# Current Continuity in Auroral System Science: Defining Electron Precipitation

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# Background

- Local auroral coupling of the ionosphere and magnetosphere (MI) is an open area of study (Wolf, 1975; Cowley, 2000; Lotko 2004; Amm et al. 2008).
- MI coupling demands self-consistent, topside maps of field-aligned current (FAC) and  $\mathbf{E} \times \mathbf{B}$  plasma flow that agree with ionospheric conductivity patterns created by charged particle auroral precipitation.
- Discrete auroral precipitation provided by the auroral acceleration region creates arc-scale morphology in the ionospheric conductivity volume to which the MI coupling is highly sensitive.
- Quasi-static ionospheric plasma flow, FAC, and conductivity have a 2D topside relation given by Eq. 6.12 in Kelley (2009):
  - $j_{\parallel}(x, y) = \Sigma_P \nabla \cdot \mathbf{E} + \mathbf{E} \cdot \nabla \Sigma_P + (\mathbf{E} \times \mathbf{b}) \cdot \nabla \Sigma_H$ (1)

where  $j_{\parallel}$  is a horizontal map of FAC at the topside ionosphere,  $\Sigma_P$  and  $\Sigma_H$  are the height-integrated Pedersen and Hall conductivities, and E is the perpendicular ionospheric electric field

• This 2D picture can significantly hide the 3D nature of auroral current closure. • Due to the highly sensitive nature of auroral current closure to the 3D conductivity volume, how electron precipitation is defined is crucial.

### **Problem Statement**

This work focuses on an important part of auroral arc systems: impact ionization imparted by electron precipitation. This flux is often assumed to be a simple unaccelerated Maxwellian. However, for discrete aurora, this flux can be a superposition of a primary accelerated Maxwellian population along with secondary, low energy population of re-accelerated backscatter (Evans, 1974). We investigate the bearing this has on auroral current closure by comparing unaccelerated vs. accelerated Maxwellian electron precipitation, both with and without the backscatter low-energy tail.

# Ways to Define Electron Precipitation

• We look at three parallel differential electron number fluxes (see Figure 1): 1. An u

unaccelerated Maxwellian:  
$$J_{\parallel}(E; E_0) = \frac{Q_0}{2E^2} \frac{E}{E} \exp\left(-\frac{E}{E}\right)$$

 $E_0 = 2 \text{ keV}$ 

Maxwellian

 $ZE_0^- E_0$ where  $Q_0$  is the total energy flux,  $Q_0 \equiv \int_0^\infty E J_{\parallel}(E) dE$ , and  $E_0$  is the characteristic energy.

2. An accelerated Maxwellian:

$$I_{\parallel}(E;T_s,\boldsymbol{U_d}) = \begin{cases} \frac{Q_0}{T_s^2 + (T_s + \boldsymbol{U_d})^2} \frac{E}{T_s} \exp\left(-\frac{E - \boldsymbol{U_d}}{T_s}\right), & E \ge \boldsymbol{U_d} \\ 0, & \text{else} \end{cases}$$
(3)

where  $T_s$  is the characteristic energy of the source region and  $U_d$  is the acceleration region potential drop (Kaeppler, 2013).

3. A combination of a primary accelerated Maxwellian from the source region and a low-energy tail (LET) from re-accelerated backscattered electrons of ionospheric origin (Evans, 1974).



### From Precipitation Spectra to Impact Ionization

- Calculations of altitude dependent impact ionization rates are done using two methods by Fang et al. (2008) and by Fang et al. (2010):
- Method A: Fang et al. (2008) parameterize with isotropic, unaccelerated precipitation using both multistream (Lummerzheim & Lilensten, 1994) and two-stream (Solomon & Abreu, 1989) transport models.
- Method B: Fang et al. (2010) do the same but with isotropic, monoenergetic precipitation which allows for building up arbitrary precipitation spectra using multiple of their calculations (see Figure 2).



 $E_0 = 2 \text{ keV}$  **Figure 2**: A comparison of  $E_0 = 4 \text{ keV}$  impact ionization rates using methods A and B. Top: Three unaccelerated Maxwellian spectra given by Eq. (2) with characteristic energies,  $E_0$ Lines indicate the full spectra used in method A and points show what energies,  $E_{mono}$ , are used in composing the same spectra for method B. Bottom: Ionization rates vs. calculated using altitude method A (dashed) and method B (solid). This uses an MSISE-90 (Hedin, 1991) atmosphere at 10 UT on Feb 1, 2015, at 65.8° N and 207.7° *E*, with daily and 3-month-avg. F10.7 cm radio emissions of 143.7 and 137.6, and a daily Ap-index of 20.

# Accelerated vs. Unaccelerated Impact Ionization

- Figure 3 shows how impact ionization altitude profiles change when comparing unaccelerated vs. accelerated Maxwellians.
- Defining an "X keV" auroral arc with either  $E_0$  or  $U_d$  can change peak ionization rate altitudes by 10-20 kilometers.
- Depending on the source characteristic energy,  $T_s$ , an accelerated "4 keV" arc
- can vary peak ionization by 10-20 km in altitude. • Only when  $E_0 = U_d = T_s$  is an unaccelerated spectrum comparable to an accelerated one



**Figure 3: Left:** Unaccelerated (dashed) vs. accelerated with  $T_s = 800 \text{ eV}$  (solid) spectra and resulting ionization rate altitude profiles for different peak flux energies. Right: Unaccelerated with  $E_0 = 4 \text{ keV}$  (dashed) vs. accelerated with  $U_d = 4 \text{ keV}$  (solid) for different source characteristic energies.

### Implementing Arbitrary Spectra into GEMINI

- To look at auroral current closure, we use multi-fluid model runs provided by GEMINI (Zettergren & Semeter, 2012; Zettergren & Snively, 2019). For details see github.com/gemini3d.
- This model can simulate the ionosphere at auroral arc scales (see Figure 6). • GEMINI solves for static current continuity to account for changes in model parameters impacting conductivities as it steps forward in time.
- The model is forced with a 2D topside map of FAC or flow (see Figure 7). • Additionally, in its present version, it is driven with 2D topside electron precipitation maps of  $Q_0$  and  $E_0$  covering impact ionization via method A.
- We implemented <u>method B</u> into GEMINI and tested it (see Figure 4).
- Figure 5 shows GEMINI results after running the three spectra in Figure 1.



Figure 4: Example test of the implementation into GEMINI of method B (solid) against the original method A (dashed). These tests are done at 4 orders of magnitude in both  $E_0$ and  $Q_0$  ( $Q_0 = 0.1, 10, 100 \text{ mW}$ )  $m^2$  not shown).

Figure 5: GEMINI results for a 2D (alt., lat.) simulation with a line of constant topside precipitation as defined by the spectra in Figure 1 using method B. This simulation has no FAC or plasma flow drivers. **Top:** Electron density altitude profile for the three different spectra. Bottom: Hall (solid) Pedersen (dashed) and conductivity altitude profiles.



Figure 6: The general context of this work and the GEMINI model space.

Circle Comments

East [km] Figure 7: Typical topside 2D input maps for a GEMINI simulation.

-1000 -500

east [km]



500 1000 east [km]

-1000 -500

0





unaccelerated,

Zettergren, M. & Semeter J., 2012, J. of Geophys. Research, 10.1029/2012JA017637 Zettergren, M. & Snively J., 2019, J. of Geophys. Research, 10.1029/2018GL081569