

Volcano-generated Ionospheric Disturbances: Comparison of GITM-R Simulations with GNSS Observation



COUP-10

Justin J. Tyska¹, Yue Deng¹, Cissi Y. Lin², Shunrong Zhang³

¹University of Texas at Arlington, USA, ²National Central University, Taiwan, Massachusetts Institute of Technology, USA

Introduction

It has been known for some time that explosive events can generate acoustic-gravity waves (AGWs) that then propagate by virtue of the background density profile to thermospheric heights and influence the ionosphere in a way detectable by dual frequency Global Navigation Satellite Systems (GNSS) through the measurement of Total electron content (TEC).

Volcanic Ionospheric Disturbances (CVIDs) observed following main eruption phases typically follow one of two types, thought to be indicative of the eruption dynamics

Type 1 (T1)

- Vulcanian eruption style: Sudden, intense, explosion
- Shock-Acoustic Dominate
- "N-Shaped" TEC waveforms
- 8-11 min arrival times to IT

Type 2 (T2)

- Plinian/Sub-Plinian eruption style: Continuous/multiple explosions
- AGW/GW dominate
- Quasi-Periodic TEC waveforms
- 11-50 min arrival times to IT

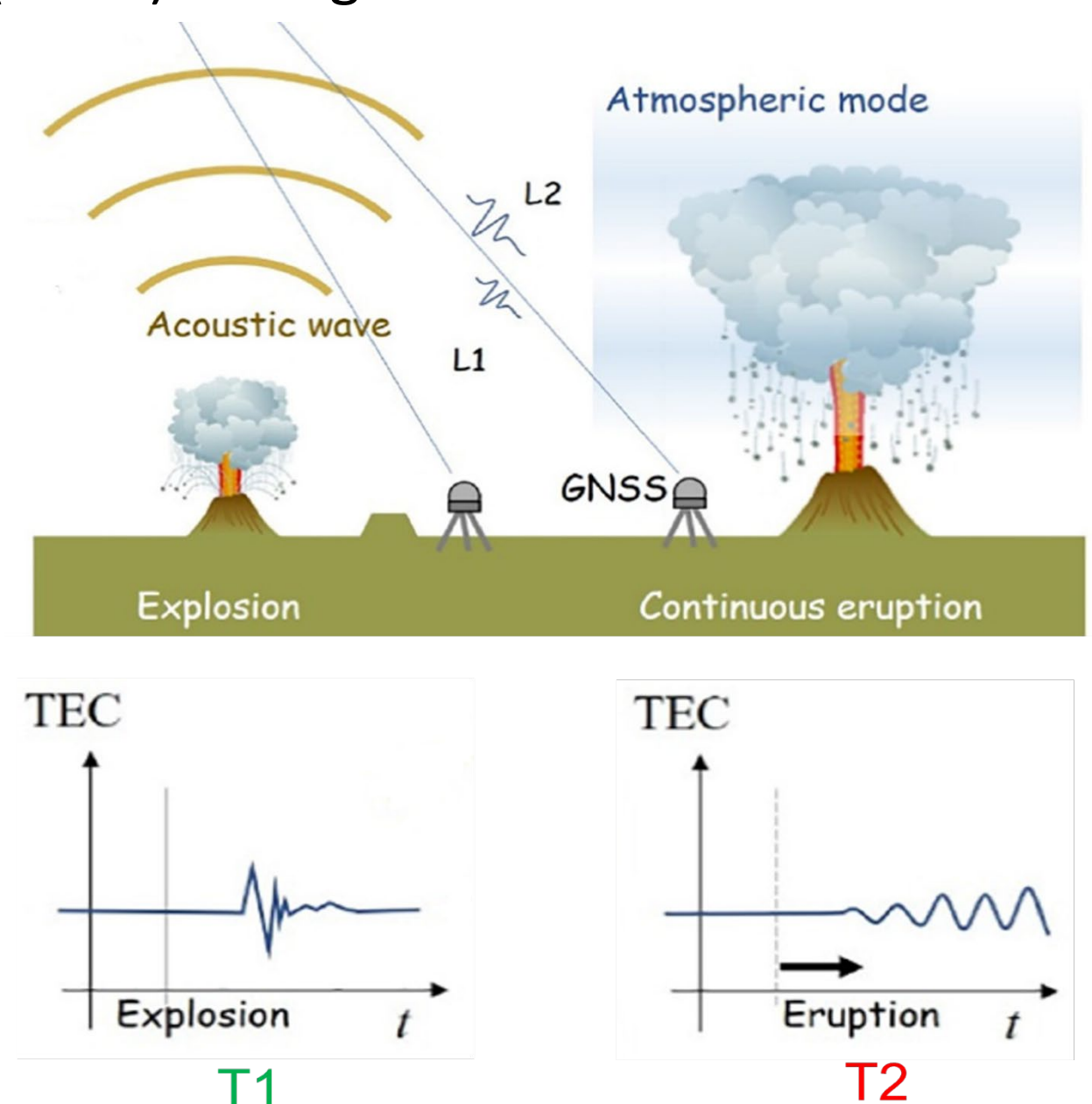


Fig 1. Schematic of Atmospheric response to different eruption styles (re-rendered)[1]

Motivation

- There are a multitude of observational studies concerning ionospheric disturbances induced by volcanic eruptions, but little has been done to recreate such events through simulation, especially in a global circulation model (GCM).
- We believe Data-model comparisons can provide valuable in-depth analysis that can not only be used to access and improve the performance of models but also be used to better understand physical process in complex case studies.

The focus of this study is to simulate ionospheric Total Electron Content (TEC) variations induced by the 2015 Calbuco eruption using the Global Ionosphere-Thermosphere model (GITM) and subsequently compare these simulations with GNSS data of the first eruption phase. Additionally, we investigate the difference in simulated TEC variations induced by different primary AGW forcing mechanisms.

Calbuco 2015 Event and GNSS data (Phase 1)

General event info [2]

- Location:** Southern Chile (~41.3 S, 72.6 W)
- Vent:** ~2 km above sea level
- Eruption Type:** Plinian (continuous)
- Time:** April 22nd, 21:04 UT (18:04 LT)
- Volcanic Explosivity Index:** 4
- Plume Height:** ~15 km

GNSS data [3]

Duration: 1.5 Hours

Form:

- Quasi-Periodic (T2)
- multiple wave packets (near)
- Gravity Mode (Far)
- Dominant Modes:**
- 4.8, 5.2 mHz (near)
- 1.0 mHz (Far)
- Magnitude :**
- ~0.6/0.25 TECU (max)
- ~0.45 TECU (mean)
- Apparent Phase Speeds:**
- 900 m/s (Acoustic packets)
- 260 m/s (Gravity packet)

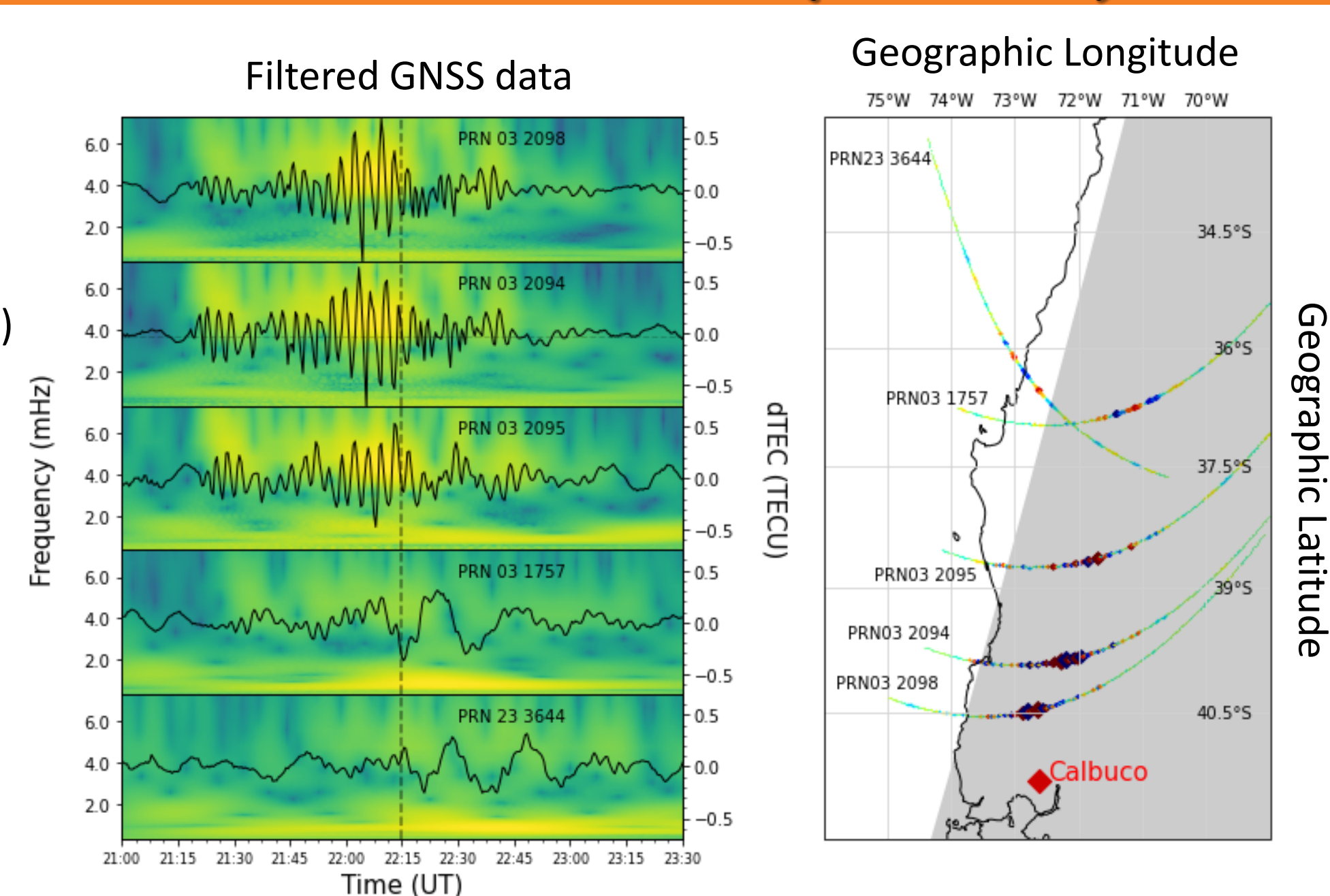


Fig2. Example of GNSS data showing TEC variations caused by Calbuco Eruption

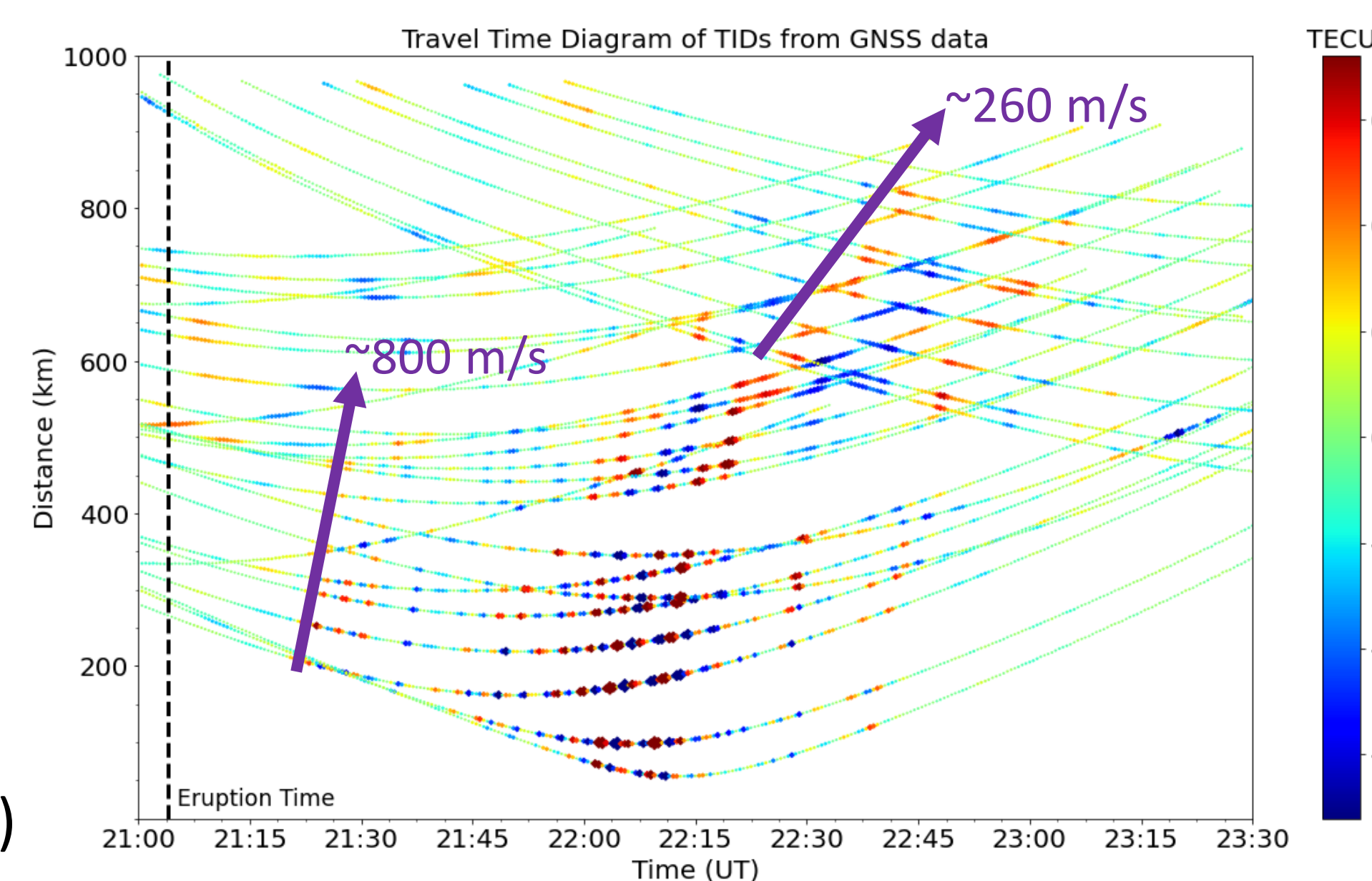


Fig 3. Travel time diagram of GNSS observations for first phase of Calbuco Event

Volcano Model and Propagation to 100 km

- Point source pressure oscillation at local acoustic cut-off frequency ($\omega_a \sim 2.9$ mHz, $A = 6.0e6$) [4]

$$p'(t) = A e^{-0.5H\bar{p}^1} \left\{ \delta(t - t_0) - \frac{\omega_a t_0 / H}{(\bar{t}^2 - t_0^2)^{1/2}} H(t - t_0) \right\}$$

- Assume a mass injection rate:

$$F_M(t) = B(t - t_0) e^{-\frac{(t-t_0)^2}{2\sigma_t^2}} \quad B = 5000 \quad \sigma_t = 58.5$$

- Simplified Volcano Forcing Function:

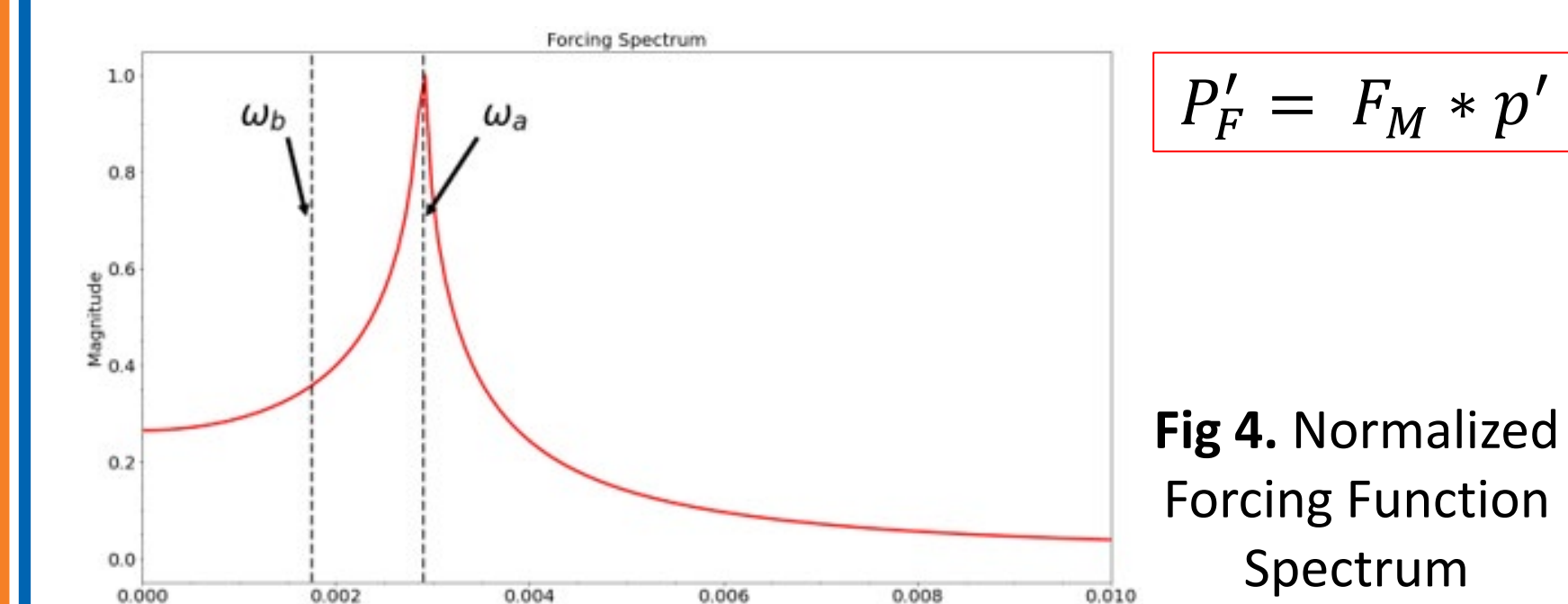


Fig 4. Normalized Forcing Function Spectrum

- Spherical Propagation to GITM-R lower boundary [5]

$$w = \left(\frac{\rho_{0km}}{\rho_{100km}} \right)^{1/2} \frac{1}{2\pi r_e} \int G_p^w \bar{P} e^{i(kr - \omega t + \phi)} d\omega$$

Direct Propagation (DP) [5]

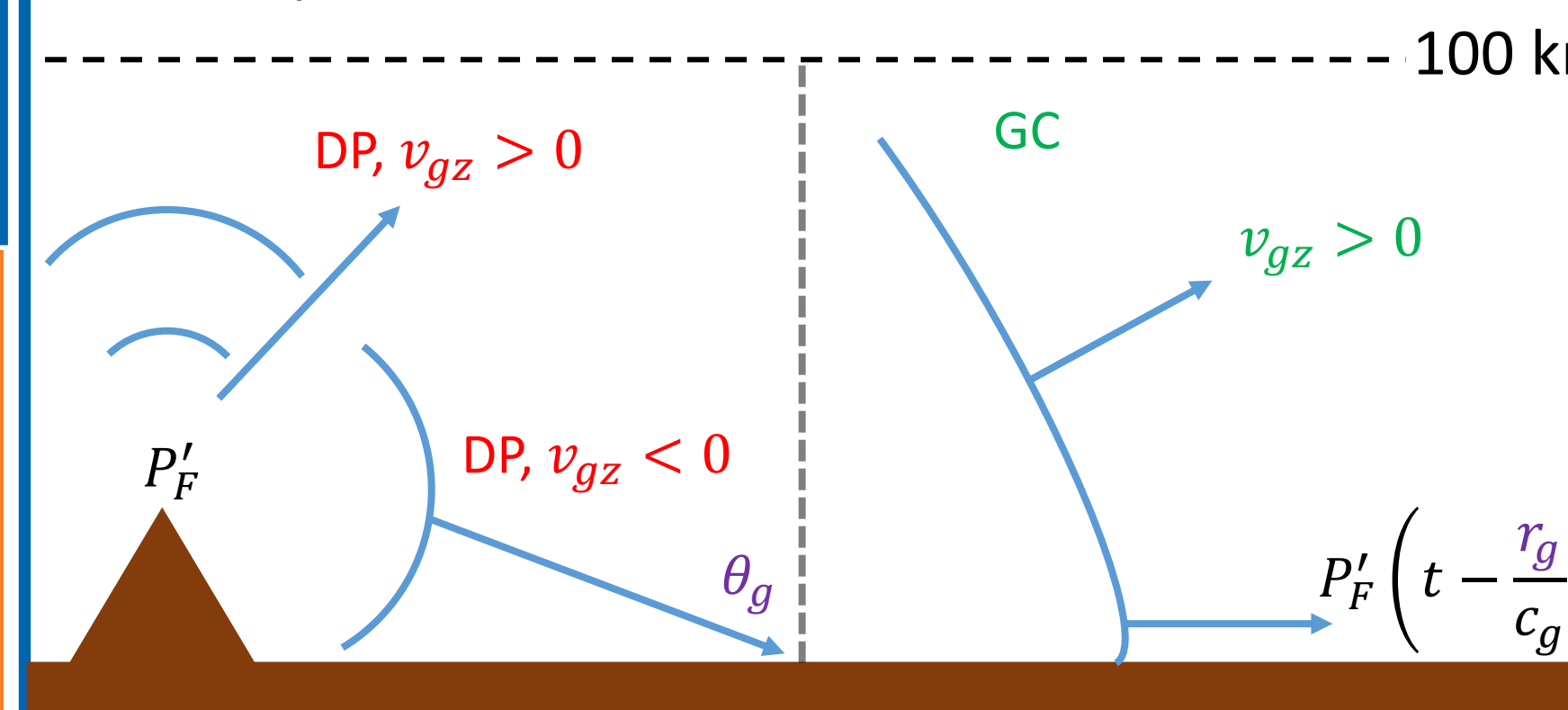
$$k^2 = m^2 \tan^2 \theta$$

$$m^2 = \frac{\omega^2 - \bar{\omega}_a^2}{c^2} - k^2 \frac{\omega^2 - \bar{\omega}_b^2}{\omega^2} \quad \phi = 0$$

Ground Coupled (GC) [6]

$$v_s = \frac{c_g}{\sin \theta_g}$$

$$m^2 = \frac{\omega^2 - \bar{\omega}_a^2}{c^2} - \frac{\omega^2 - \bar{\omega}_b^2}{v_s^2} \quad \phi = (k - k')r_g - m z_s$$



Global Ionosphere Thermosphere Model w/ Local Mesh Refinement [7]

- GCM that models the Earth's IT system by solving Navier-Stokes for Neutrals and Simplified MHD for the plasmas
- Self consistently solves for constituent densities, temperatures, and dynamics

Why GITM-R for Meso-scale TAD-TID dynamics?

- Allows for non-hydrostatic solutions
- **Local Mesh Refinement feature**
- Layered patches of increased resolution, imbedded and coupled together.
- realistic specification of regional boundaries

Simulation Set-up

- 4 Hour Simulation time (20:30 – 24:30 UT)
- Three imbedded layers
- **Regional Layer 2**
Lon: [55° W, 90° W]
Lat: [34° S, 59° S]
Resolution: 0.1° x 0.1°
- Specification of Ionosphere state
- 8 hr pre-run with solar wind drivers (B, v, N_p, T)
- F10.7 of 150

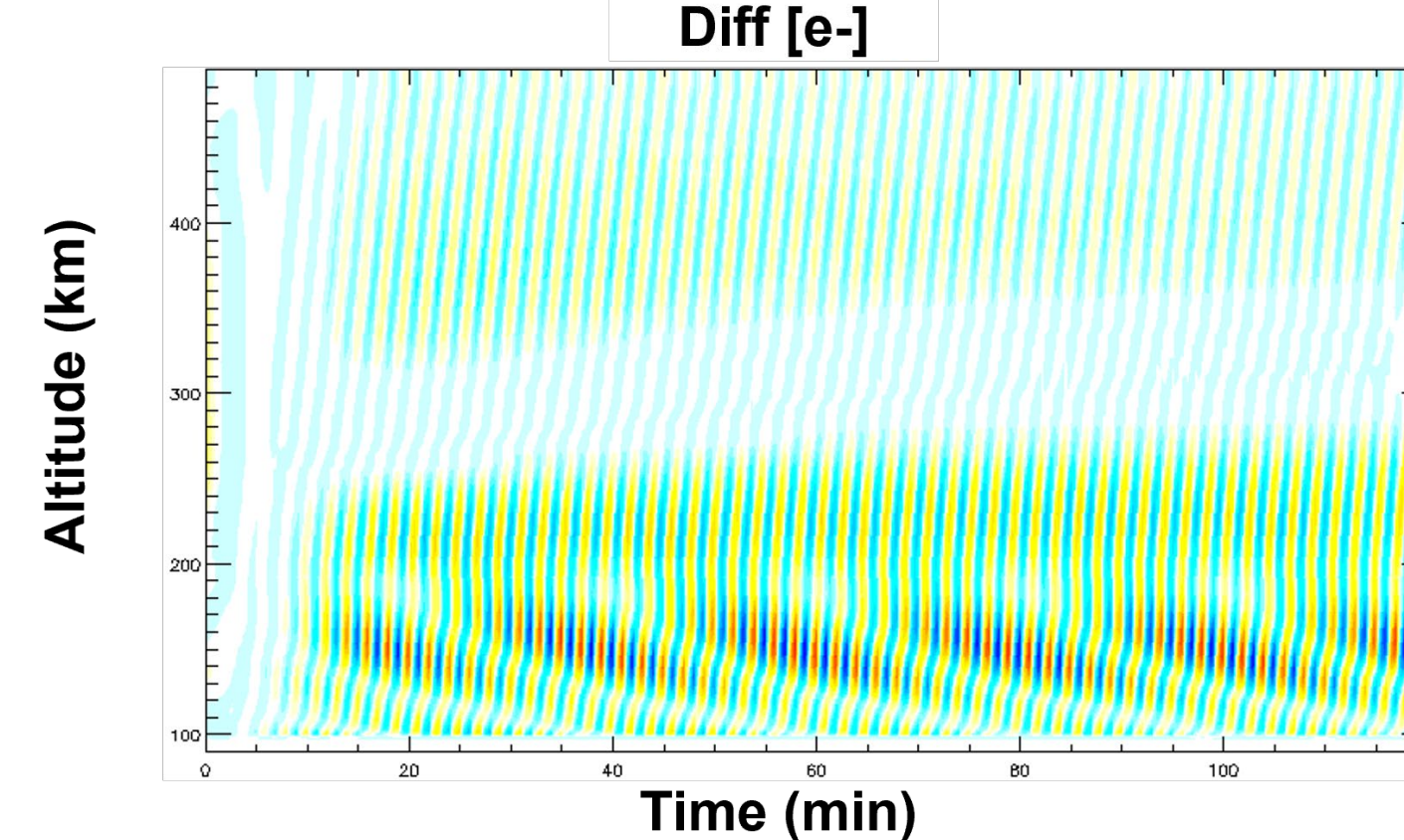


Fig 5. Example of AGWs in GITM-R simulation

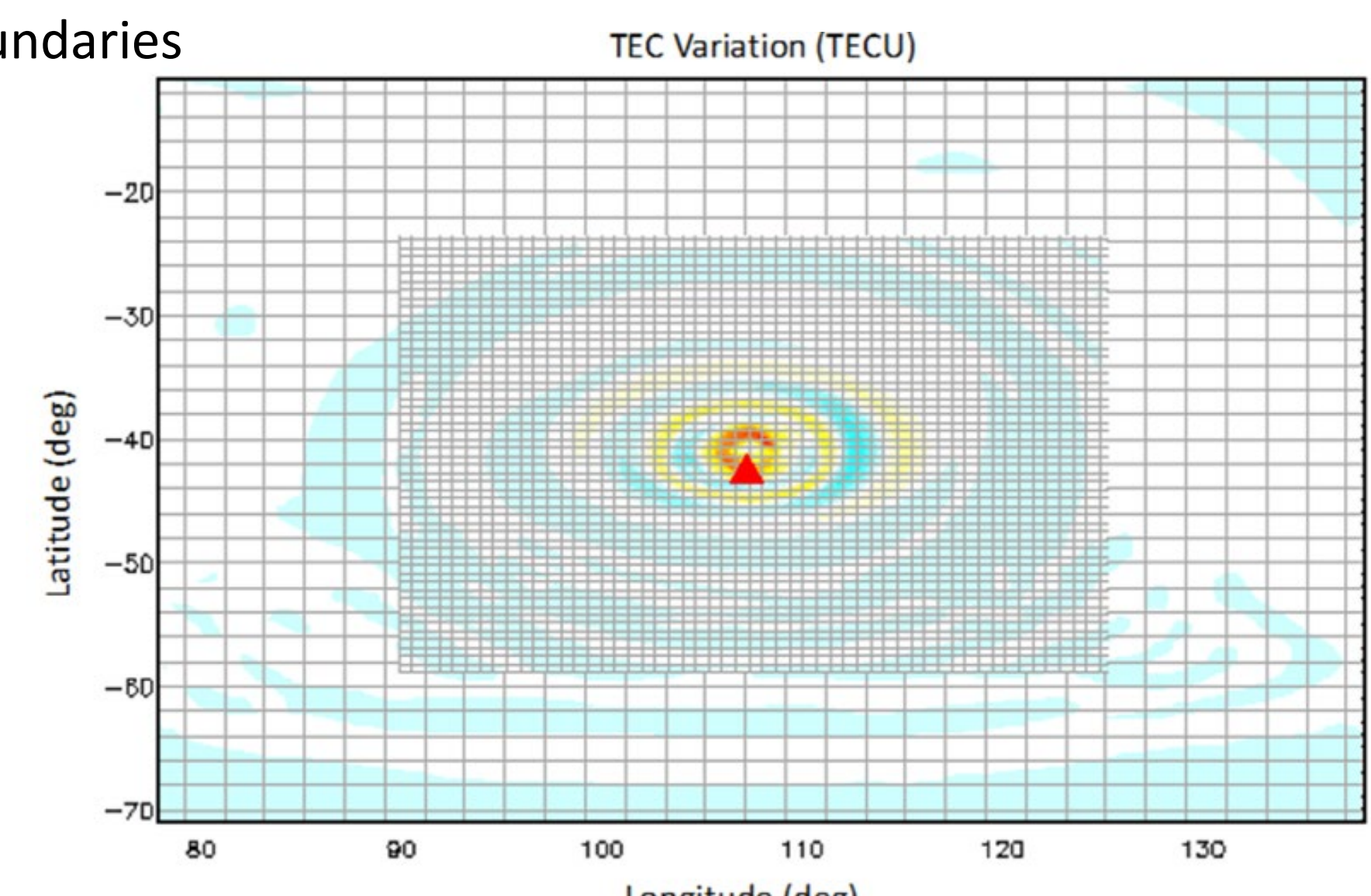


Fig 6. Grid configuration showing imbedded layers

Data-Model Comparisons

Direct Comparison

DP sim

- Travel time and magnitude of first wave packet and GW packet match well
- Magnitude distribution is recreated (EIA)
- GW perturbation take affect at larger distances
- GW packet arrival time is ~15 mins off

GC sim

- Multiple wave packets (4 total)
- Initial acoustic perturbation is offset
- Arrival of second wave packet matches
- Maximum of second wave packet occurs before data
- GW packet not noticeably present

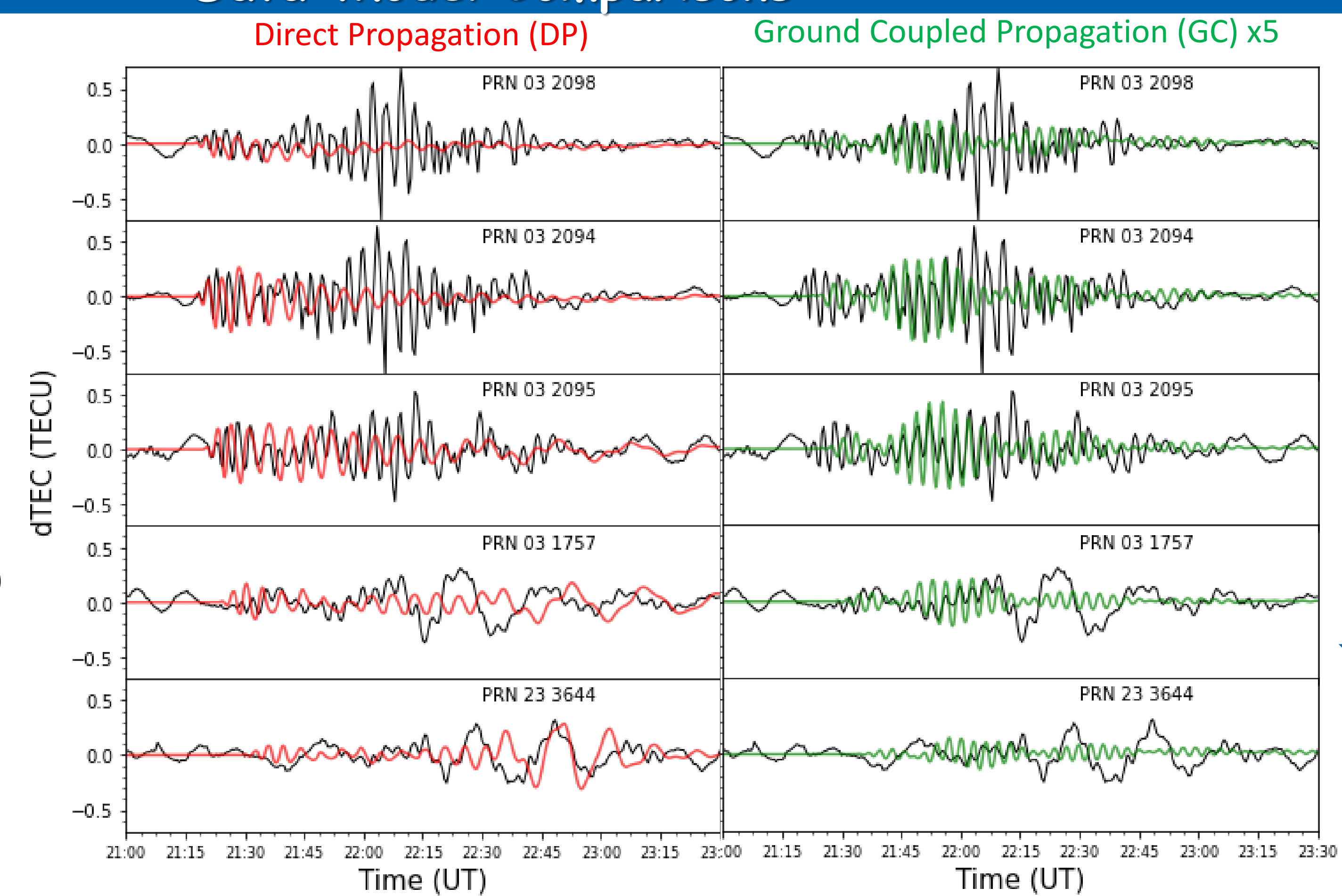


Fig 7. Direct comparison of GNSS observation and GITM-R simulation

Spectral Comparison

- DP settles to near forcing frequency (dashed)
- DP GW mode has slightly larger period
- GC dominate modes matches much better to data
- GC GW mode is present but small

Phase Speed Comparison

- DP AGWs have phase speed ~730 m/s while GW mode ~190 m/s (not shown)
- GC AGWs have phase speeds ~730 m/s with GW mode ~260 m/s

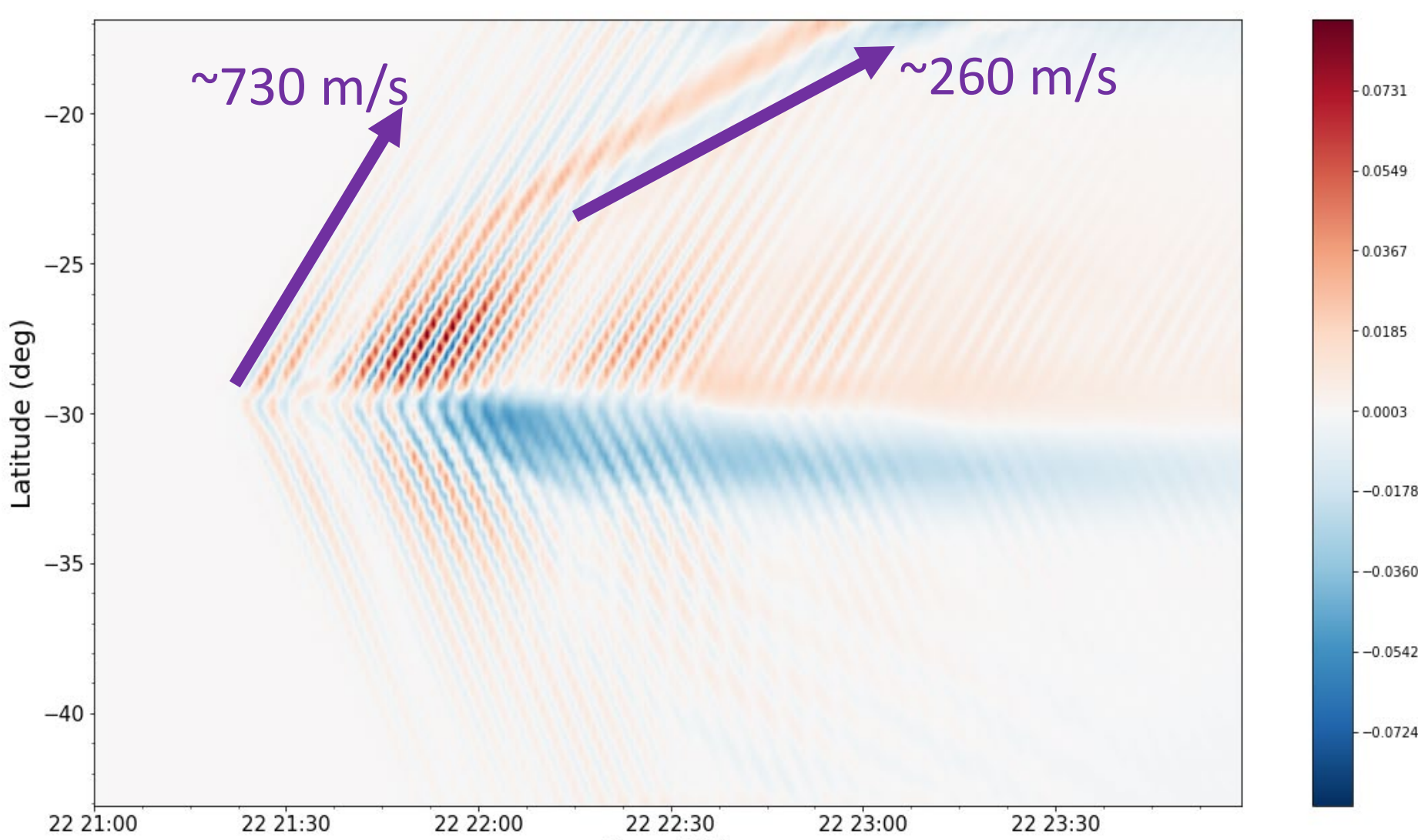


Fig 6. Latitude vs Time slice of GC dTEC at longitude of volcano

Discussion and Conclusions

- GITM-R was shown to recreate important features of the observed GNSS data such as travel times and relative magnitudes
- GITM-R was able to reproduce spectral peaks of ~5.0 mHz and ~1 mHz typically seen in CVIDs
- The simplified specification of GC propagation to force GITM-Rs lower boundary was able to improve data-model comparison of this event.
- The First wave packet is likely acoustic waves propagating from directly above the vent
- The Second wave packet may be a localized forcing due to the passage of a ground coupled airwave
- Comparison of GW mode frequencies, temporal locations, and phase speeds between the specifications suggest they may have been launched between 2-100 km

Future Works

- Reduced order model of atmospheric resonance to better match GC magnitude and spectral content with observation

References

- [1] Cahyadi, M. N. et al. (2020), doi:10.1016/j.jvolgeores.2020.107047
- [2] Shults, K. et al. (2016), doi:10.1002/2016JA023382
- [3] Zhang, Shun-Rong, data from Haystack Observatory, MIT
- [4] Kanamori, H et al. (1994) Excitation of oscillation by volcanic eruptions, J. Geophys. Res. 99
- [5] Meng X, et al. (2018), doi:10.1029/2018JA025253
- [6] Godin O. A. et al. (2020), doi:10.1186/s40623-020-01260-9
- [7] Deng Y., et. Al. (2020), Global Ionosphere-Thermosphere Model with Local Mesh Refinement, J. Geophys. Res., Submitted.