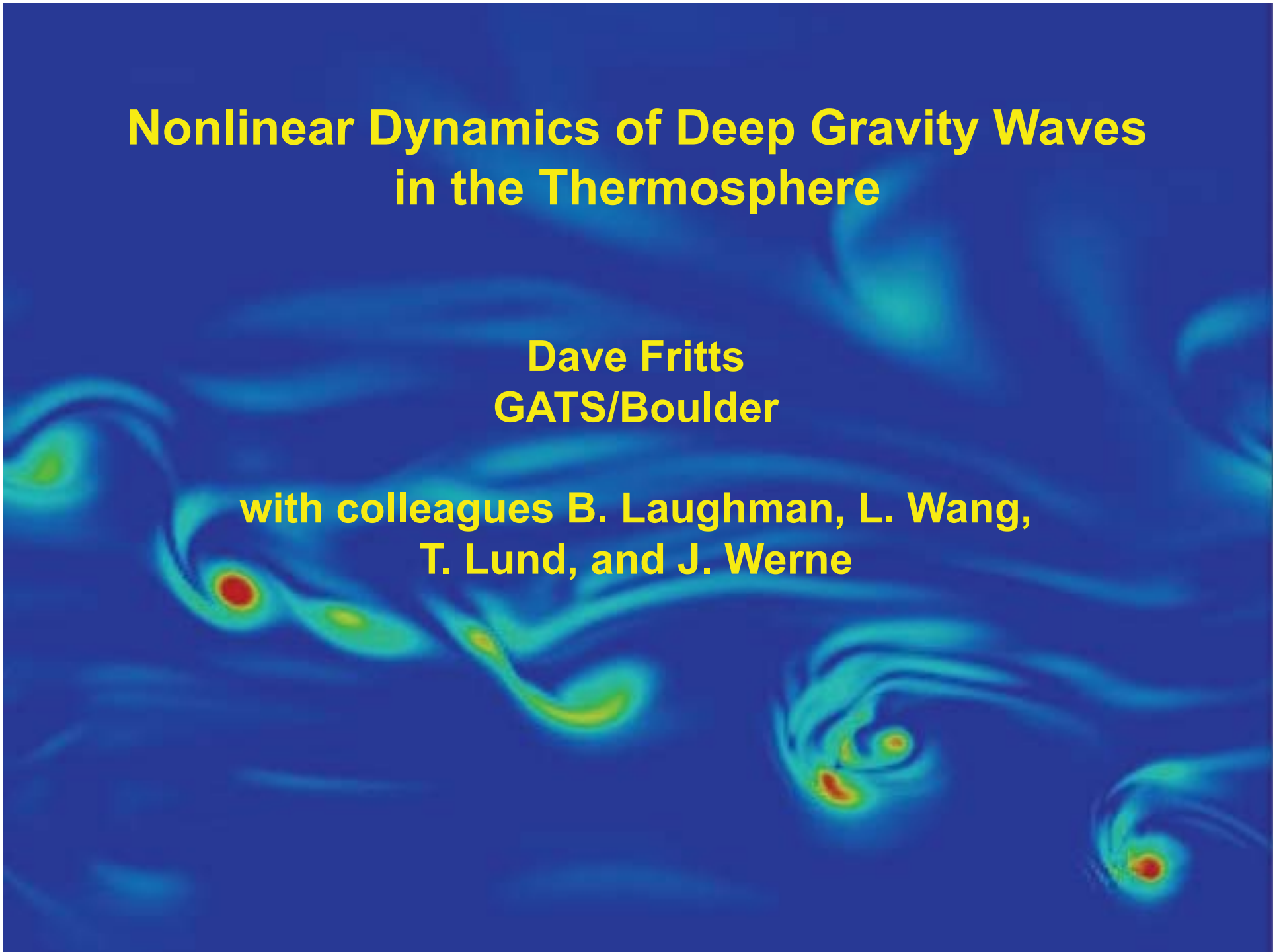


# **Nonlinear Dynamics of Deep Gravity Waves in the Thermosphere**

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# Numerical Models Employed for GW Studies

## 1. Pseudo-spectral model solving Boussinesq Navier-Stokes eqns.

### Capabilities:

- very high resolution (and high Reynolds numbers,  $Re = \lambda_z^2 / T_b \nu$ )
- accurate descriptions of instabilities & turbulence
- valid for  $\lambda_z^2 \ll (4\pi H)^2$   
or  $\lambda_z \sim 10$  km or less

(KH instability for  $Ri = 0.05$ )

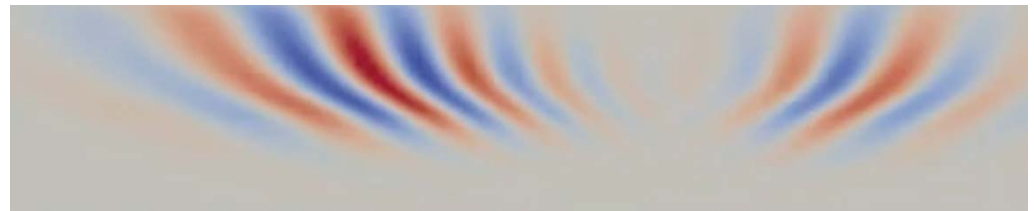
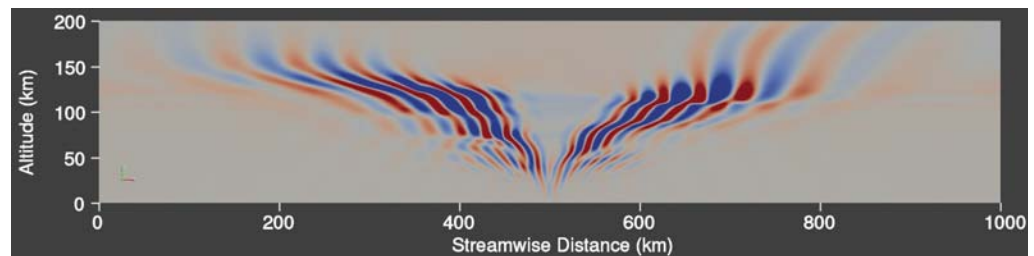


## 2. Finite-volume model solving Compressible or Anelastic Navier-Stokes eqns.

### Capabilities:

- high resolution (and  $Re$ )
- NL dynamics spanning  
 $\sim 15$  to  $20 H$  ( $\sim 0$ - $400$  km)

(GW responses to convective plume and tsunami)



transient response, portion of 3200 x 250 km domain

# Why are Gravity Waves Nonlinear at High Altitudes?

Large-scale and large-amplitude GWs yield

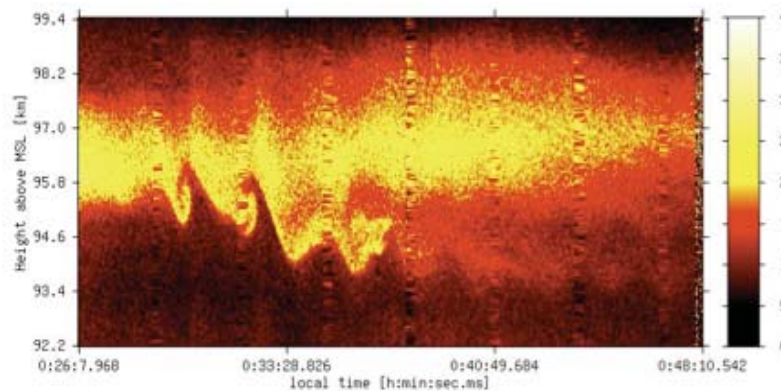
- large Reynolds numbers,  $Re = \lambda_z^2 / T_b \nu$ ,  $\Rightarrow$  flow instabilities
- large momentum transport and effects

$z \sim 80 - 100$  km  $\Rightarrow$   
high  $Re$  ( $10^4 - 10^6$ )  
dynamics  
(strong instabs.)

$z \sim 120 - 160$  km  $\Rightarrow$   
low  $Re$  ( $10^2 - 10^3$ )  
dynamics  
(weaker instabs.)

$z > 200$  km  $\Rightarrow$   
V Low  $Re$  ( $1 - 10^2$ )  
dynamics  
(large amplitudes,  
but no neutral  
instabilities)

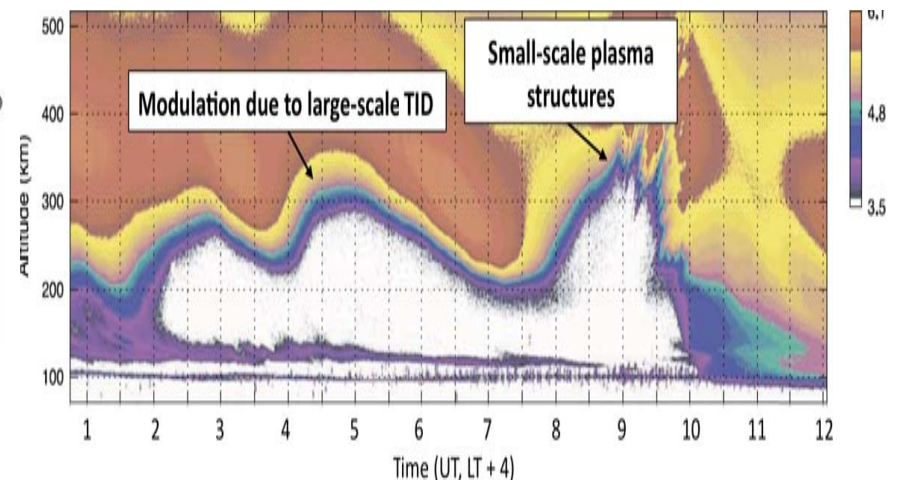
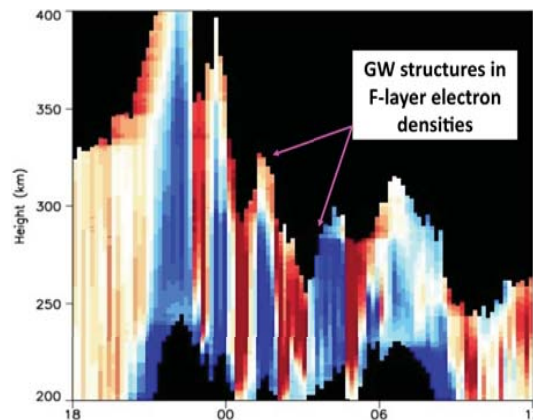
KH instabilities,  $\lambda_h \sim 5-10$  km



GW breaking in NLC,  $\lambda_h \sim 30$  km



conv. GWs in  $N_e$  densities    auroral GWs in the Arecibo ISR  $N_e$  densities



## 3 highlights

- Multi-scale GW dynamics (high  $Re$ ) ~90 km
- GW self acceleration and effects (large GW amps.) ~100 – 140 km
- GW instability in the LT (large GW amps., lower  $Re$ ) ~120 – 160 km

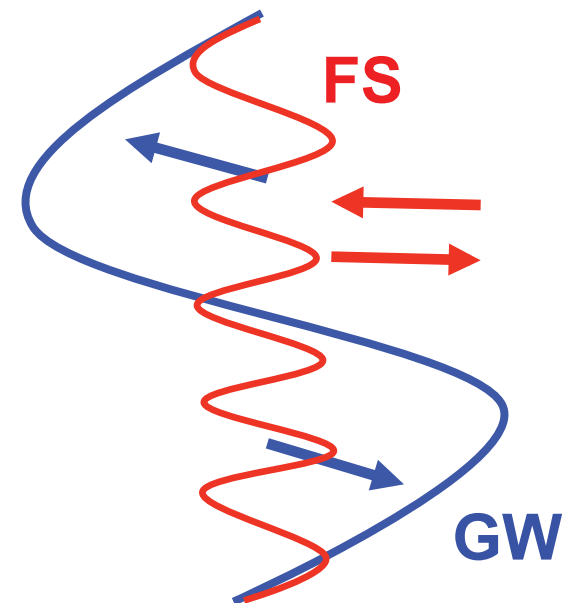
# 1. Multi-Scale GW Dynamics ~80 – 100 km

Superpositions of larger- and smaller-scale waves are ubiquitous  
- these superpositions are largely unexplored at present

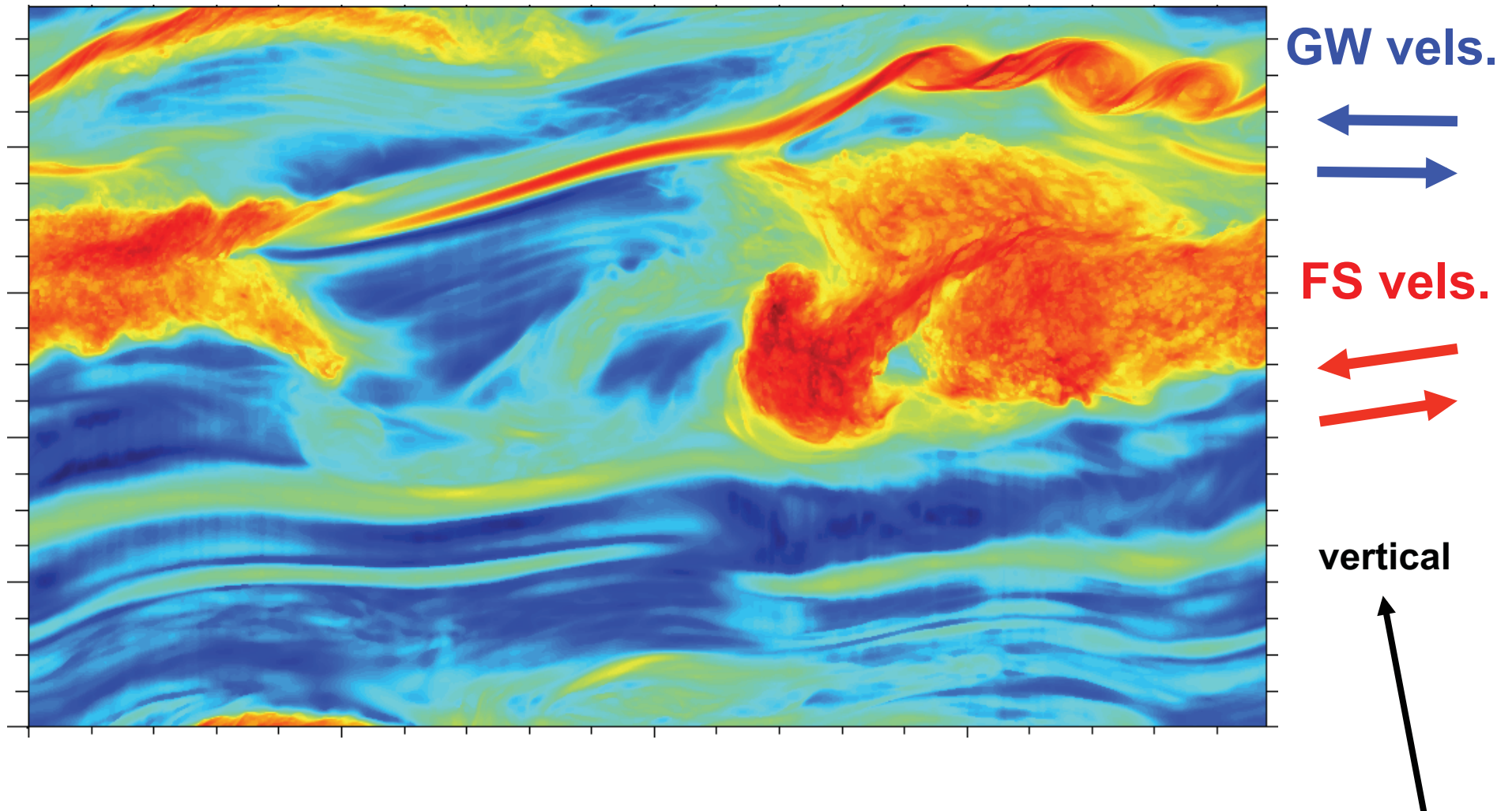
Consider a superposition of

- a “stable” GW ( $a = 1/2$ ,  $Ri_{\min} \gg 1$ ),  
with  $\lambda_{zGW} = 10$  km

- a marginally stable  
fine structure (FS) with  
 $u'_{FS} \sim 2N \sin(m_{FS} z)$ ,  
 $\lambda_{zFS} = 2$  km  
( $Ri_{\min} = 1/4$ )



**GW – fine structure superposition (aligned shears)**  
both GW and fine structure are individually stable  
(red is high vorticity)

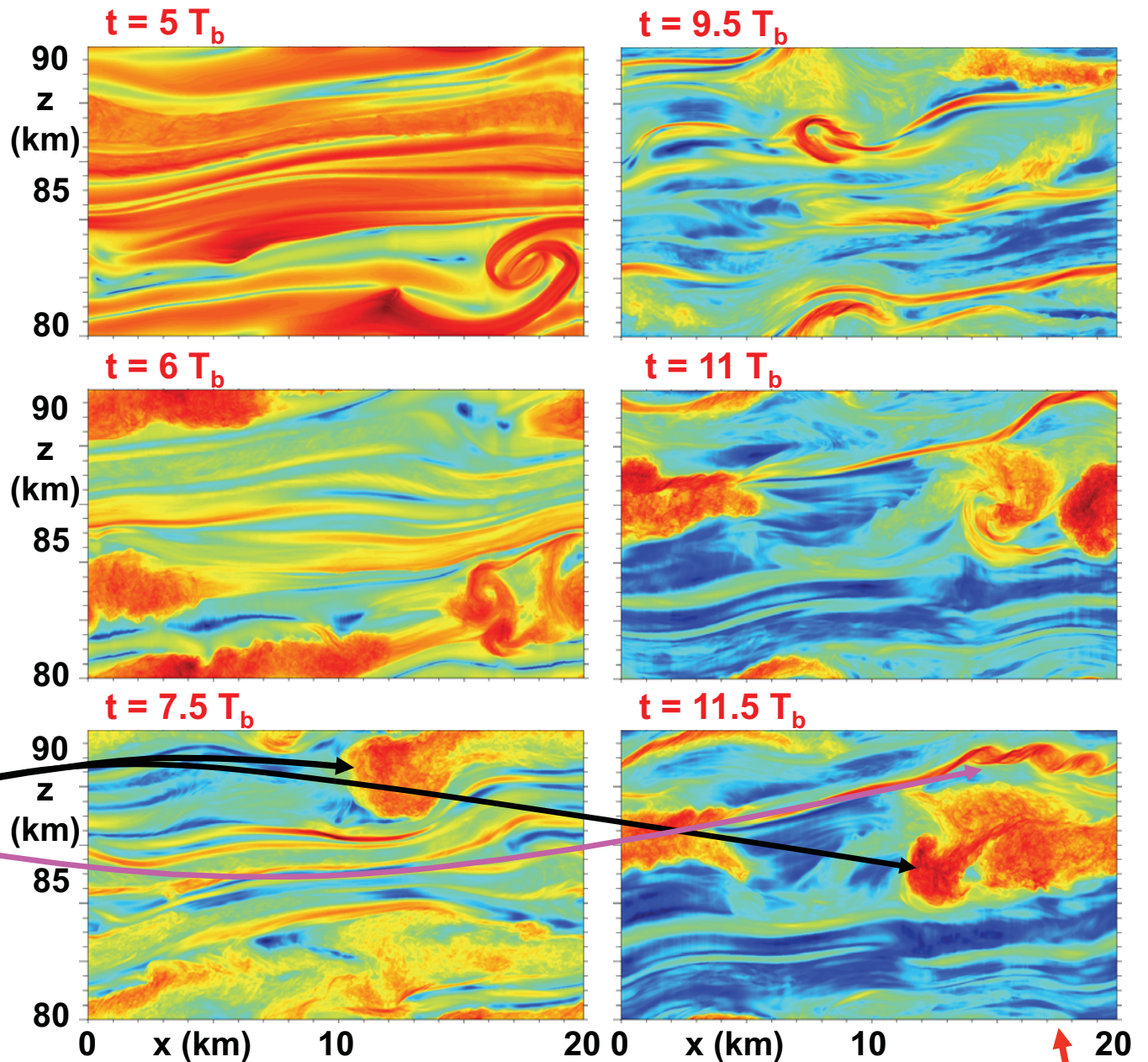


# GW – fine-structure interactions at $z \sim 90$ km ( $Re \sim 50,000$ )

energy dissipation rate ( $\epsilon$ ) in x-z cross sections (along the GW and FS motions)

GW-FS interactions lead to

1. strong wave-wave interactions
2. strong, intermittent, GW breaking, KH instability
3. layering of winds, temperatures, and turbulence

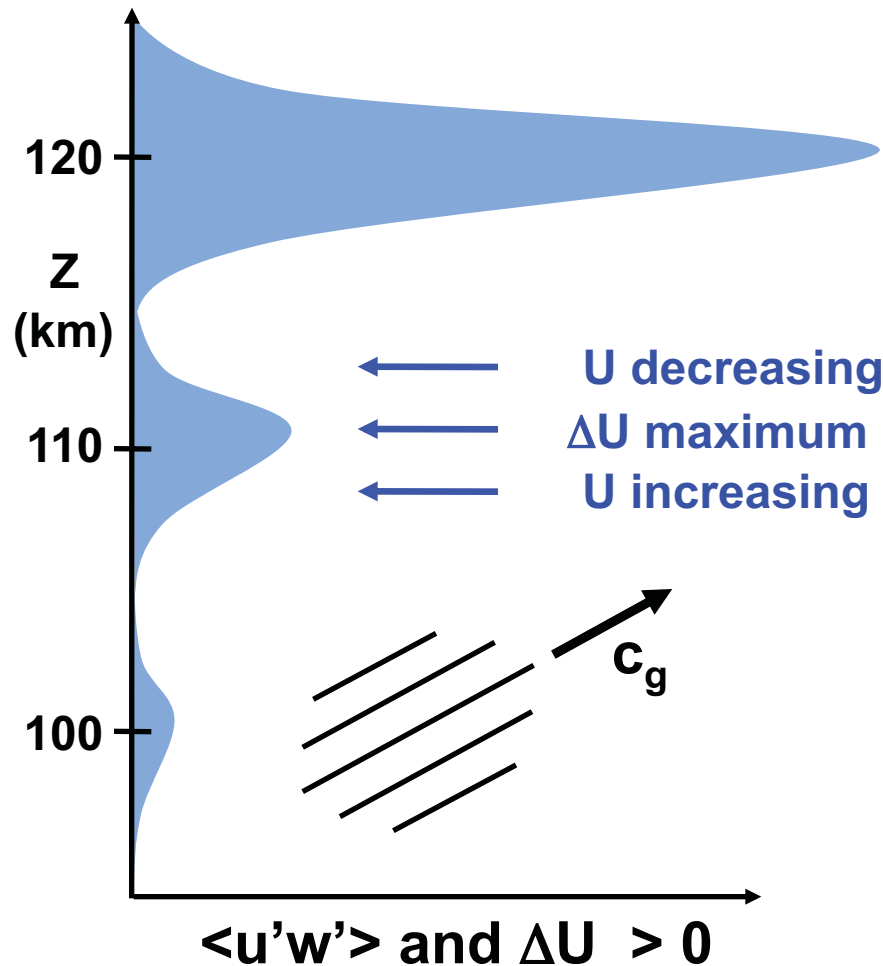


vertical

## 2. GW Self Acceleration and Effects

Decreasing mean density

- =>
- GW amplitudes increase as  $a \sim 1/\rho^{1/2}(z)$
  - $\Delta U \sim 1/\rho(z)$
  - large flow accelerations, altered GW phase structures



=> self-acceleration  
& phase distortion

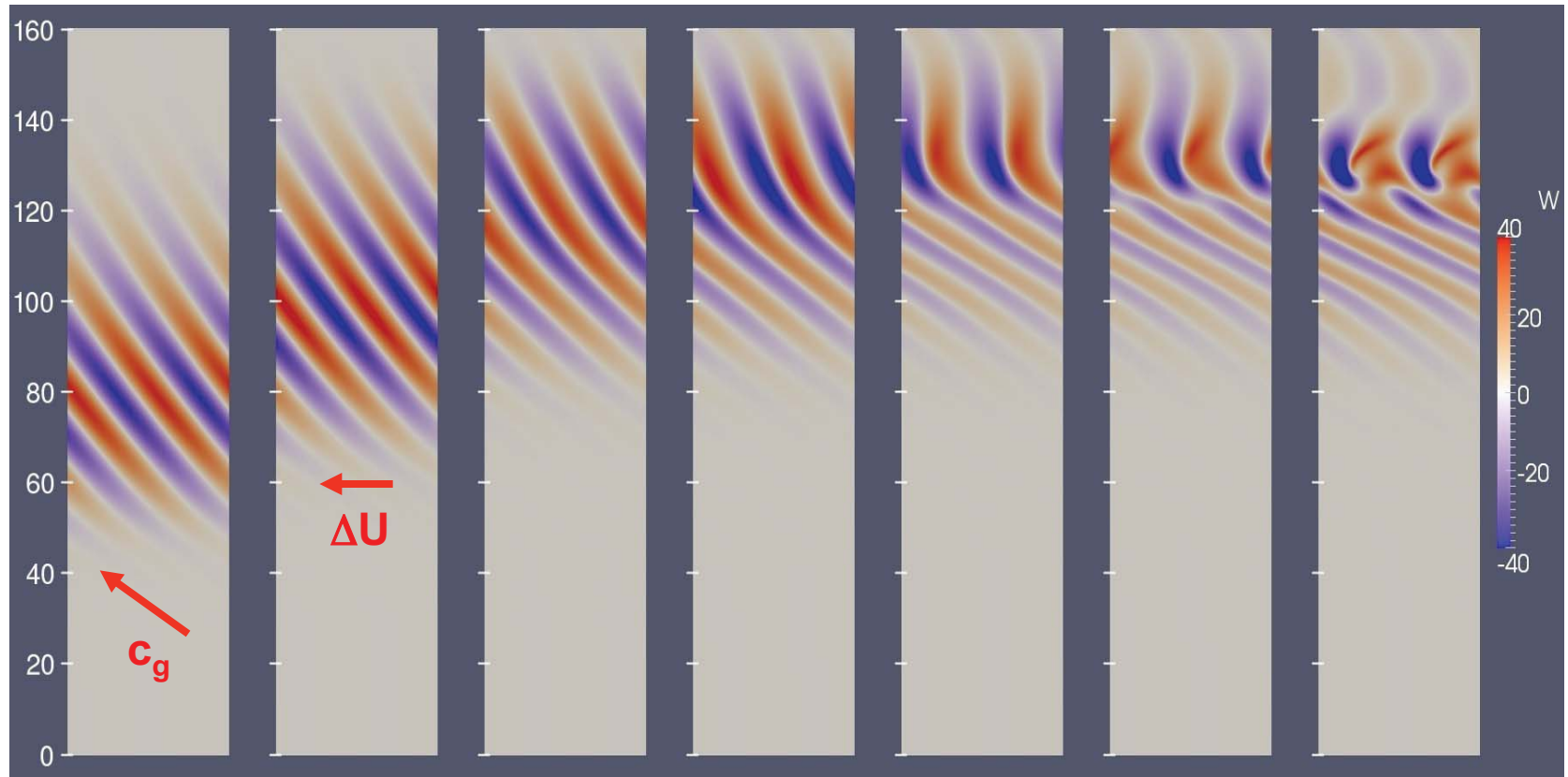
=> leading edge of GW packet  
accelerates

trailing edge decelerates



## “Self-acceleration” of a localized GW packet

- =>
- steepening phase structures at leading edge
  - altered GW group velocities and GW instability



$\Delta t \sim 5$  GW periods

GW with initial  $\lambda_x = \lambda_z = 20$  km,  $N = 0.02$  s $^{-1}$ ,  $\omega = N/1.4$

### 3. GW instability in the TI

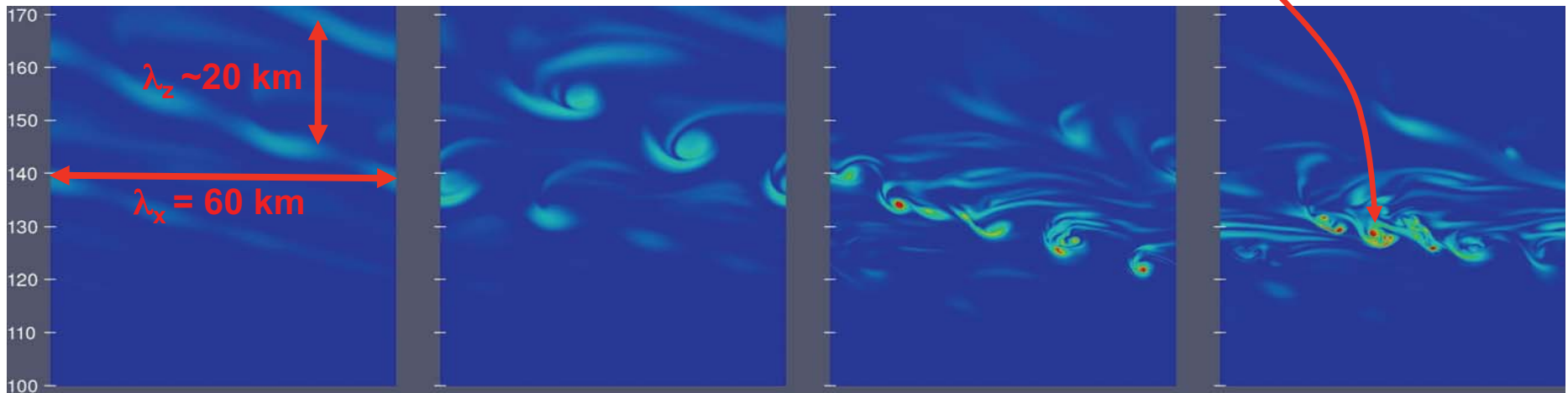
#### large-scale GWs

- penetrate to high altitudes
- achieve large amplitudes
- maintain moderate Reynolds numbers

⇒ **Momentum transport and instabilities**  
**both play large roles**

**GW breaking @  $z \sim 100\text{-}160$  km assuming constant  $U(z)$   
 $u' \sim c-U \sim 50$  m/s,  $\omega \sim N/3$ ,  $Re \sim 10^3$  at  $z \sim 160$  km**

**minimum scales of turbulence  $\sim 1$  km (130 km)**



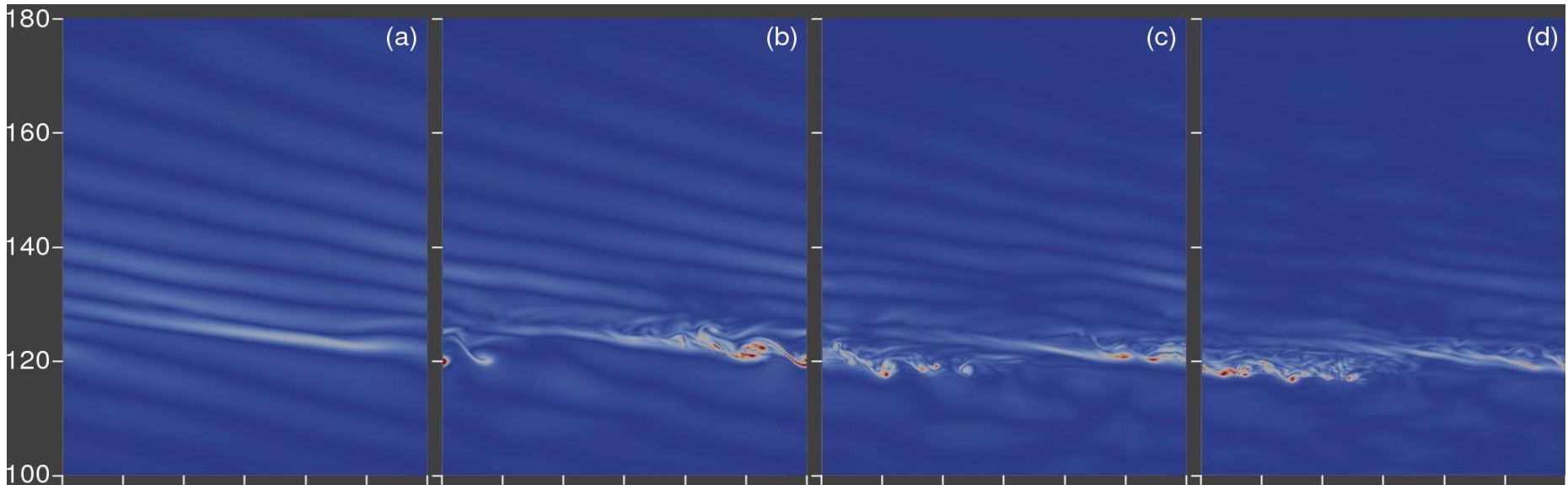
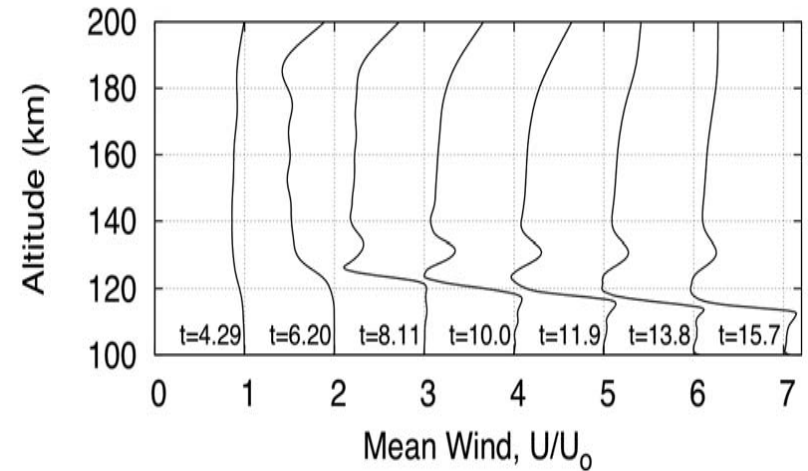
- => - turbulence should be expected intermittently well above  $\sim 100$  km when GW amplitudes and wavelengths are sufficiently large**
- turbulence scales are much larger than we see in the MLT due to the larger viscosity and thermal diffusivity and lower  $Re$**

**Same GW evolution, but now allowing mean flow changes**

**this case exhibits large induced mean winds**

**=> confinement of shears and instabilities to lower altitudes**

**evolving  $U(z,t) \Rightarrow$  accelerations towards GW phase speed**



- **instability character is now much different, due to GW critical level approach**
- **turbulence scales are smaller due to lower altitude, higher  $Re$**

## Summary

Current models are now able to address nonlinear dynamics in the MLT and TI **at realistic scales and Reynolds numbers**

**GW momentum transport, transience, and instabilities all play major roles in the thermosphere**

Multi-scale simulations reveal **complex interactions and impacts on the motion field at large and small scales**

**Momentum transport by large-amplitude GWs induce strong GW self acceleration effects at high altitudes**

Large-scale GWs can **induce instabilities extending well into the lower thermosphere for sufficient GW scales and amplitudes**