# **Development and Validation of an Airglow Photometer for Upper**

# **Atmospheric Chemistry**

Yi-Chung Chiu<sup>1</sup>, Yi Duann<sup>1</sup>, Loren Chang<sup>1</sup>, J. B. Nee<sup>2</sup>

<sup>1</sup>Institute of Space Science, National Central University, Taiwan <sup>2</sup>Department of Physics, National Central University, Taiwan Contact : yc.small.phi@gmail.com



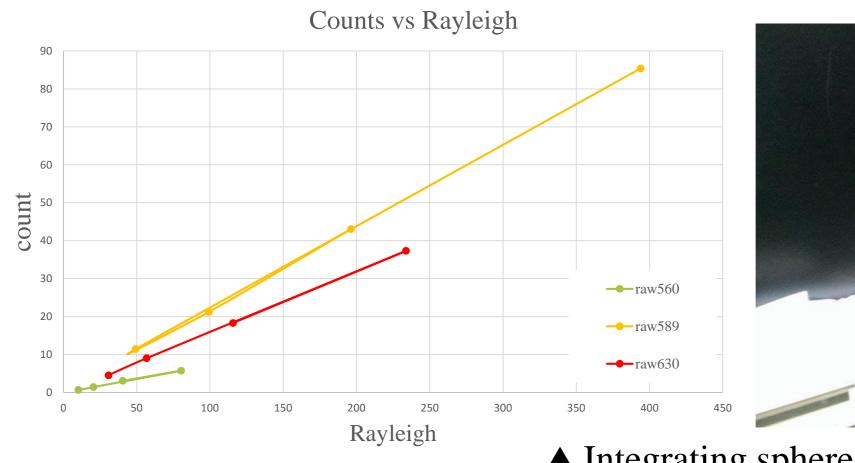
## Abstract

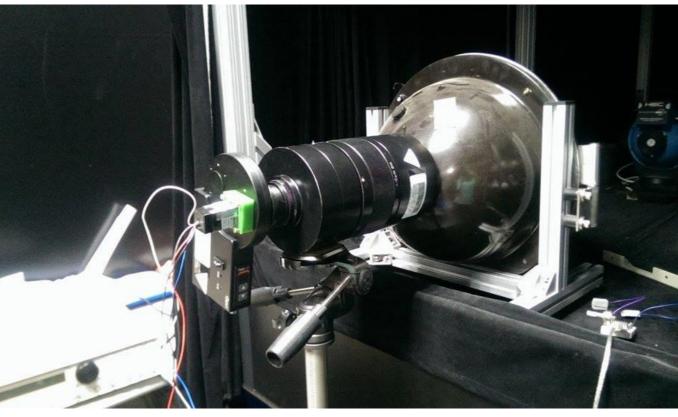
NCU-GISS

Airglow is a phenomenon caused by chemical reactions in the mesosphere and thermosphere, and can also be used as a proxy to measure important parameters in upper atmosphere. Due to this reason, we have built a simple airglow photometer system and plan to deploy it for long term ground-based observations at Lulin Observatory in Taiwan (120°52'25''E, 23°28'07''N). We have selected three airglow emission channels (557.7nm, 589nm, 630nm) as our observation target, which have been the subject of several past studies. In order to get reasonable data from our airglow observations, we need to validate all parts of our system, design, calibration, and the data processing procedure. We have used an integrating sphere to determine the response of our photometer to different levels of irradiance at different channels, and have performed three nights of airglow observations. In the future, we will use our observation results to compare with some atmospheric events and determine chemical changes in the mesosphere and thermosphere. After demonstrating that the data we use is reliable, we will provide long-term observations and monitoring of airglow emission rates and chemical processes over Taiwan.

Background & Motivation			Instrument				
Oxygen	<ul><li>[O] affects:</li><li>Satellite drag in LEO.</li><li>Atmospheric vertical structure.</li></ul>	OI Red Line (630 nm) $O^+ + O_2 \rightarrow O_2^+ + O$ $O^+ + O_2 \rightarrow O_2^+ + O$		Photometer block diagram and channel information. $\mathbf{\nabla}$			

concentration	Ionospheric plasma lifetime.	$O_2^+ + e^- \to O({}^1D) + O({}^3P)$	Telescope lens									
Airglow emissionProxy for [O]: 		$O({}^{1}D) \to O({}^{3}P) + hv(630 nm)$ OI Green Line (557.7 nm) $0 + 0 + M \to O_{2}^{*} + M$ $O_{2}^{*} + 0 \to O_{2} + O({}^{1}S)$ $O({}^{1}S) \to O({}^{1}D) + hv(557.7)$		Filter wheel	Front side system	Preamplifier	Photon counter	<ul> <li>Data flow</li> <li>Operation control</li> <li>Power control</li> </ul>				
				Photomultiplier tube (PMT) Power supply Computer								
		Na D Line 589.0 & 589.6 nm $Na + O_3 \rightarrow NaO + O_2$ $NaO + O \rightarrow Na({}^2P_{1/2}) + O$ $\rightarrow Na({}^2P_{3/2}) + O$ $Na({}^2P_{1/2}) \rightarrow Na({}^2S) + hv(589.0 nm)$ $Na({}^2P_{3/2}) \rightarrow Na({}^2S) + hv(589.6 nm)$ Airglow emissions observed $\blacktriangle$		Channels	Wavelength (nm)	Bandwidth (nm)	Transmission	Function				
	Our objectives:			1	530	10	50 %	Clear sky				
Ground based	• Develop the photometer system.			3	560	10	65 % 33 %	OI green line Na D line				
observation	• Long term ground based observation.			4	630	10	60 %	OI red line				
				5	Black mask	_	-	Dark current				
				6	-	-	_	-				
Calibration		Observation										
Intensity of our target. Rayleigh-Radiance transform. Rayleigh transform. Rayleigh transform. Rayleigh transform. Rayleigh transform.												



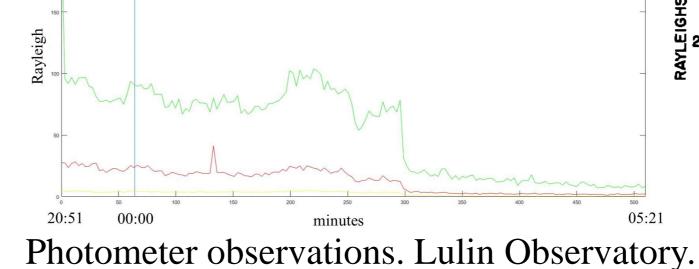


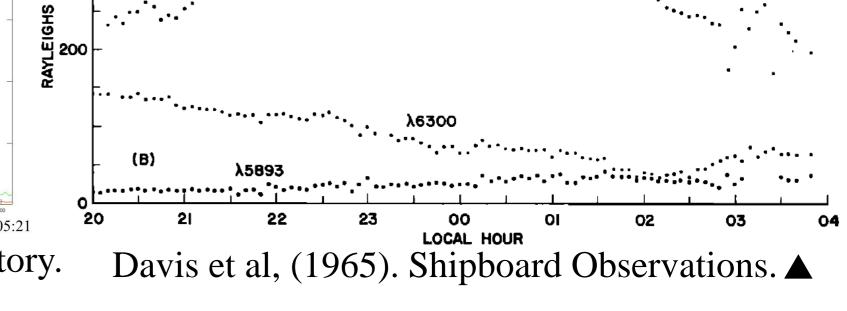
 $\blacktriangle$  Integrating sphere calibration  $\blacktriangle$ 

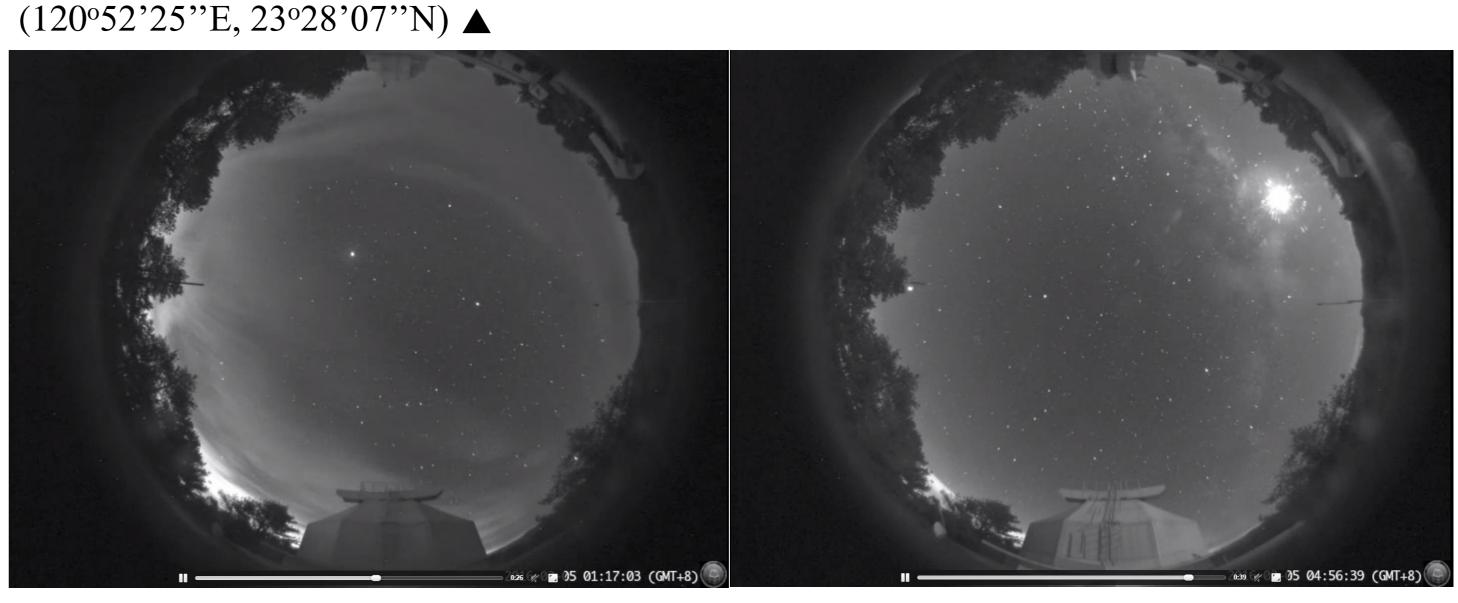
- The airglow intensity is so weak that we should use photon counting, by measuring the electrical pulses from the PMT.
- We also need to know the absolute airglow intensity in Rayleighs, so we used an integrating sphere to obtain calibration curves in different airglow channels.
- Calibration curves show that our system is more sensitive to Na D line airglow emission and least sensitive to OI green line emission.

#### Summary

- Completed assembly, calibration, and functional testing of an airglow photometer for mesosphere and thermospheric monitoring.
- Cloudiness and ambient temperature are important factors in our observation.
- The variation for each channel is similar to previous observations.
- We shall build up a long term ground station at Lulin Observatory, and the system will



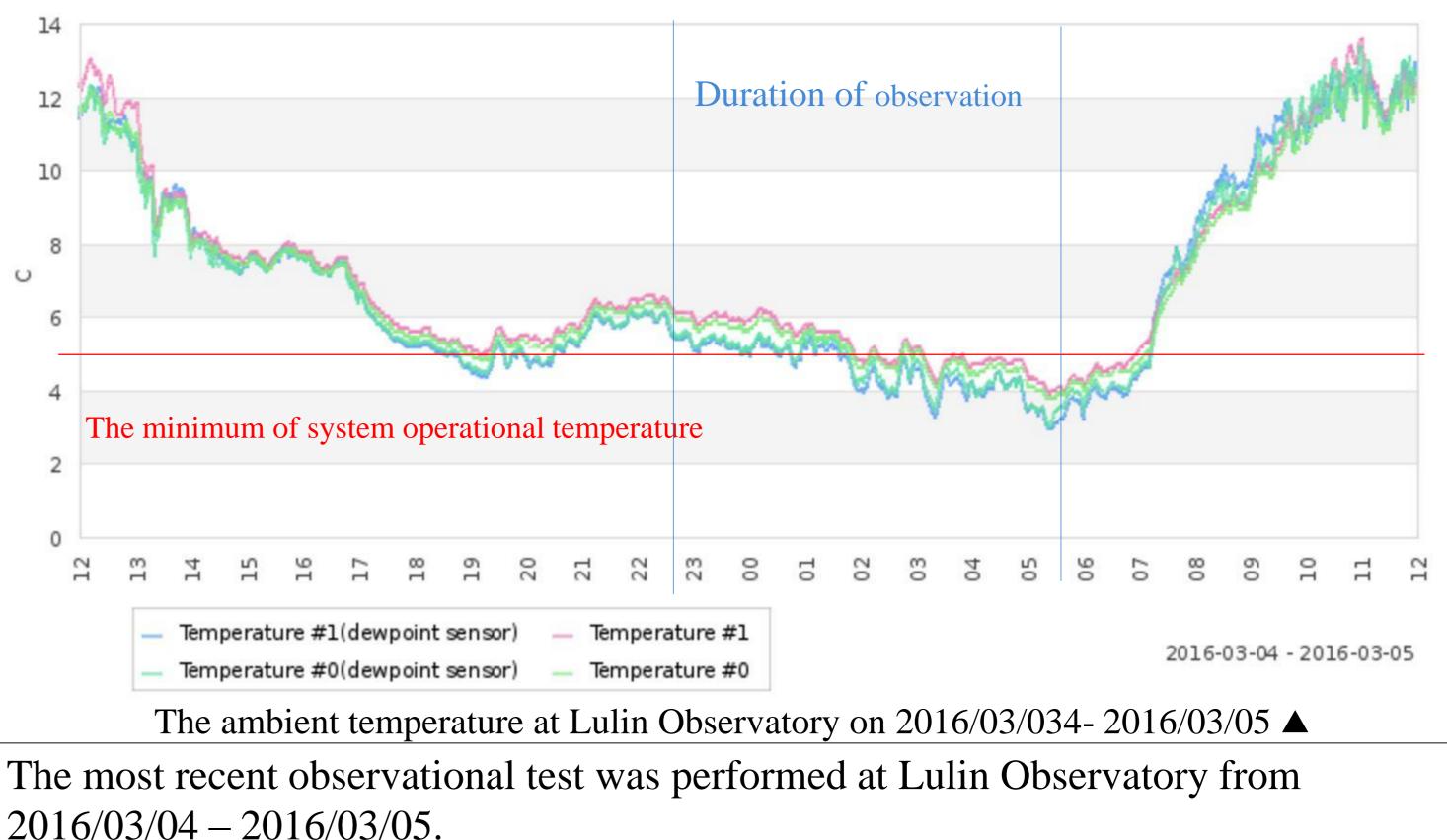




All sky image before 02:00 LT ▲

All sky image after 02:00 LT▲

Ambient Temperature



### be the remote controlled.

### Reference

- Davis, T. N., & Smith, L. L. (1965). Latitudinal and seasonal variations in the night airglow. *Journal of Geophysical Research*, 70(5), 1127-1138.
- Smith, L. L., & Steiger, W. R. (1968). Night airglow intensity variations in the [OI] 5577 A,[OI] 6300 A, and NaI 5890-96 A emission lines. *Journal of Geophysical Research*, 73(7), 2531-2538.
- Jana, P. K., & Nandi, S. C. (2006). Ozone decline and its effect on night airglow intensity of Na 5893 Å at Dumdum (22.5 N, 88.5 E) and Halley Bay (76 S, 27 W). *Journal of earth system science*, *115*(5), 607-613.
- Wu, Shin-Chun (2014). Ground observation by using six channels airglow photometer. National Central University Department of Physics Dissertation

## Acknowledgement

This research is funded by MOST103-2111-M-008-019-MY3 from the Ministry of Science and Technology. We thank the National Applied Research Laborites Instrument Technology Research Center help us to do the system calibration.

- The related differences of functional test results are similar as previous observations (Davis et al, 1965).
- The decrease airglow intensity after 02:00 LT is not related to cloud cover since the all sky camera shows fewer clouds after 02:00 LT. It is related to the lower ambient temperature after 02:00 LT.