

## INTRODUCTION

The objectives of this work are:

To present a novel technique for estimation of [H] in the topside ionosphere using charge-exchange driven transport of H<sup>+</sup> and O<sup>+</sup> between plasmasphere-ionosphere

To investigate the climatology of the estimated H densities in comparison to NRLMSIS00 specification.

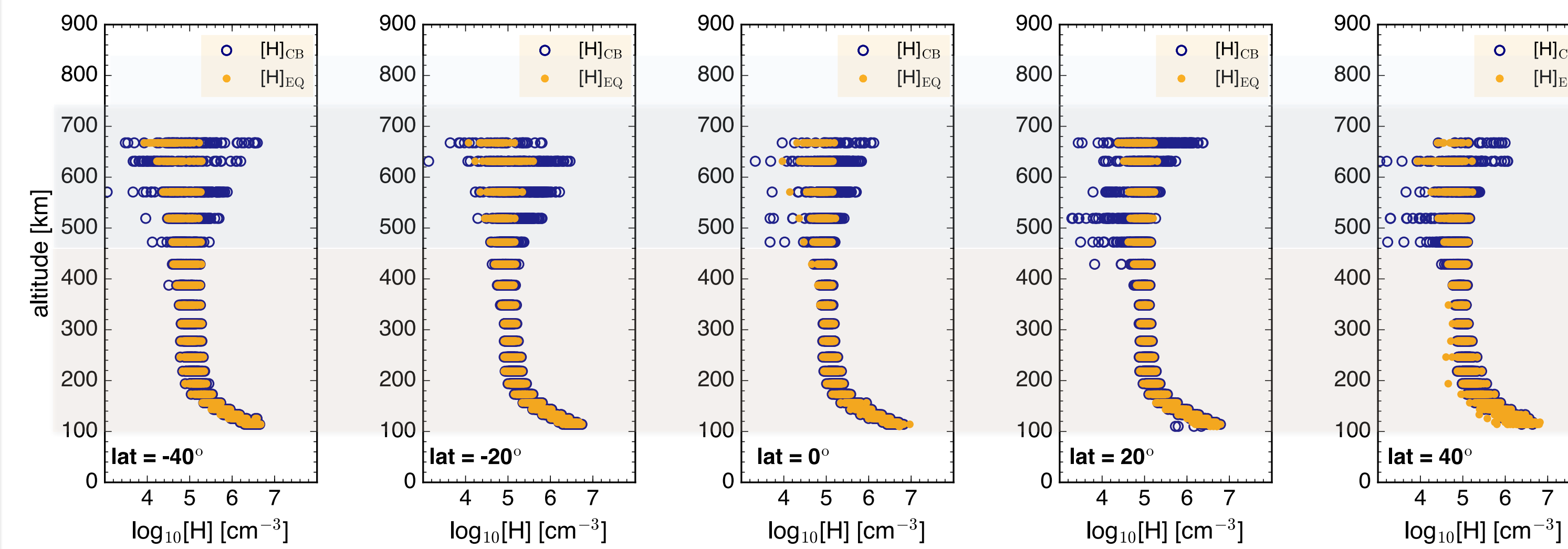
What is the major advantage of this work?

Presents a new technique for [H] estimation with implications for neutral density models.

Motivates several important questions in MIT coupling which can be addressed using upcoming space missions.

## NEUTRAL HYDROGEN DENSITY ESTIMATION

### REGIONS OF VALIDITY



### PROTON TRANSPORT

$$\nabla \cdot \Phi + \frac{\partial [H^+]}{\partial t} \neq 0$$

$$[H]_{CB} = [H]_{EQ} + [H]_{NEQ}$$

### CHARGE-EXCHANGE EQUILIBRIUM

$$\nabla \cdot \Phi + \frac{\partial [H^+]}{\partial t} \approx 0$$

$$[H]_{CB} = [H]_{EQ}$$

### OBSERVATIONS

#### VARIATION OF [H] WITH LATITUDE

Charge-exchange equilibrium is valid from ~110 – 450 km, whereas proton transport becomes increasingly significant at altitudes >450 km.

Ionosphere acts as a proton source for the plasmasphere during the day at altitudes >450 km.

#### VARIATION OF [H] WITH ALTITUDE

The derived [H]<sub>EQ</sub> estimate below ~450 km can be reliably used to:

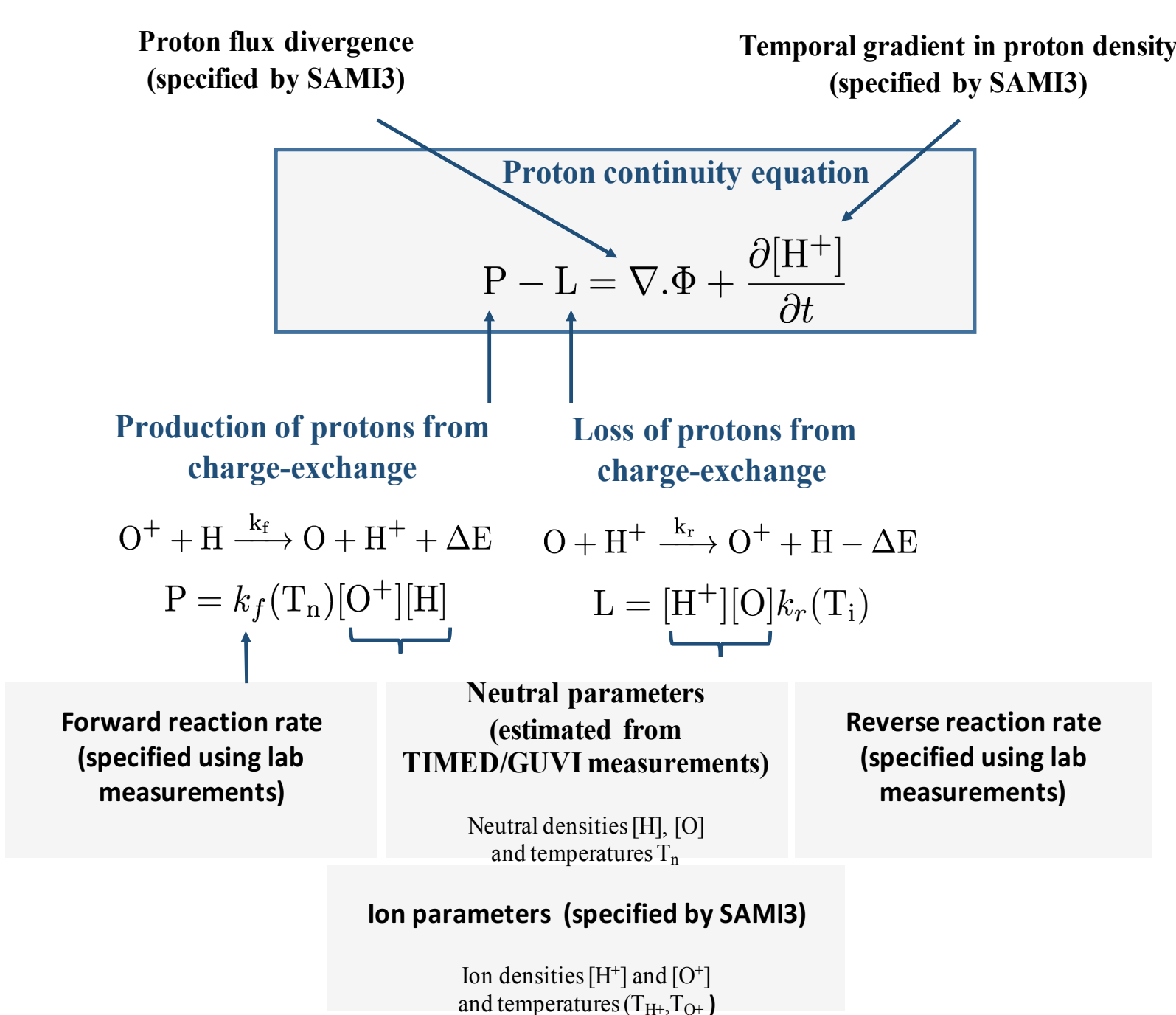
- 1) Evaluate potential bias in neutral atmosphere models such as MSIS00 which derive [H] assuming charge-exchange equilibrium,
- 2) To constraint the inverse-theoretic [H] estimates from radiative transfer models,
- 3) Evaluate thermal Jeans' escape flux at the terrestrial exobase (~400-500 km)

## METHODOLOGY

## PARAMETER SPECIFICATION

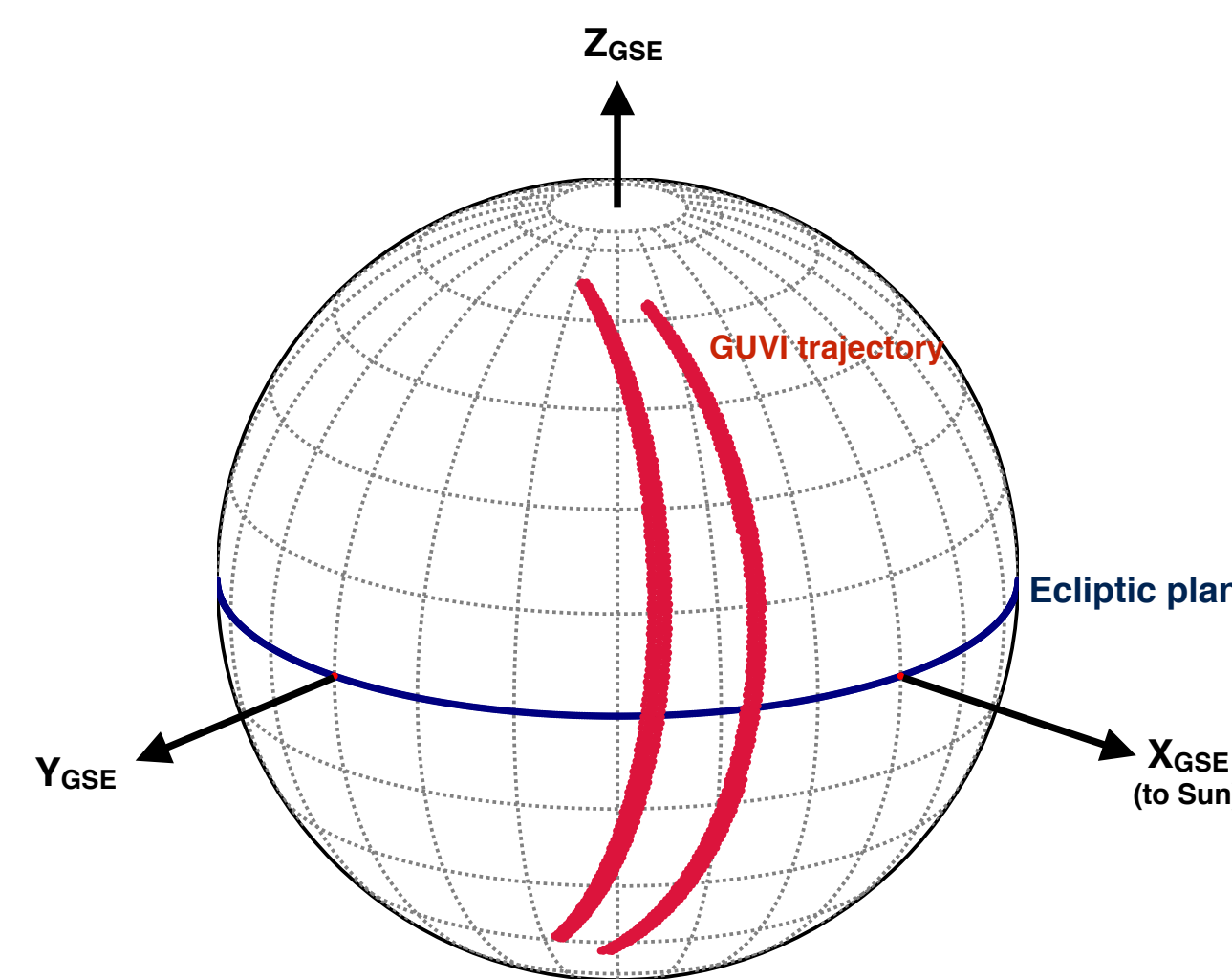
### H<sup>+</sup> CONTINUITY

### TIMED/GUVI



### Thermospheric density and temperature:

Neutral oxygen density [O] and neutral temperatures T<sub>n</sub> estimated for GUVI limb scans using radiative transfer inversion of 135.6 nm emission [Meier et al., 2015].



In one day GUVI has 15 orbits with ~135-150 limb scans over the dayside. TIMED/GUVI trajectory for all 15 orbits is shown (in red) in Figure 1 above.

Each limb scan provides [O] for 21 altitude steps uniformly spaced from 110 - 670 km.

We use height-resolved [O]<sub>GUVI</sub> profiles from 2012 scans spanning 15 orbits on 26 March, 2003.

### SAMI3

#### Ionospheric state parameters:

SAMI3 is a 3-dimensional global ionospheric model that provides the densities, temperature and velocities for seven ion species (H<sup>+</sup>, He<sup>+</sup>, N<sup>+</sup>, O<sup>+</sup>, NO<sup>+</sup>, N<sub>2</sub><sup>+</sup> and O<sub>2</sub><sup>+</sup>) in the altitude range 85 km to 20,000 km [Huba et al., 2000, 2008].

We use co-located estimates of H<sup>+</sup> density [H<sup>+</sup>], temperature (T<sub>H+</sub>) and 3-dimensional velocity (v<sub>H+</sub>) from SAMI3.

### [H] ESTIMATION

$$[H]_{CB} = \frac{k_r(T_i) [H^+]}{k_f(T_n) [O^+]} [O]_{GUVI} + \frac{\nabla \cdot \Phi + \frac{\partial [H^+]}{\partial t}}{k_f(T_n) [O^+]}$$

$$\underbrace{\hspace{10em}}_{[H]_{EQ}} \quad \underbrace{\hspace{10em}}_{[H]_{NEQ}}$$

$$\nabla \cdot \Phi = \underbrace{\frac{\partial (v_{H^+} [H^+])}{\partial z}}_{\text{zenith}} + \underbrace{\frac{\partial (v_{H^+} [H^+])}{\partial u}}_{\text{zonal}} + \underbrace{\frac{\partial (v_{H^+} [H^+])}{\partial v}}_{\text{meridional}}$$

### ANALYSIS INTERVAL

26 March, 2003 (Average Ap = 8 nT, F10.7 = 130 sfu)

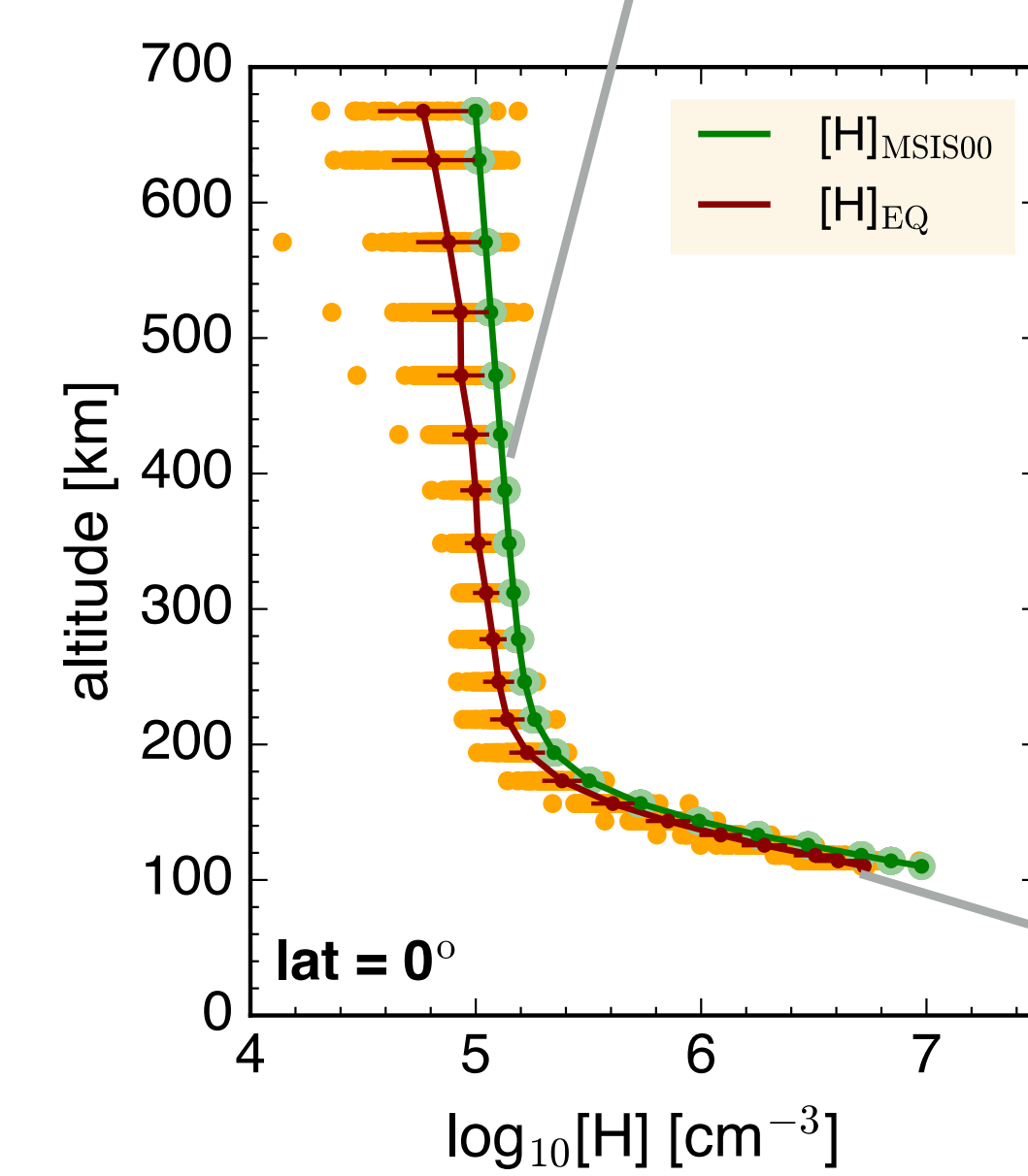
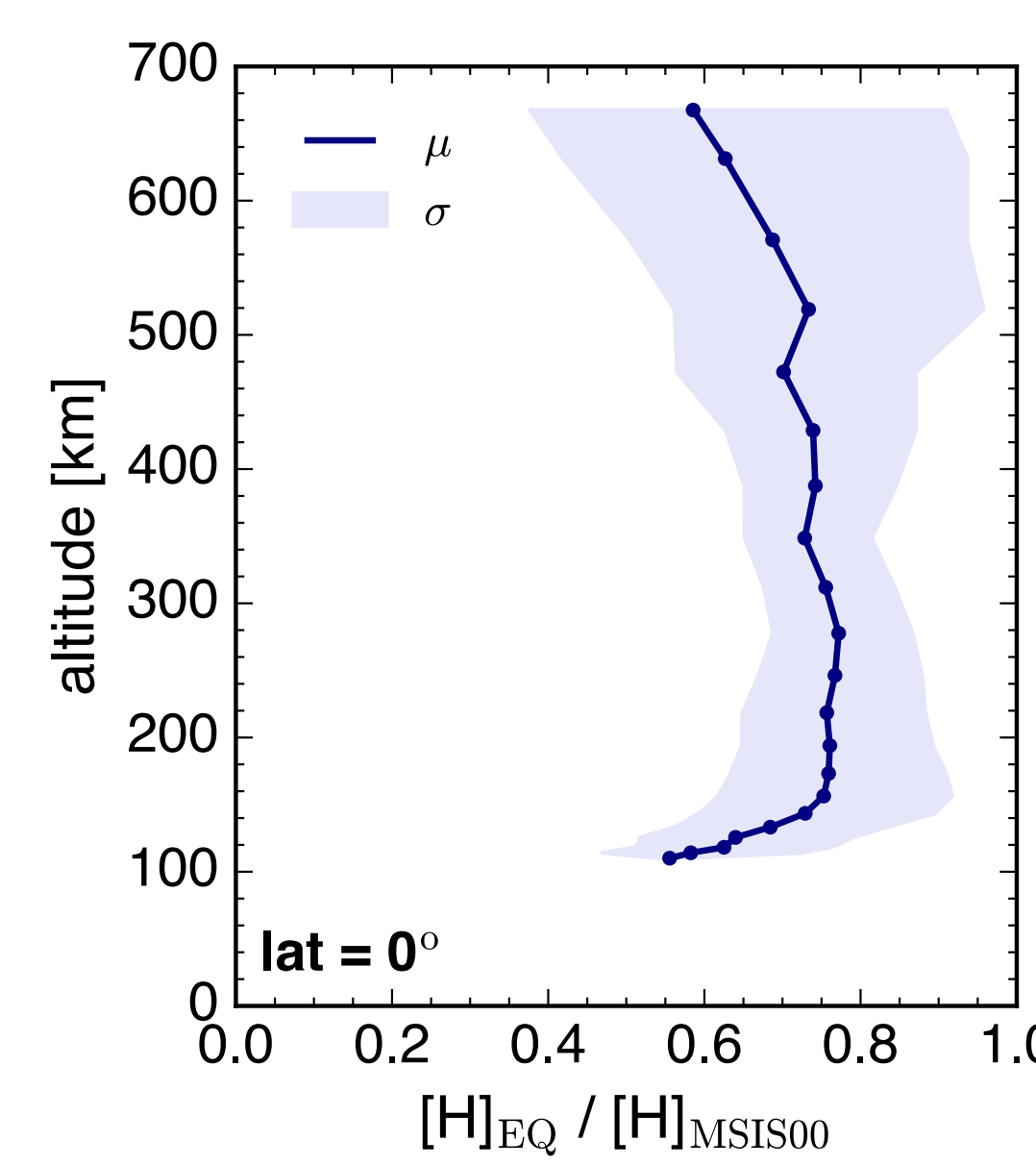
Quite-time dayside interval in solar maximum.

## VARIABILITY OF ESTIMATED [H] AND COMPARISON WITH NRLMSIS00

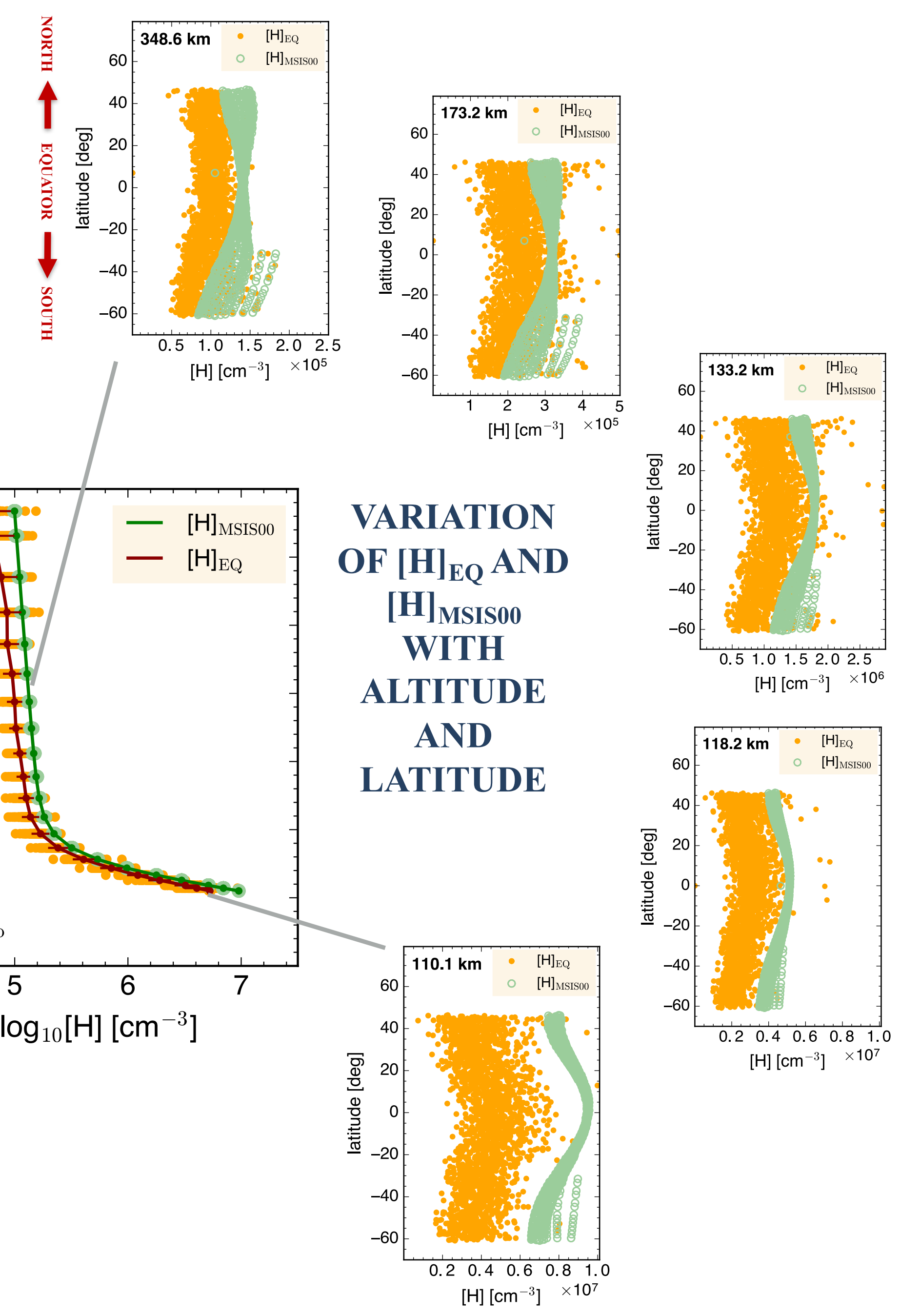
### OBSERVATIONS

SOUTH HEMISPHERE ← EQUATOR → NORTH HEMISPHERE

### POTENTIAL SCALING FACTOR FOR [H]<sub>MSIS00</sub>



### VARIATION OF [H]<sub>EQ</sub> AND [H]<sub>MSIS00</sub> WITH ALTITUDE AND LATITUDE



↑ INCREASING ALTITUDE

#### VARIATION OF [H] WITH LATITUDE

[H]<sub>EQ</sub> and [H]<sub>MSIS00</sub> have the same average climatology and relative magnitude independent of latitude.

#### VARIATION OF [H] WITH ALTITUDE

The difference between [H]<sub>EQ</sub> and [H]<sub>MSIS00</sub> is largest from 100-150 km and stays about constant between 150-450 km.

#### COMPARISON WITH MSIS00

At all latitudes from -40° to 40°, MSIS potentially overestimates [H] during the day by ~20-25% for altitudes ~150-450 km and up to ~40% for MLT altitudes spanning 110-150 km.

#### SOURCES OF UNCERTAINTY

- 1) [O]<sub>GUVI</sub> is underestimated
- 2) SAMI3 underestimates H<sup>+</sup> and/or overestimates O<sup>+</sup>.

## SUMMARY AND CONCLUSIONS

Proposed technique for [H] estimation is very promising and motivates several important questions which can be addressed using upcoming space missions and ionospheric models.

Proposed technique has strong potential to address long-standing discrepancies between models and observations in quantification of plasmasphere-ionosphere transport through model-independent calculations.

1. In-situ mass spectrometer measurements of ion densities and velocities from constellation satellite missions like NASA Dellinger and Exocube would be highly useful for model-independent implementation of [H] estimation.
2. Derivation of neutral hydrogen density [H] from inversion of Lyman-alpha radiances from upcoming mission like NASA ICON and GOLD would allow for MSIS-independent implementation of [H] estimation.
3. Derivation of [H] and/or [O] using self-consistent solution of both, O<sup>+</sup> and H<sup>+</sup> continuity balance would allow for validation of new inversion algorithms for ICON/GOLD, MSIS models as well as impact all IT studies.