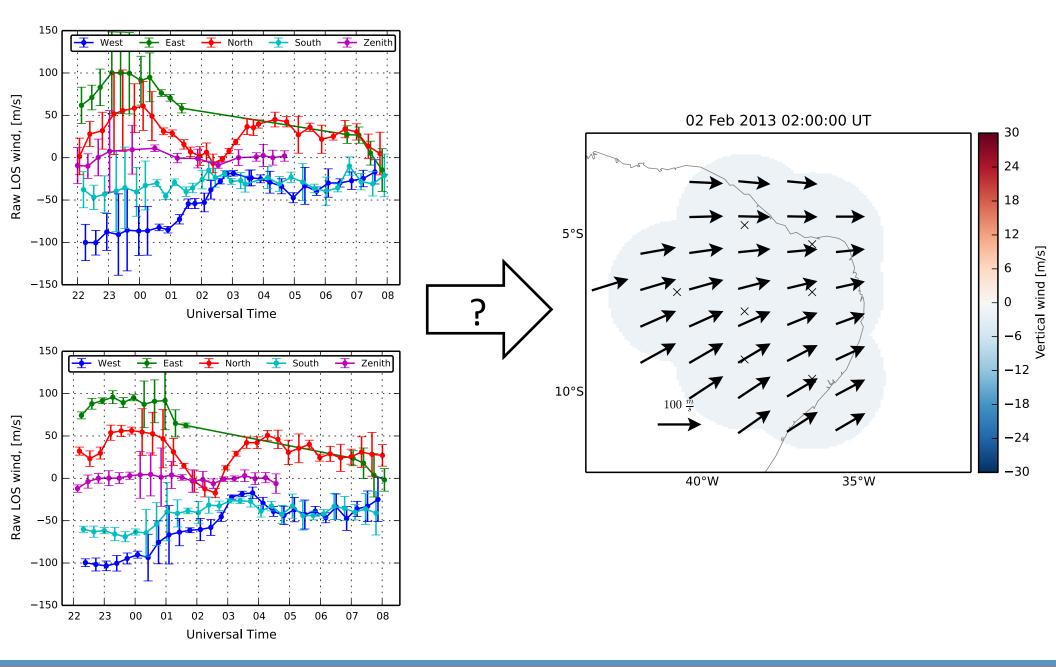


Observe temporal and spatial distributions

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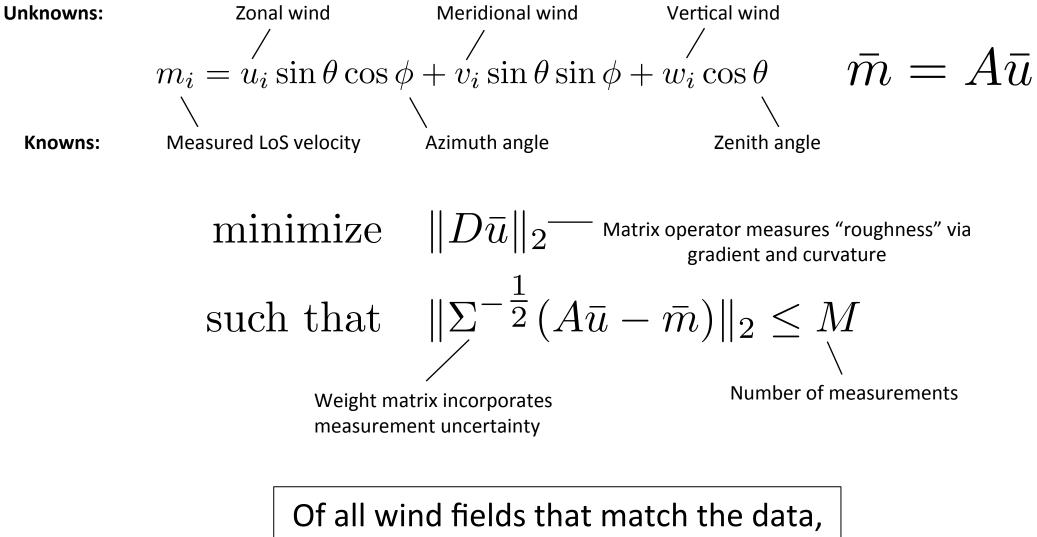
Wind Field Estimation







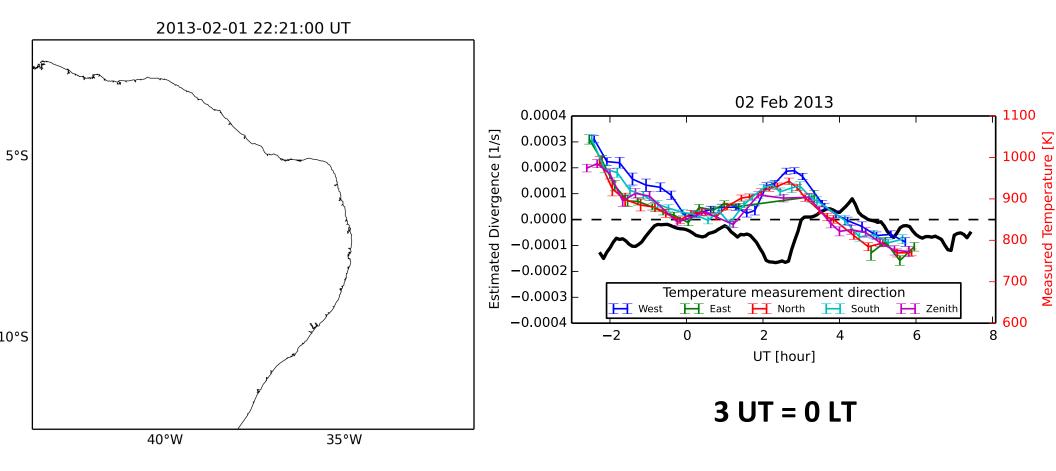
Wind Field Estimation



choose the smoothest.



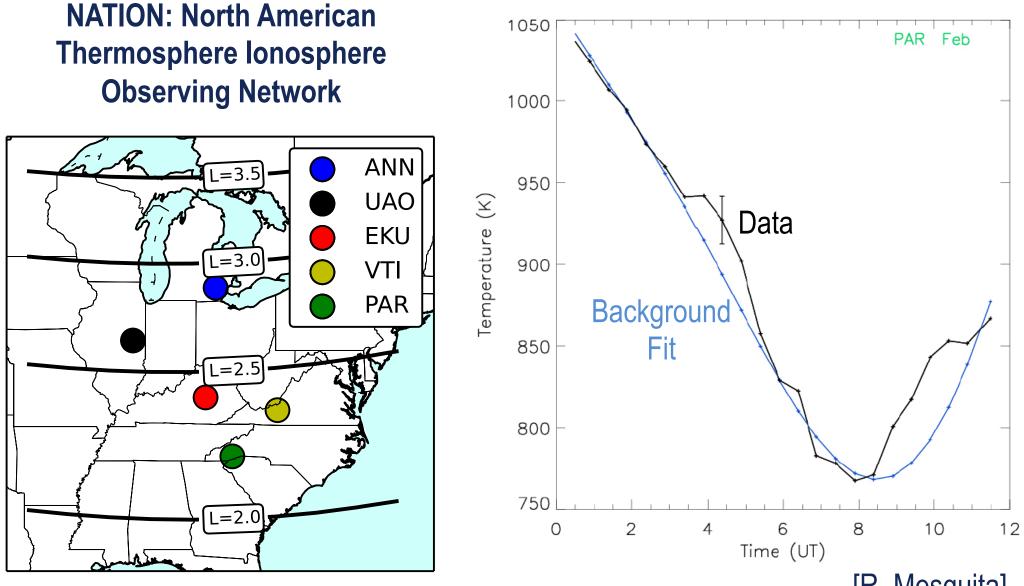
Midnight Temperature Maximum



Direct evidence of convergent wind field during MTM



Midnight temperature maximum... at midlatitudes?



[R. Mesquita]

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Midnight temperature maximum... at midlatitudes?

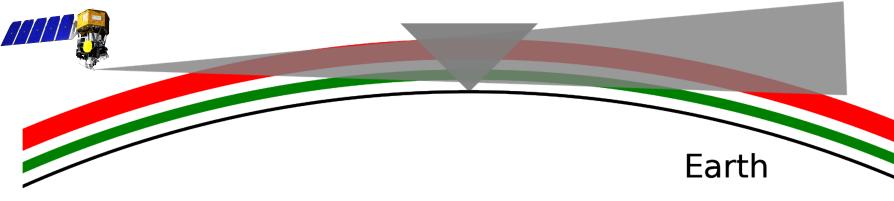


- Direct observation of winds, propagation
- Consistent with timing reported by *Hickey et al.* [2014] (ion and neutral temperatures)
- Propagation to 45°N

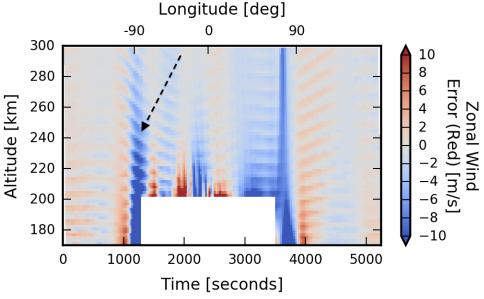




Ionospheric Connection Explorer



- MIGHTI is ICON's wind instrument
- System simulations suggest:
 - Expected Precision: 1-4 m/s (1σ)
 - At 5x500km (Vert x Horz) res (Red 30x500)
 - Systematic Error: **3-6 m/s** (80th prctile)



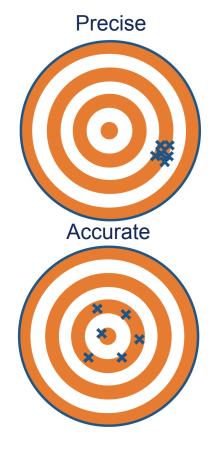
Bias

Error

- 2nd moment: precision, statistical uncertainty
- 1st moment: accuracy, bias
- Data providers rarely report 1st moment
 - Important for assimilation and data fusion
- Geophysical data often have a bias towards "smoothness"
- Can quantify with resolution matrix:

if
$$y = Ax + n$$

and $x^* = A^p y$
then $R = A^p A$



0336.00 UT

J. Gjerloev, CEDAR Prize Lecture 2016]



Conclusions

- Networks of FPIs allow us to see spatial structure, revealed by statistical estimation techniques
 - Direct evidence of converging wind field during MTM
 - Extension up to midlatitudes
- ICON/MIGHTI will provide global coverage
- Bias is just as important as uncertainty, especially for disparate data sets



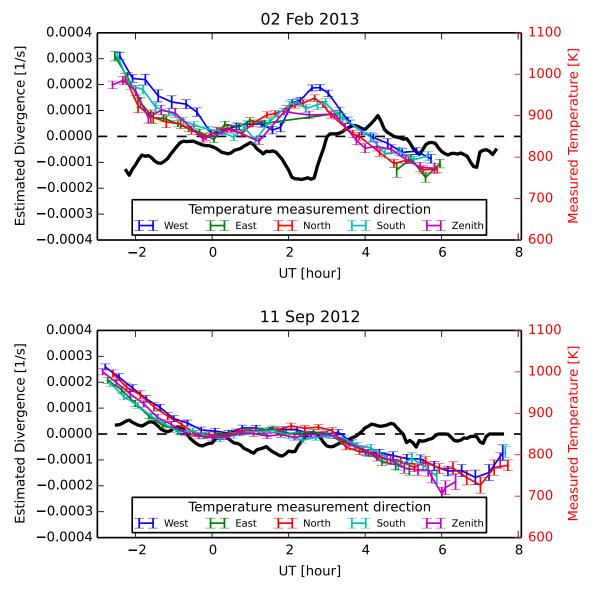


Backup Slides





Application 1: Midnight Temperature Maximum

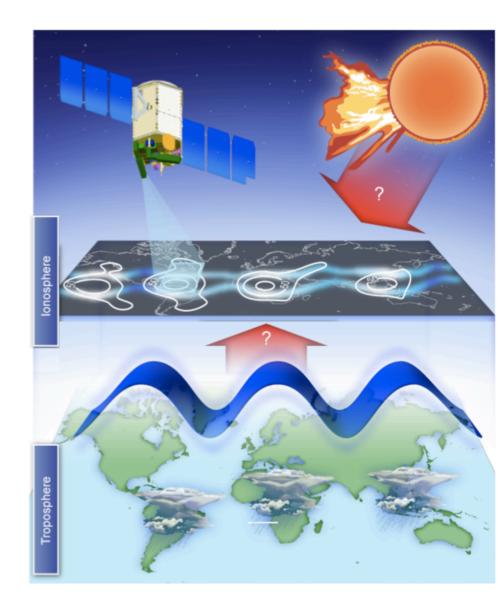


- Significant variability in amplitude of MTM
- Thought to be related to terdiurnal tide



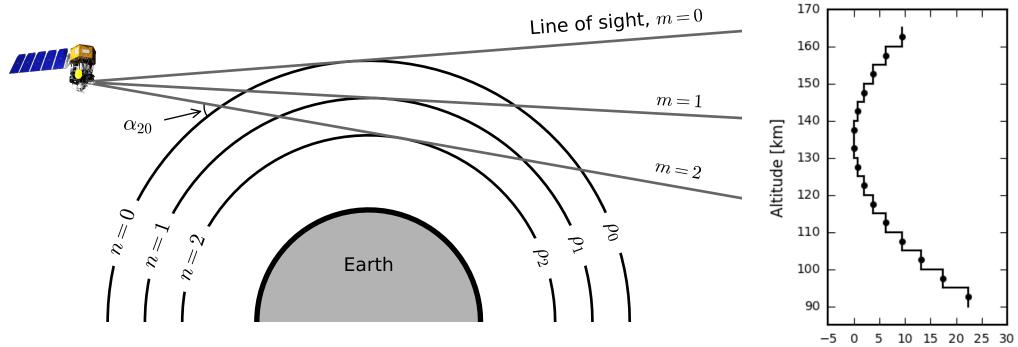
ICON Mission Goal

- Identify drivers of extreme ionospheric variability
 - Cause of structure at low latitudes during quiet times
 - Coupling processes with large-scale atmospheric waves
 - Competition of these internal drivers with external drivers during geomagnetic storms



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Problem: Scrambled spectra from different altitudes



Wind [m/s]

NOIS

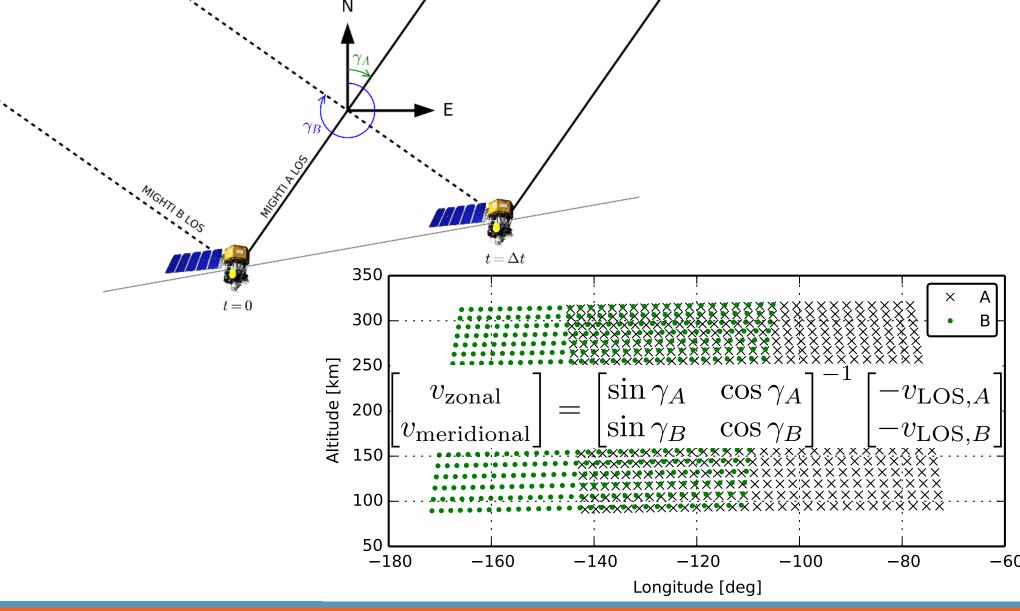
Solution: Abel-like inversion

Assume spherical symmetry and piecewise constant altitude dependence

$$g = \int_0^\infty f(r) \, ds \longrightarrow g = \sum_{n=0}^{N-1} f_n \, \Delta s_n \longrightarrow \mathbf{g} = \mathbf{S}\mathbf{f}$$

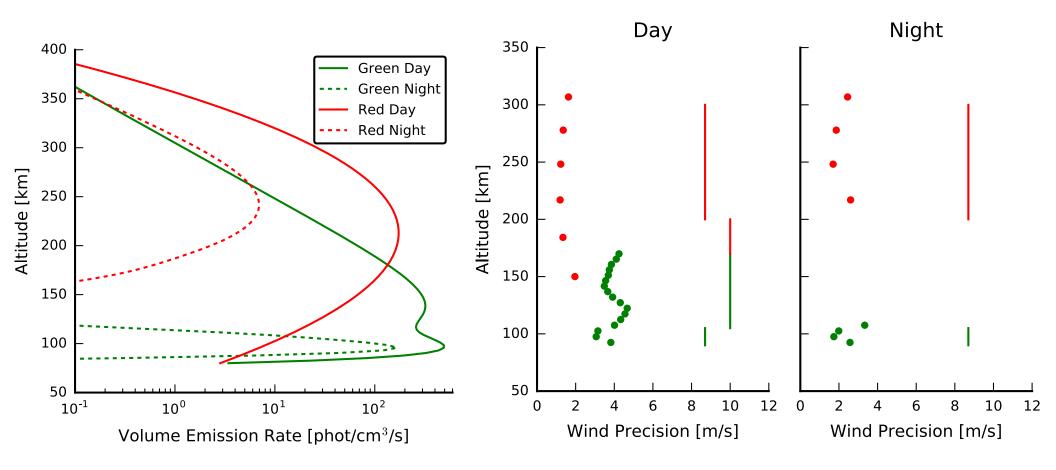


Problem: Only one component measured Solution: Use two orthogonal sensors



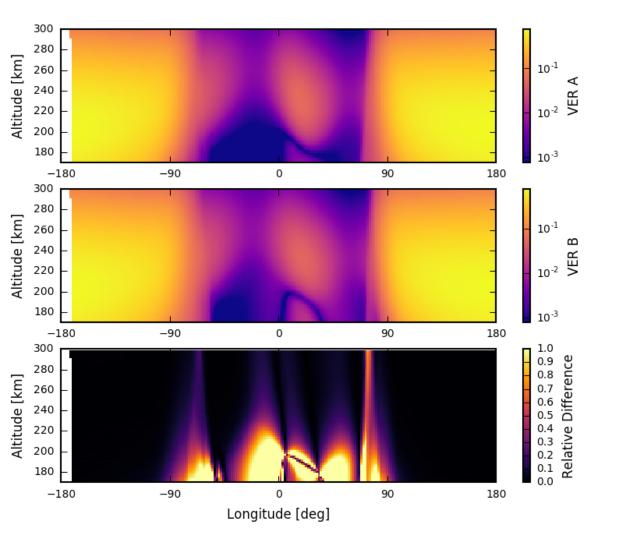
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Expected Precision

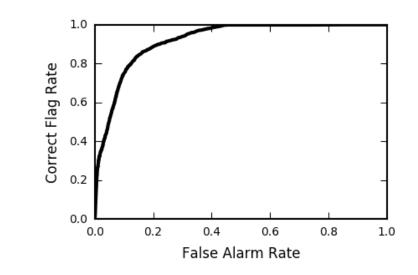




Spherical Asymmetry Flag – Red

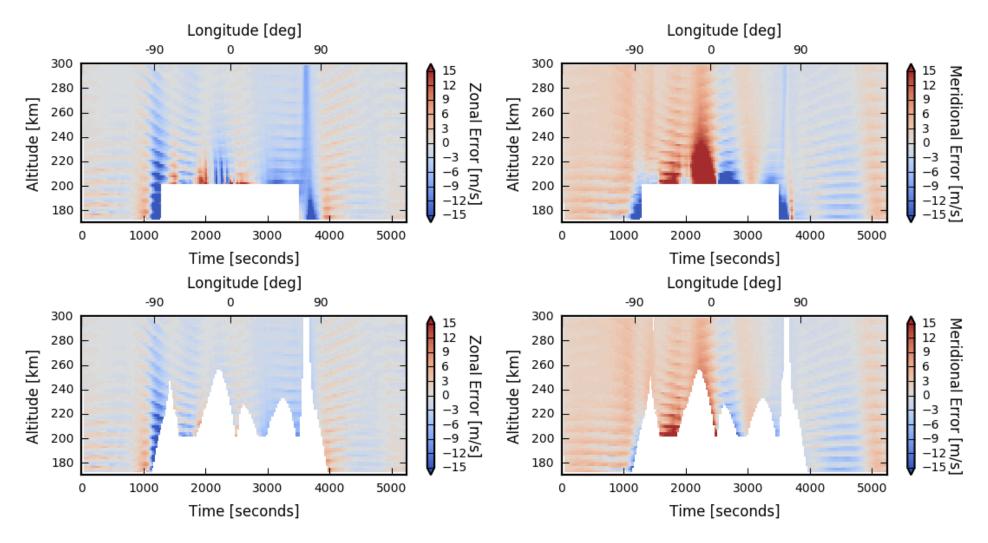


- Compare VER estimates from A and B
- Use threshold to flag expected errors in wind estimate



LINOIS



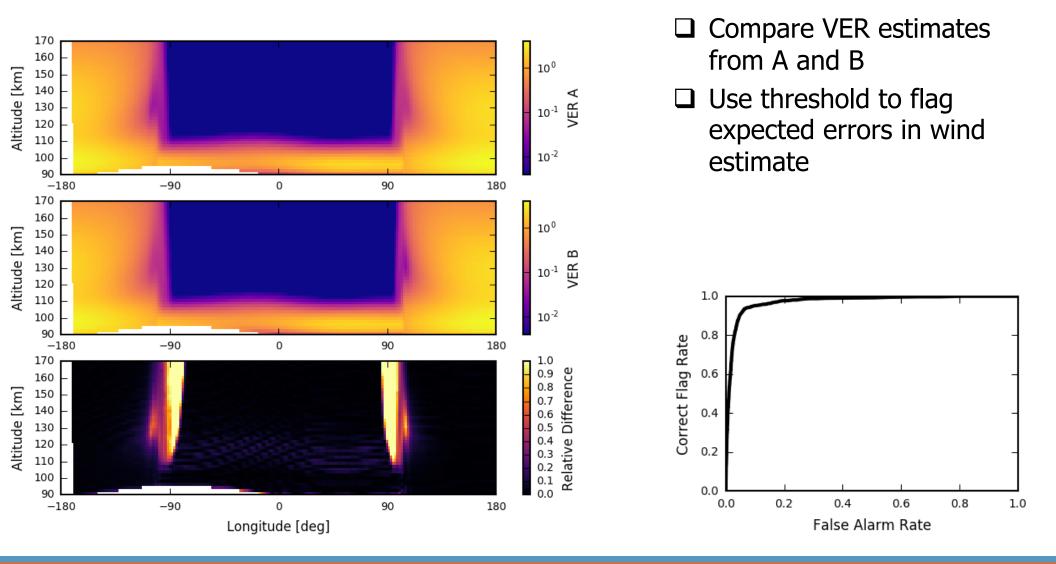


 $6.5 \rightarrow 3.9 \text{ m/s rms error}$

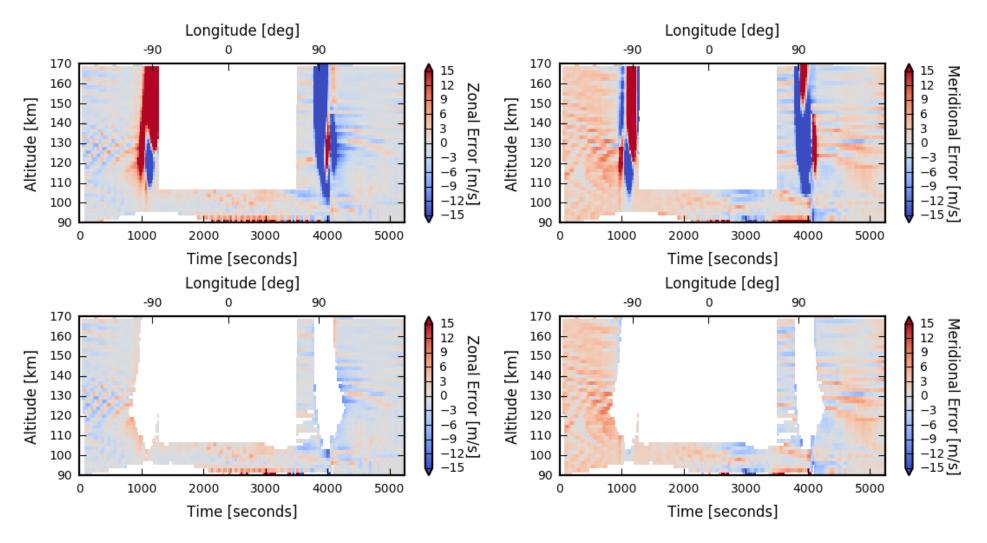




Spherical Asymmetry Flag – Green



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 $26.0 \rightarrow 4.5 \text{ m/s rms error}$





Airglow above 300 km

- Assuming:
 - Constant wind & temp. above 300km
 - Exponential falloff in VER (H=26km)
- Effective path length in top layer:
- New algorithm works well for contrived cases
- Doesn't make a difference in our realistic end-to-end simulation

$$w_{m0}^{\text{above}} = \int_{\sqrt{r_0^2 - r_m^2}}^{\infty} e^{-\frac{1}{H}\left(\sqrt{x^2 + r_m^2} - r_0\right)} dx$$

cases $\sqrt{r_0^2 - r_m^2}$

