Mapping ionospheric electrodynamics with AMPERE and SuperDARN data

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Uncertainty quantification
 Specifying conductance
 Validation of results



earth• connections

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Mapping Procedure

- Assimilative technique with Empirical Orthogonal Function (EOF) based error covariance
- Traditionally solved in terms of electrostatic potential; can also be solved in terms of magnetic vector potential



Evaluating Conductance Models

Want quantitative metric for selecting 'best' conductance model Solar Noon Evaluate by using mapping procedure with just one data type, predict observations of other type e.g., fit electrostatic potential with SuperDARN data, predict AMPERE observations (using assumed conductance), Compare predictions with actual observations Pick conductance model that minimizes discrepancy between **E** obs & ΔB obs urora Best results: diffuse aurora from Ovation Prime or SM 3.0 11.15 **Night-side** [*Newell et al.*, 2010; *Mitchell et al.*, 2013], w/ night-side $\Sigma_n = 4$ S

		O-P	FAC adj	>2, O-P	>4, O-P	>4, O-SM	>4, no aur	>6, O-P
∆B → E:	Med err [m/s]	523	513	172	147	146	147	142
E → ΔB:	Med err [nT]	33.2	33.3	33.5	34.7	34.6	34.7	36.7

Assimilative Mapping Example

- Solving with both data sets together
- Electrostatic potential map based on solving in terms of electrostatic potential (more weight for SuperDARN data)
- Field-aligned current map based on solving in terms of magnetic vector potential (more weight for AMPERE data)



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Assimilative Mapping – Benefits of using data sets together

- Analysis error decreases by 10% for Φ, 40% for A
- AMPERE fills in SuperDARN coverage gaps & vice versa
- Electric potential & FAC maps more compatible?
- feature correspondence 12 MLT 12 MLT 12 MLT b. 50 а. C. 1515 Independent 6 18 6' 18 mW/m 21 3.6 2.4mV/m nΤ 34 0.75 -0.8 -40 1.2 40 741 14 8 2 -0.4 12 MLT 12 ML 12 MI **d.** 50^c -50 -2.8 -4.0 е. 15° 15 Together 18 <u>'</u>6 6 18 +B₇ 43 0.52 53 -48 -0.66 $\mathbf{J}_{||}$ SII Φ

Summary of SuperDARN-AMPERE Mapping

- Mapping ionospheric electrodynamics using optimal interpolation data assimilation method with SuperDARN & AMPERE data
 - Solving for both electrostatic potential & magnetic vector potential
 - Using magnetic potential improves specification of FACs

Challenges

- Uncertainty quantification
 - EOF analysis of model errors
 - Sub-resolution variance of obs. (but what about biases?)
- Specifying conductance
 - Empirical models & obs.-based adjustments tested quantitatively
 - Best results w/ Ovation Prime (or SM) diffuse aurora & night-side $\Sigma_P = 4$ S
 - But even 'best' model is not representative of instantaneous reality...
- Validation of results
 - Improvement in qualitative appearance
 - Quantitative reduction in analysis error
 - Other approaches to validation?

[Extra Slides]

SuperDARN data

- 19 radars in the Northern Hemisphere, ~50° and poleward
- 'Grid' level data
 - Spatial resolution: 110 km
 - Temporal resolution: 2 min
- Line-of-sight (LOS) observations of
 ExB plasma drifts at ~300 km altitude





Active Magnetosphere and Planetary Electrodynamics Response Experiment

Iridium for Science

- Magnetometer on every satellite
- 6 orbit planes

 (12 cuts in local time)
 ~11 satellites/plane
- 9 minute spacing re-sampling cadence
- 780 km altitude, circular, polar orbits



[Brian Anderson]

CEDAR 2014, Hi-Lat Data Assim

AMPERE Data: 26 Nov – 2 Dec, 2011



AMPERE Data: 26 Nov – 2 Dec, 2011



Assimilative Mapping Procedure

- Use the optimal interpolation (OI) method of data assimilation
 - Optimally combine information from observations and a background model, taking into account error properties of both
- Background model:
 - Electrostatic potential: SuperDARN CS10 empirical convection model
 - Magnetic vector potential: AMPERE mean
- Error properties of background models estimated using Empirical Orthogonal Function (EOF) analysis
- Observational errors based on local small-scale variance values

$$\mathbf{x}_a = \mathbf{x}_b + \mathbf{K} \left(\mathbf{y} - \mathbf{H} \mathbf{x}_b \right)$$

$$\mathbf{K} = \mathbf{P}_b \mathbf{H}^T (\mathbf{H} \mathbf{P}_b \mathbf{H}^T + \mathbf{R})^{-1}$$

- \mathbf{x}_a analysis
- y observations
- \mathbf{x}_{b} background model
- H~- forward operator [physics + Σ]
- $\mathbf{K}~$ Kalman gain
- \mathbf{P}_b background model error covariance
- ${\bf R}~$ observational error covariance

AMPERE data: Uncertainty Estimation

- Estimate uncertainty by looking at variation in values below resolution of basis functions (~2.5°Λ, 1h MLT)
- For vectors with <5 neighbors, use median of variance values calculated as above, multiplied by data quality flag (~0.7 – 1.3)



November 29, 2011 14:40:00 - 14:50:00 UT



Empirical Orthogonal Function (EOF) analysis

- A variant of principal component analysis (PCA)
- Technique to estimate dominant modes of variability in a data set (Decompose into eigenmodes & eigenvalues)
- Calculated in terms of magnetic vector potential: $\Delta \vec{B}_{hor} = \nabla \times \vec{A}_r$

$$A_{r}(\mathbf{x},t) \approx A_{r}^{mean}(\mathbf{x}) + \sum_{v} \alpha^{(v)}(t) EOF^{(v)}(\mathbf{x})$$

$$A_{r}^{\mathsf{bor}} = \nabla \times \Delta \vec{B}_{hor}$$

$$A_{r}^{\mathsf{bor}} = -\nabla^{2}A_{r}^{\mathsf{bor}}$$

Data quality issues

- Difficult to obtain clean EOFs using just 1 week of data with low SNR
- Higher quality data in across-track direction vs. along-track direction (likely due to attitude determination errors)
- Better EOFs using just across-track data



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EOF Properties: Winter/Summer

- EOFs calculated independently for North/Winter & South/Summer hemispheres
- Dawn-dusk asymmetry in mean different in two hemispheres
- Winter hemi. shows more night-side features, summer more day-side



Universal EOFs

- Ideal to have set of universal EOFs that can be used to describe any AMPERE data
 - Obtained by combing results from North/Winter & South/Summer



EOF properties: Power Spectrum

- Mean accounts for more of total ΔB² in summer than in winter
- Mean + 6 EOFs accounts for < 50% of total ΔB^2
 - → Significant amount of small-scale variability &/or noise in data









