



P. Joshi<sup>1</sup>, L. Waldrop<sup>1</sup>, C. Brum<sup>2</sup>, M. Sulzer<sup>2</sup>, N. Aponte<sup>2</sup>, S. Gonzalez<sup>2</sup>, P. Santos<sup>2</sup> and E. Robles<sup>2</sup>

1: Department of Electrical and Computer Engineering, University of Illinois at Urbana-Champaign, USA; 2: Arecibo Observatory, Puerto Rico



## INTRODUCTION

The objectives of this work are:

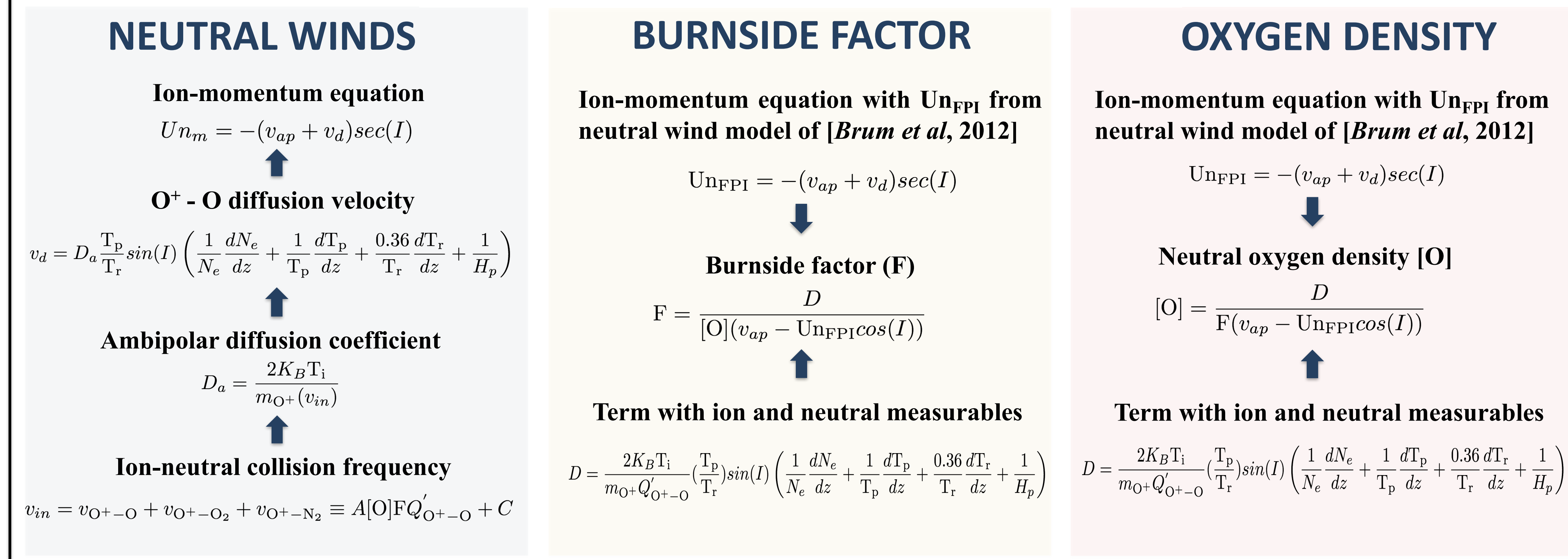
To derive the notorious "Burnside" factor scaling of the O-O+ charge exchange cross section,  $Q(O-O^+)$ , from momentum balance analysis in order to improve calculations of plasma drift speeds, diffusion coefficients, and electron density distributions.

To investigate potential bias in the model specification of O density, [O], used in the analysis.

What is the unique advantage of this work?

Use of unprecedented 18-year baseline of combined incoherent scatter radar and neutral wind data acquired at Arecibo Observatory.

## ION-MOMENTUM BALANCE CALCULATIONS



## UNCERTAINTIES

Altitude gradients in the plasma drift velocity which are currently neglected due to lack of height-resolved data.

Measurement uncertainties in the ISR data, which are not considered.

Assumption of negligible vertical wind gradients in the FPI wind model of Brum et al [2012].

Empirical calculation of peak emission altitude using Link and Cogger [1988]

Temperature dependence of the functional form of cross section ( $Q_{O^+-O}$ ) by Pesnell et al [1993]

Uncertainty in the MSIS00 derived neutral temperature  $T_n$

Tropospheric scattering effects in FPI measurements.

Possible effects of horizontal density gradients and wind field gradients.

## PARAMETER SPECIFICATION

Ionospheric state parameters:

Electron density  $[N_e]$ , ion densities  $[H^+]$  and  $[O^+]$ , ion temperature  $T_p$ , electron temperature  $T_e$ , plasma temperature  $T_p = 0.5(T_i + T_e)$  and, anti-parallel plasma drift velocity  $v_{ap}$  measured by the Arecibo ISR from 1987-2004.

Thermospheric density and temperature:

Co-located neutral densities  $[H]$ ,  $[O]$ ,  $[O_2]$ ,  $[N_2]$  and temperatures  $T_n$  and, average ion-neutral temperature  $T_r = 0.5(T_i + T_n)$ .

Thermospheric winds:

$$Un_m = U_e \sin(D) + U_n \cos(D)$$

where,  $Un_m$  is the neutral wind along the magnetic meridian, D is magnetic declination,  $U_e$  and  $U_n$  are zonal and meridional winds from the neutral wind model of Brum et al [2012] which is based on the 630 nm redline emission measured by Arecibo Fabry-Perot interferometer (FPI).

Charge exchange cross section:

$$Q'_{O^+-O} = 3 \times 10^{-11} \cdot \sqrt{T_r} \cdot (1 - 0.135 \log(T_r/10^3))^2 \text{ from Pesnell et al [1993]}$$

## METHODOLOGY

STEP 1:

For each of the 39,512 individual ISR measurements available, calculate the peak 630 nm emission altitude ( $h_{peak}$ ) using Link and Cogger [1988].

STEP 2:

Find the values of needed thermospheric and ionospheric parameters at  $h_{peak}$ .

STEP 3:

Use the O-O+ momentum balance equation to derive neutral wind  $Un_m$ , Burnside factor F, and neutral O density for each measurement.

STEP 4:

Average the observations over 10 minutes of local time and calculate errors in terms of both propagated measurement uncertainty (via Reddy et al [1994]) and standard deviation of the binned data.

ASSESS SEASONAL DEPENDENCE:

Bin data by season, defined as +/- 45 days from March equinox (21 March), June solstice (21 June), September equinox (21 September) and, December solstice (21 December), respectively.

ASSESS SOLAR CYCLE DEPENDENCE:

Bin data by solar activity, defined in sfu as: 90 < F10.7 < 120, 120 < F10.7 < 150, 150 < F10.7 < 180 and, 180 < F10.7 < 210.

## RESULTS

### SEASONAL VARIATION

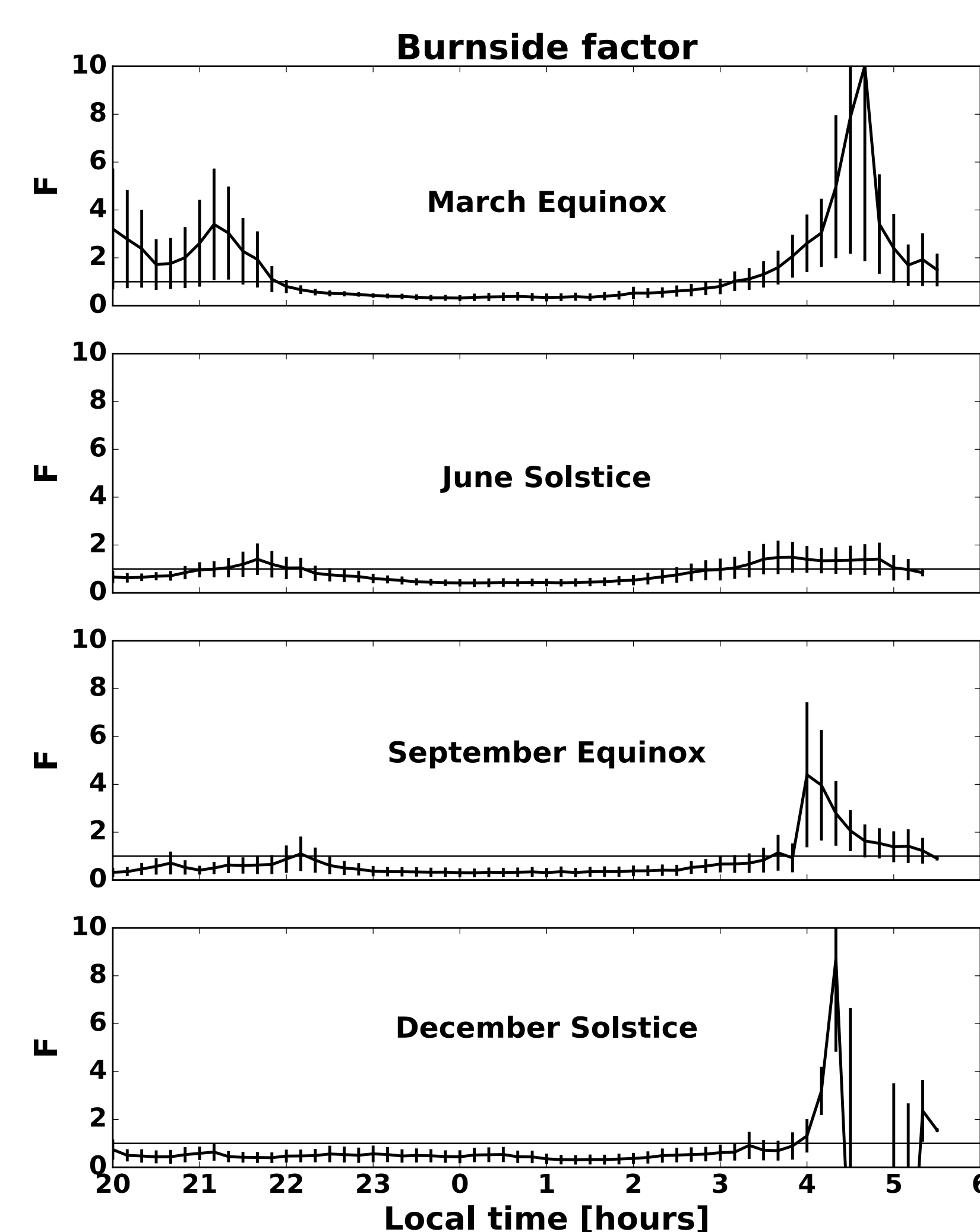


Figure 1. Local time/seasonal variation of the average Burnside factor (F) derived from the ISR-based momentum balance analysis using neutral wind ( $Un_m$ ) specified by the FPI wind model.

### SOLAR CYCLE VARIATION

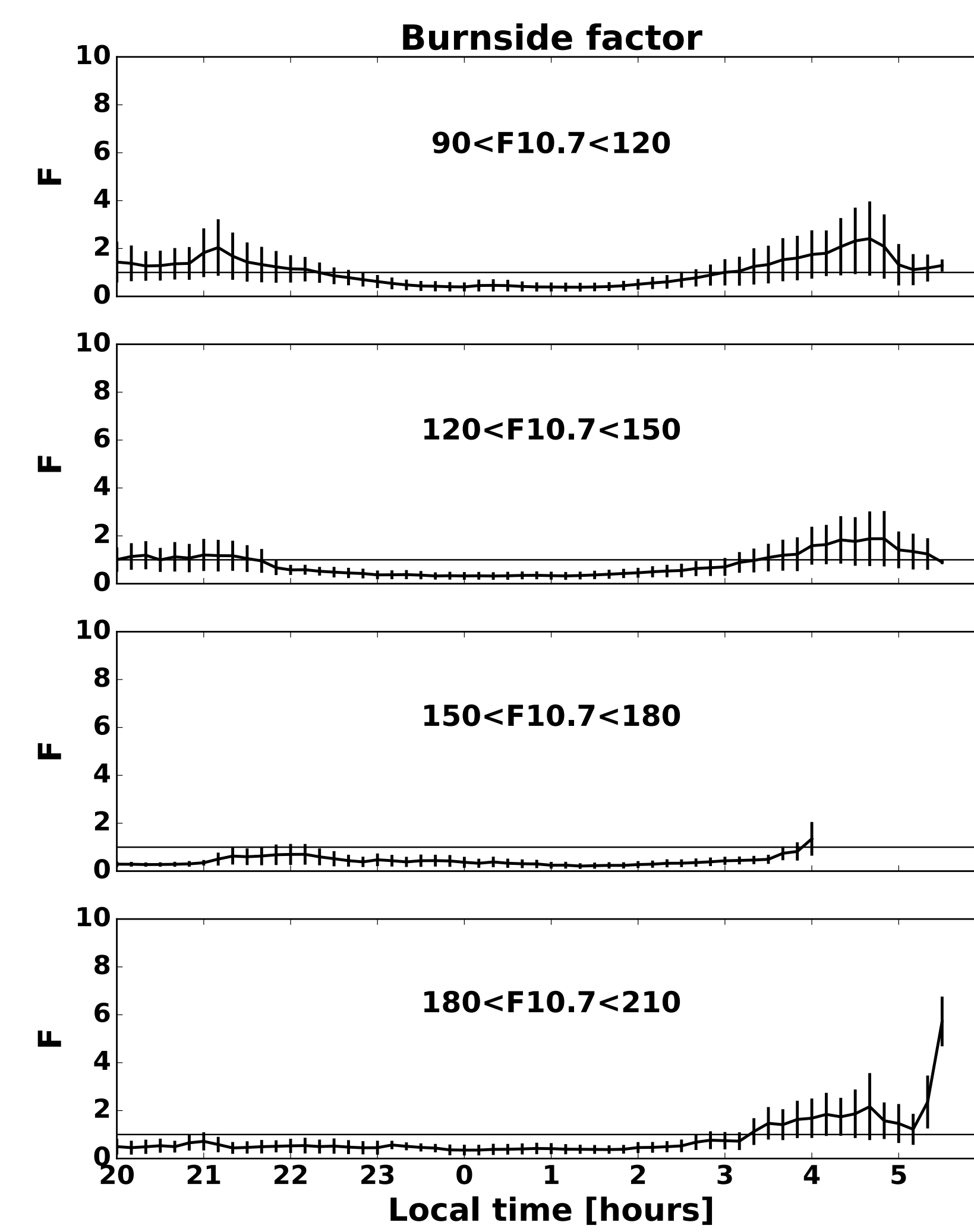


Figure 2. Local time/solar activity variation of the average Burnside factor (F) derived from the ISR-based momentum balance analysis using neutral wind ( $Un_m$ ) specified by the FPI wind model.

### FINDINGS

In Figures 1 and 2, a local time variation is observed in the Burnside factor where it is (<1) over the time from ~22 LT - 4 LT) and (>1) just post the sunset and then again near the pre-dawn interval ~4 LT - 6 LT).

The variations in the Burnside factor are consistent over all the seasons and solar activity levels.

We observe an average Burnside factor ( $F_{avg}$ ) of 0.5603 +/- 0.0336 during (22 LT - 4 LT) and 0.7112 +/- 1.049 in the predawn (4 LT - 6 LT) interval and 0.9001 +/- 0.1228 over the entire (22 LT - 6 LT) interval.

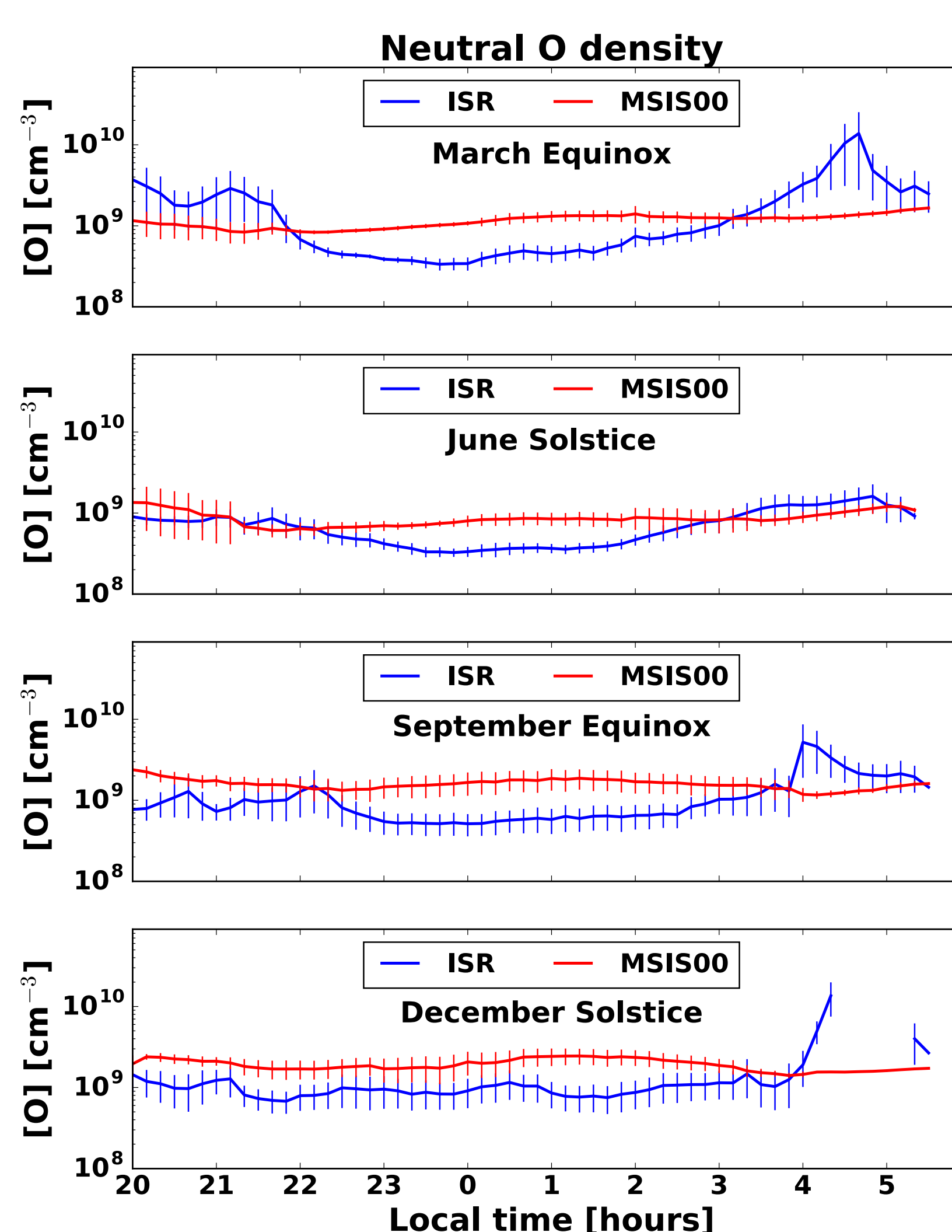


Figure 3. Local time/seasonal variation of the neutral oxygen density [O] specified by MSIS00 (red) and derived from the ISR-based momentum balance analysis (blue) using a Burnside factor  $F=1$ .

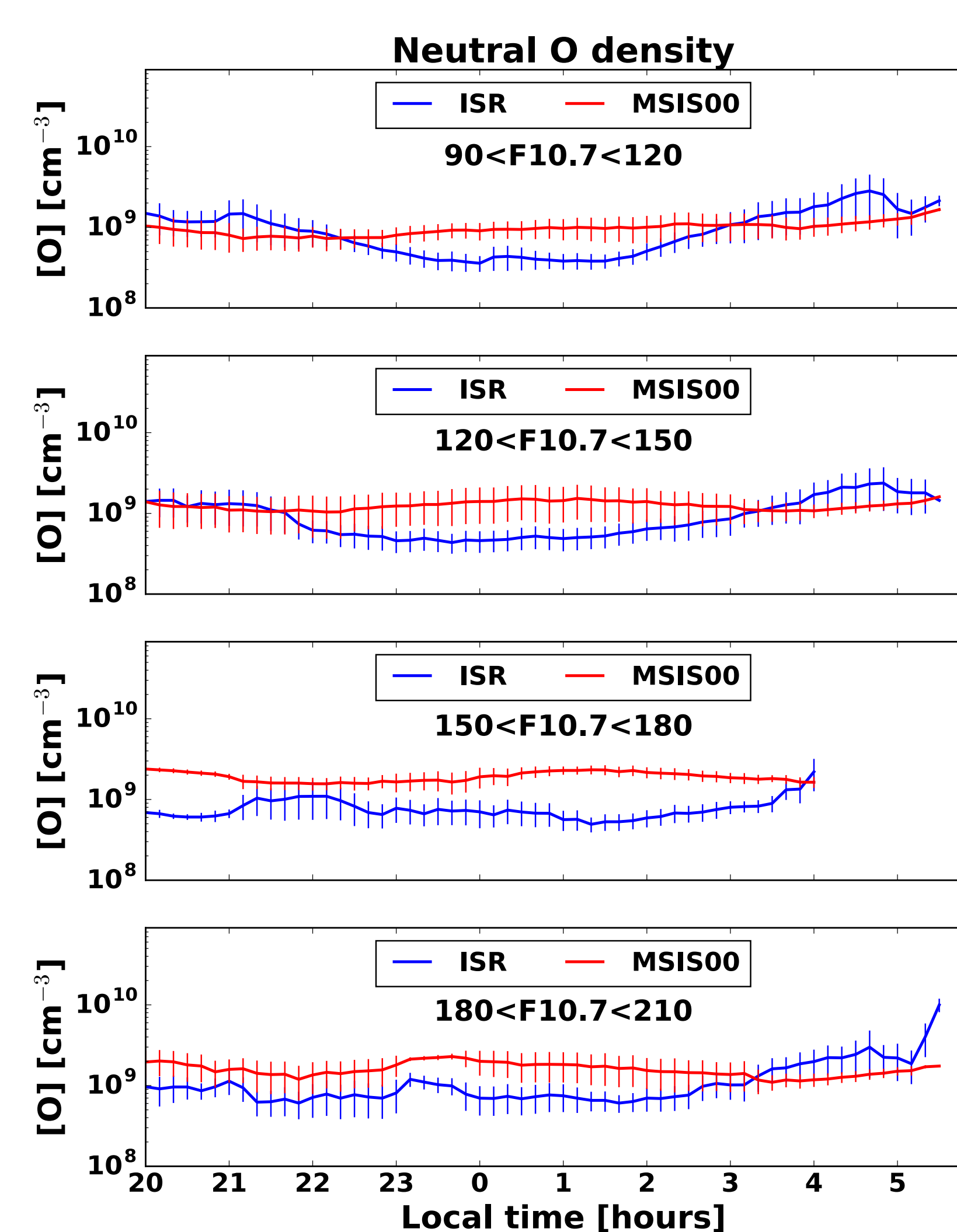


Figure 4. Local time/solar activity variation of the neutral oxygen density [O] specified by MSIS00 (red) and derived from the ISR-based momentum balance analysis (blue) using a Burnside factor  $F=1$ .

In Figures 3 and 4, a systematic bias is observed between the [O] derived from momentum balance and MSIS00 as a function of local time. The resulting variation is consistent over all the seasons as well as solar activity.

MSIS00 is observed to overestimate [O] during the time interval from ~22 LT - 4 LT) and underestimates [O] in the pre-dawn ~4 LT - 6 LT) time interval.

## CONCLUSIONS

The Burnside factor derived via momentum balance analysis of ISR and FPI data, which scales the product of  $Q(O^+-O)$  and [O], is found to vary as a function of local time and solar activity. Similarly, such systematic discrepancies are observed between  $[O]_{MSIS00}$  and [O] derived from momentum balance, assuming a Burnside factor of unity.

These results may arise from errors in the assumed functional form for  $Q(O^+-O)$ , from its miscalculation due to biased MSIS00  $T_n$  specification, from systematic bias in ISR or neutral wind data, or from bias in the specified [O] provided by MSIS00. Potential bias in MSIS00 model output would impact numerous aeronautical investigations as well as the accuracy of numerical modeling of the space-atmosphere interaction region.

Future work will involve the refinement of our momentum balance analysis to include:

- (1) coincident ISR and FPI data, rather than reliance on an empirical wind model
- (2) more accurate plasma velocity data using the dual-beam mode at Arecibo
- (3) coincident measurements of thermospheric O density

These refinements will allow for model-independent derivation of  $Q(O^+-O)$  as well as an unambiguous assessment of potential MSIS00 bias.

## REFERENCES

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