

1998 CEDAR Workshop
Boulder, Colorado
June 7-12, 1998

Tutorial Lecture

by Timothy Fuller-Rowell
Space Environment Center, NOAA

**Polar Aeronomy: Thermosphere-Ionosphere
Interactions above 100 km**

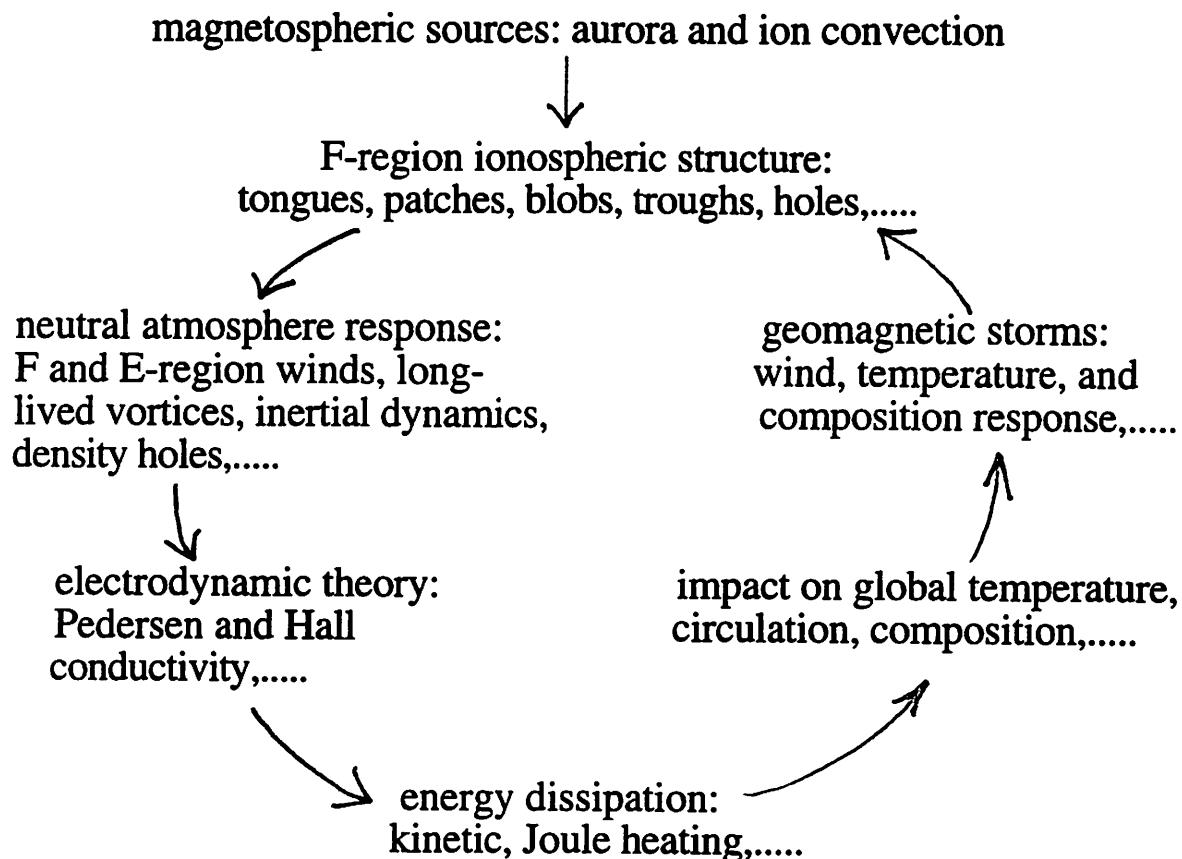
POLAR AERONOMY:

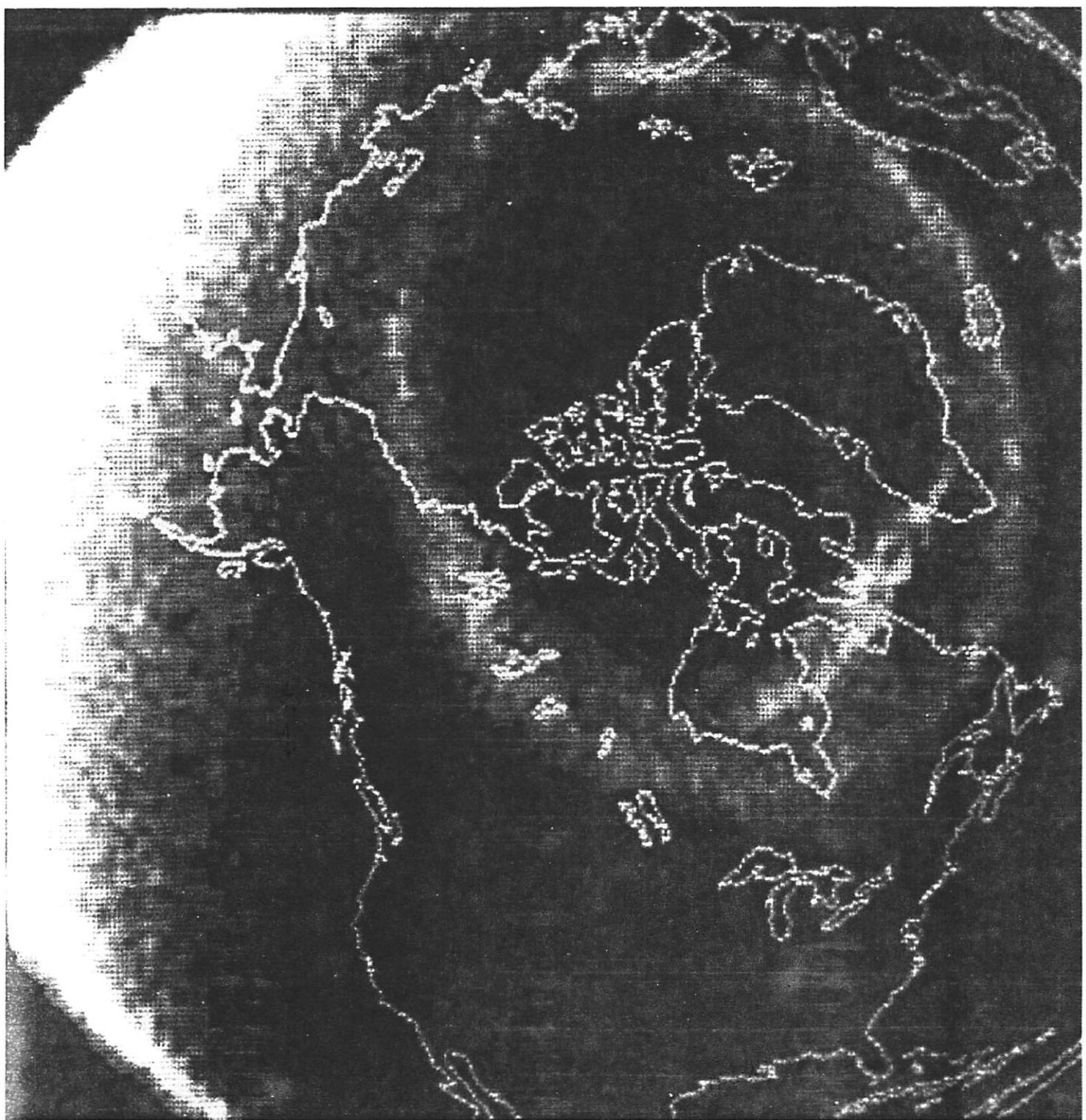
Thermosphere-Ionosphere Interactions above 100 km

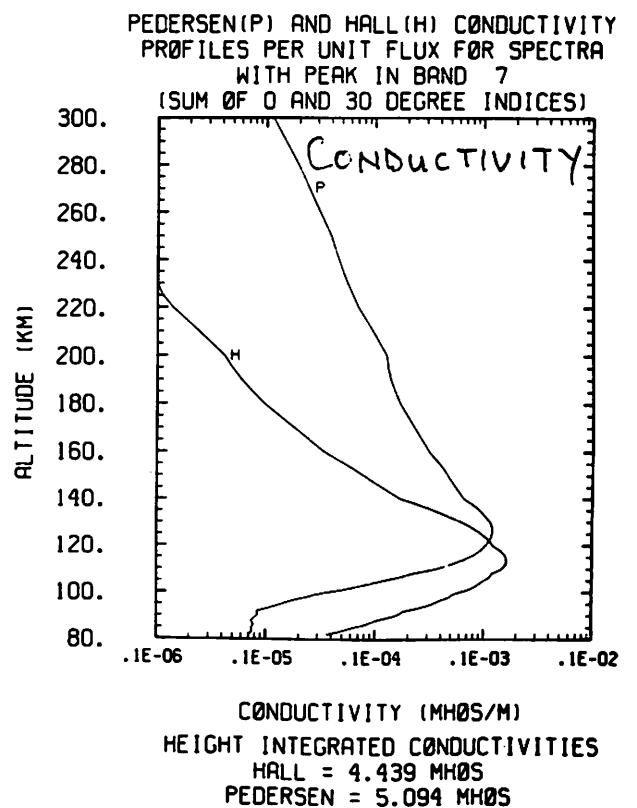
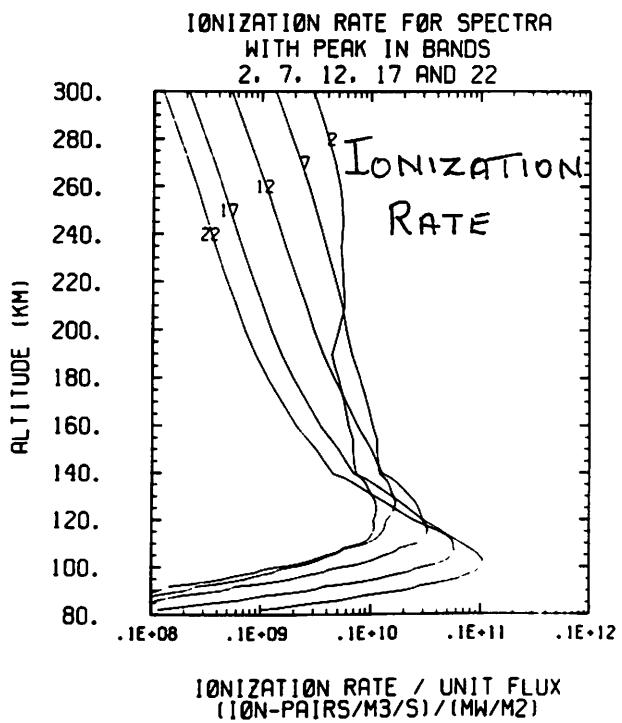
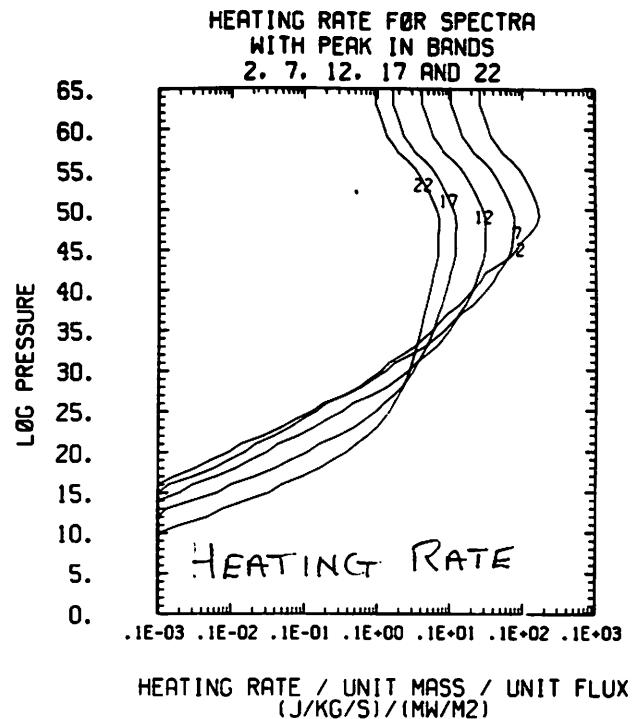
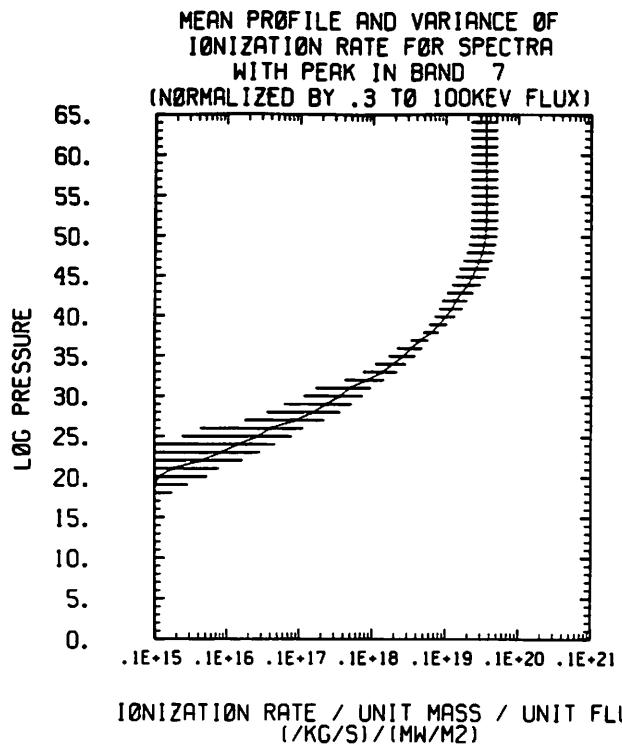
by

Tim Fuller-Rowell

CONTENTS:

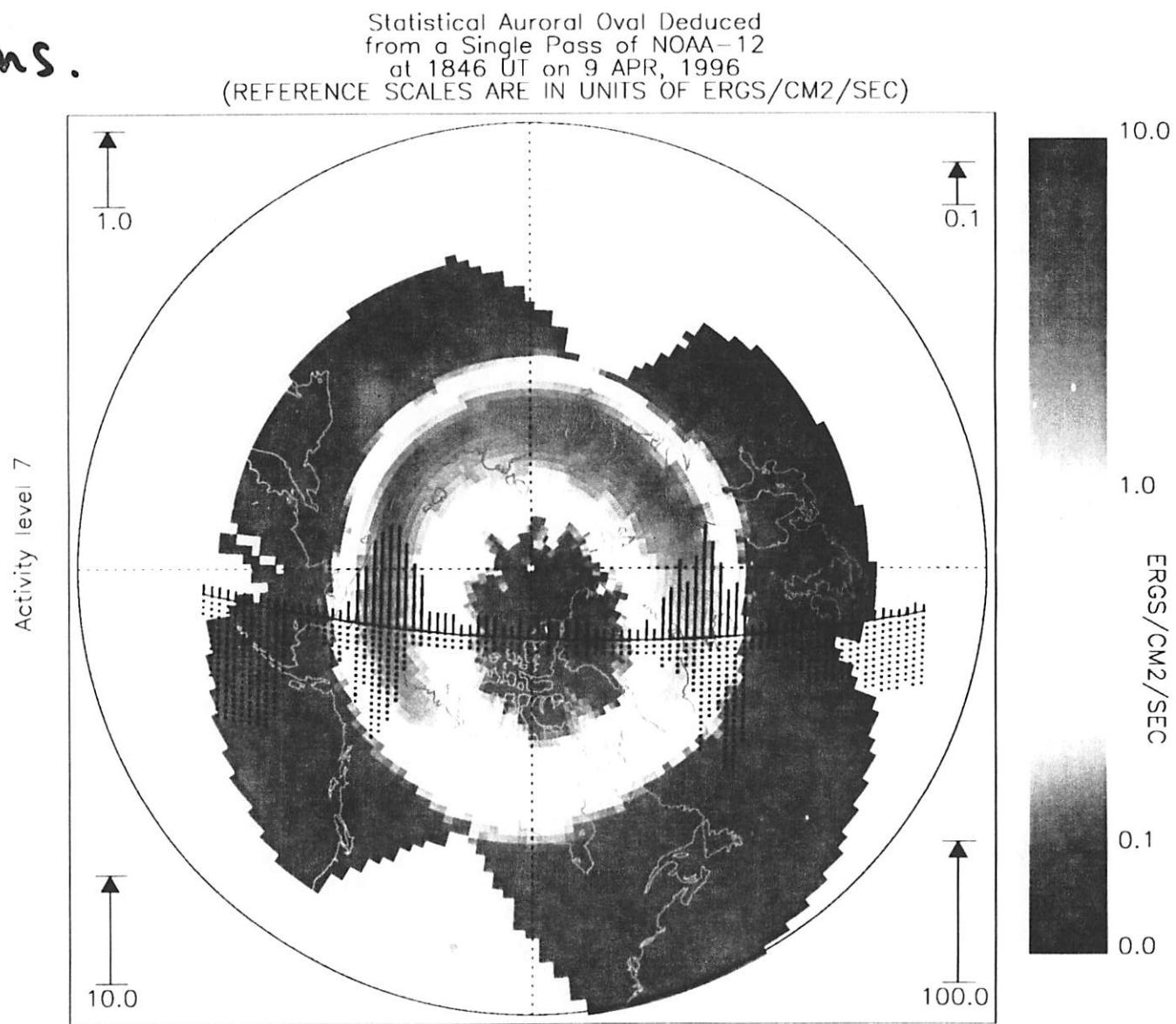




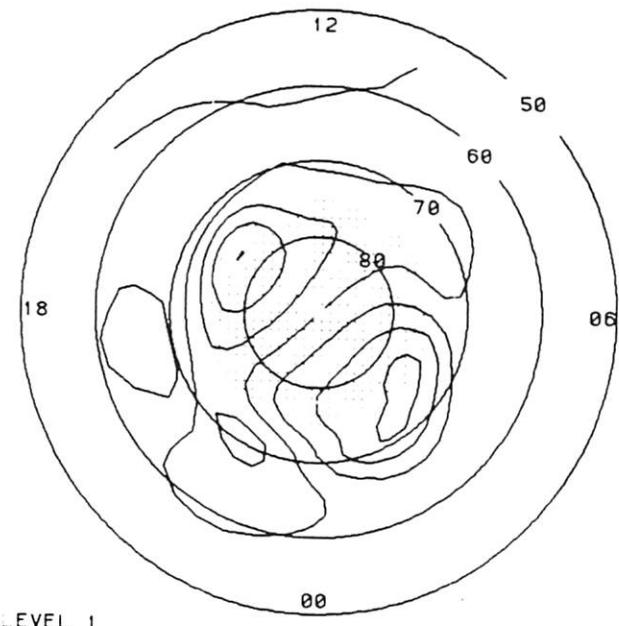


DATA :- TIROS/NOAA AURORAL PARTICLES

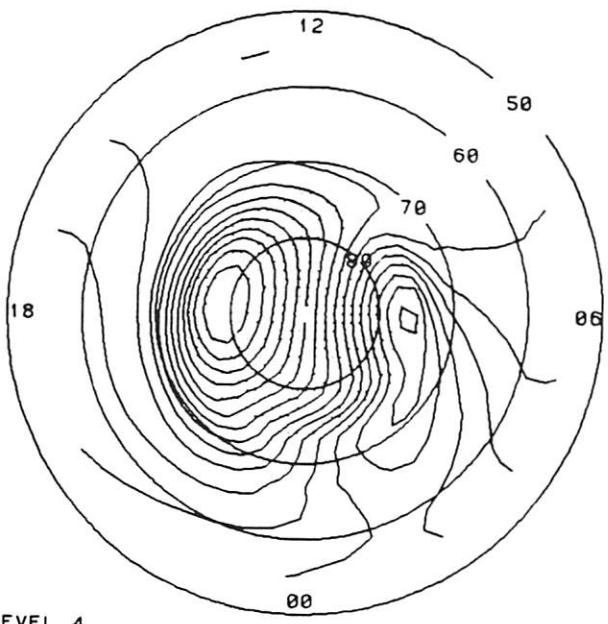
Evans.



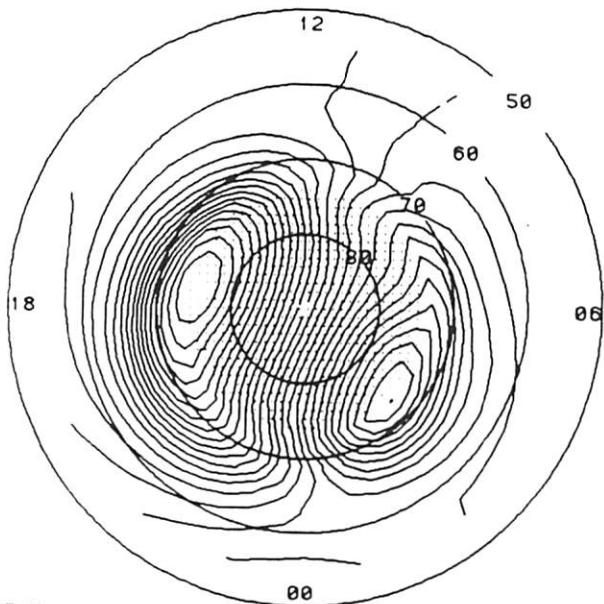
FOSTER ET AL. 1996



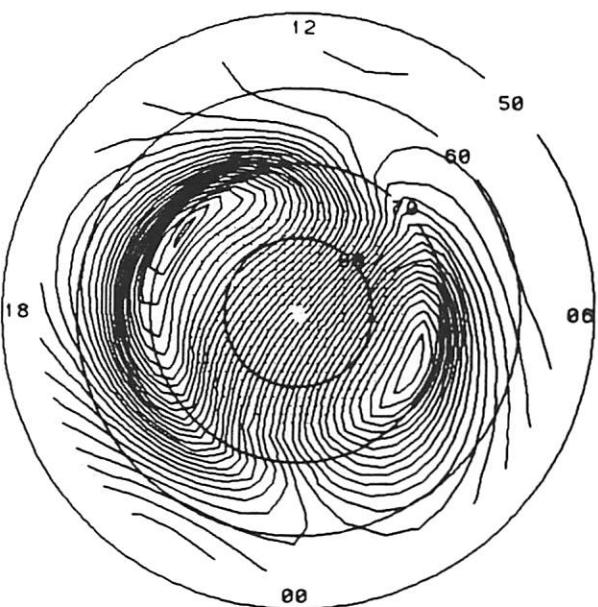
LEVEL 1



LEVEL 4



LEVEL 7



LEVEL 9

HEPPNER AND MAYNARD, 1987

HEPPNER AND MAYNARD: EMPIRICAL ELECTRIC FIELD MODELS

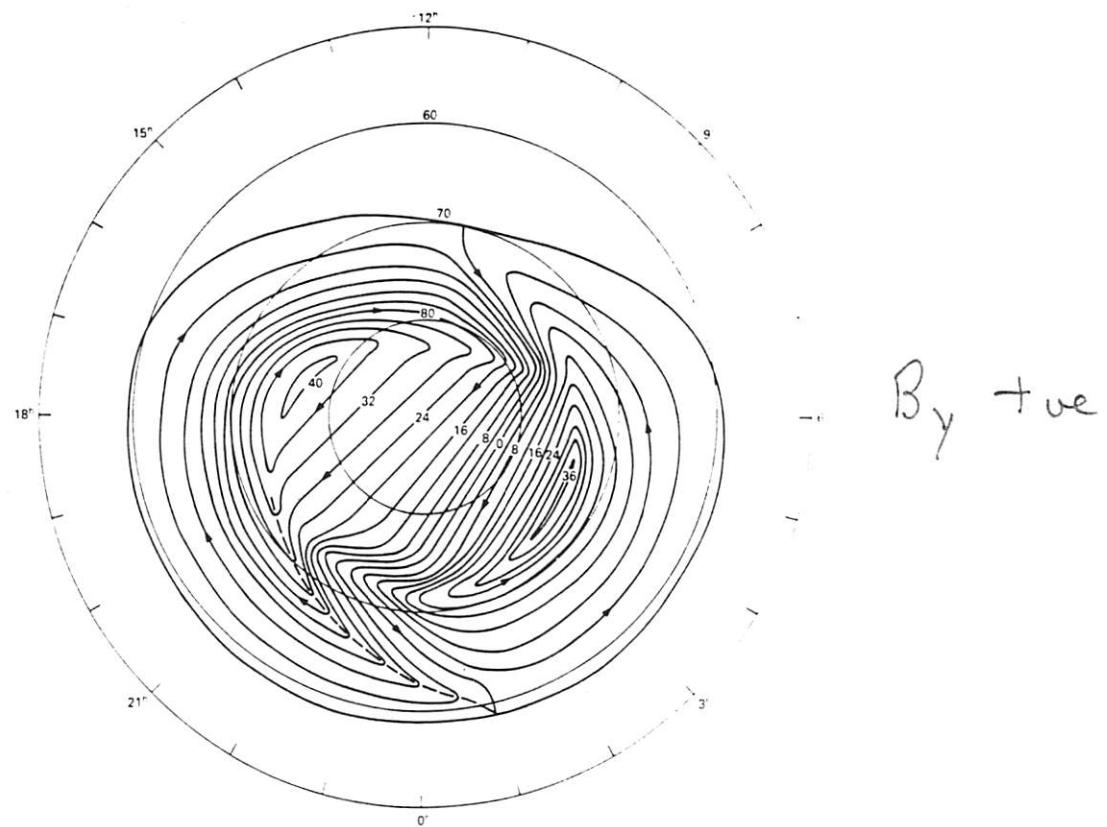


Fig. 3. Convection electric field model BC, representing the most typical pattern distribution encountered in the northern hemisphere under +Y IMF conditions and in the southern hemisphere under -Y IMF conditions ($3^{\circ} \leq K_p \leq 4^{\circ}$).

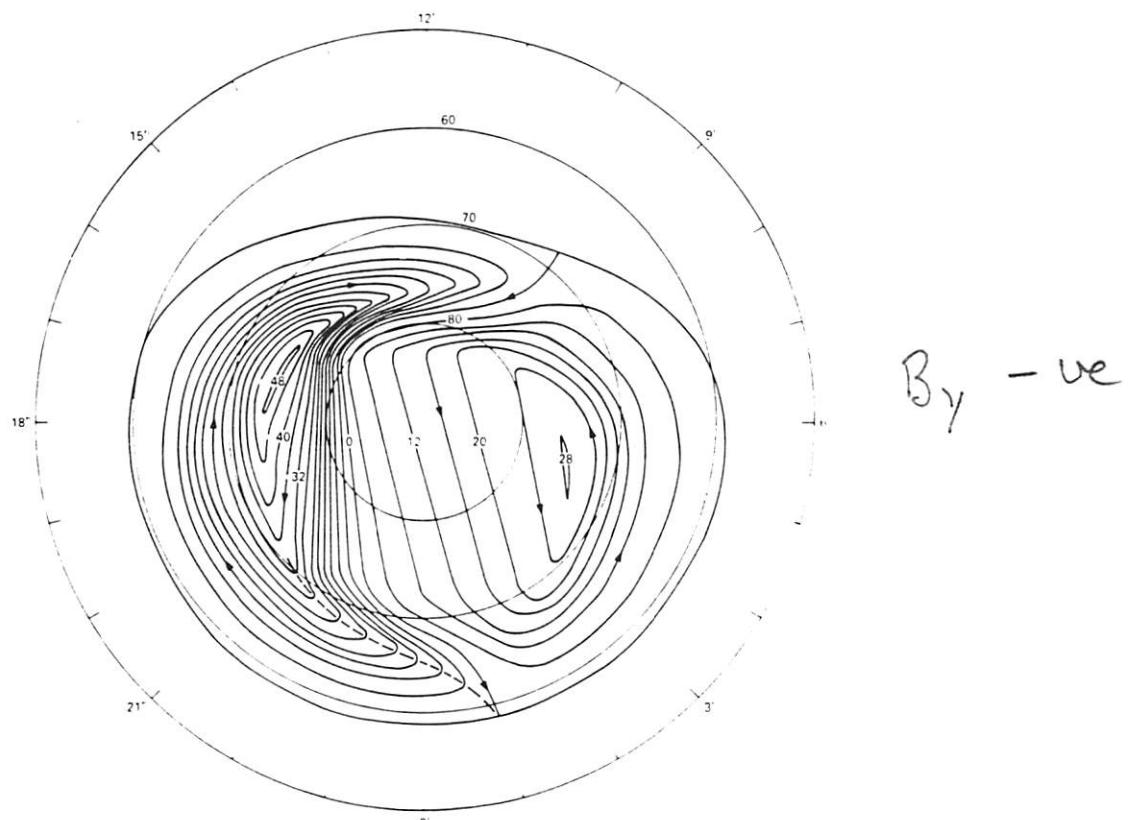
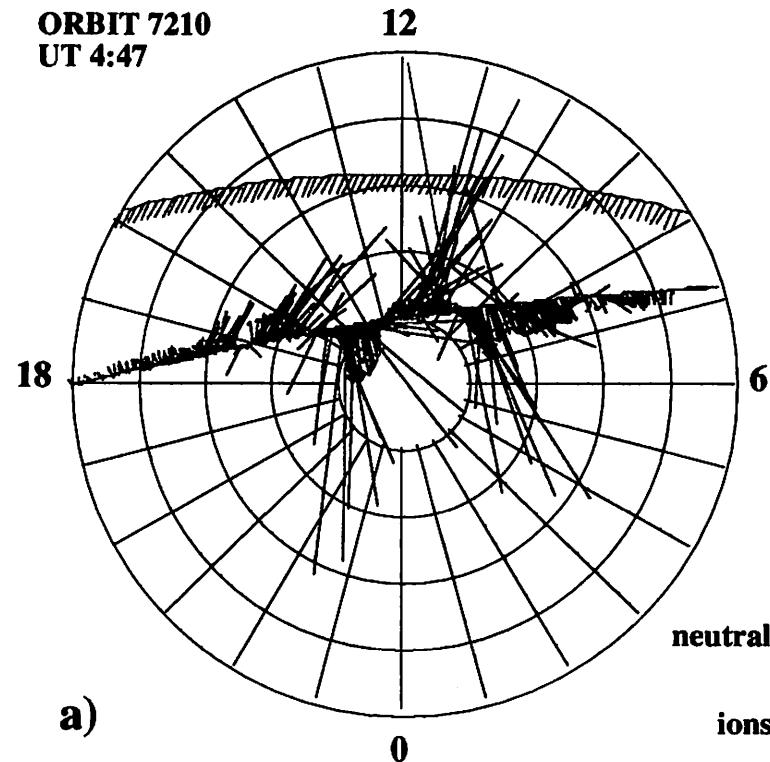


Fig. 2. Convection electric field model DE, one of two models representing pattern distributions encountered in the northern hemisphere under -Y IMF conditions and in the southern hemisphere under +Y IMF conditions ($3^{\circ} \leq K_p \leq 4^{\circ}$).

DE 2 FPI/WATS/IDM/RPA DAY 82328

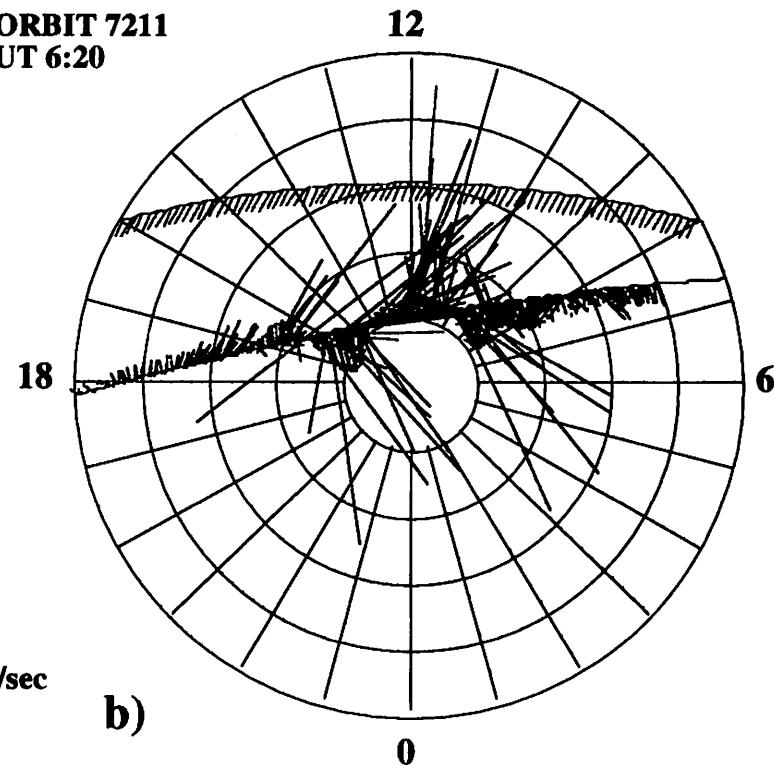
B_z north.

ORBIT 7210
UT 4:47



a)

ORBIT 7211
UT 6:20



b)

Killeen et al., JATP 1991

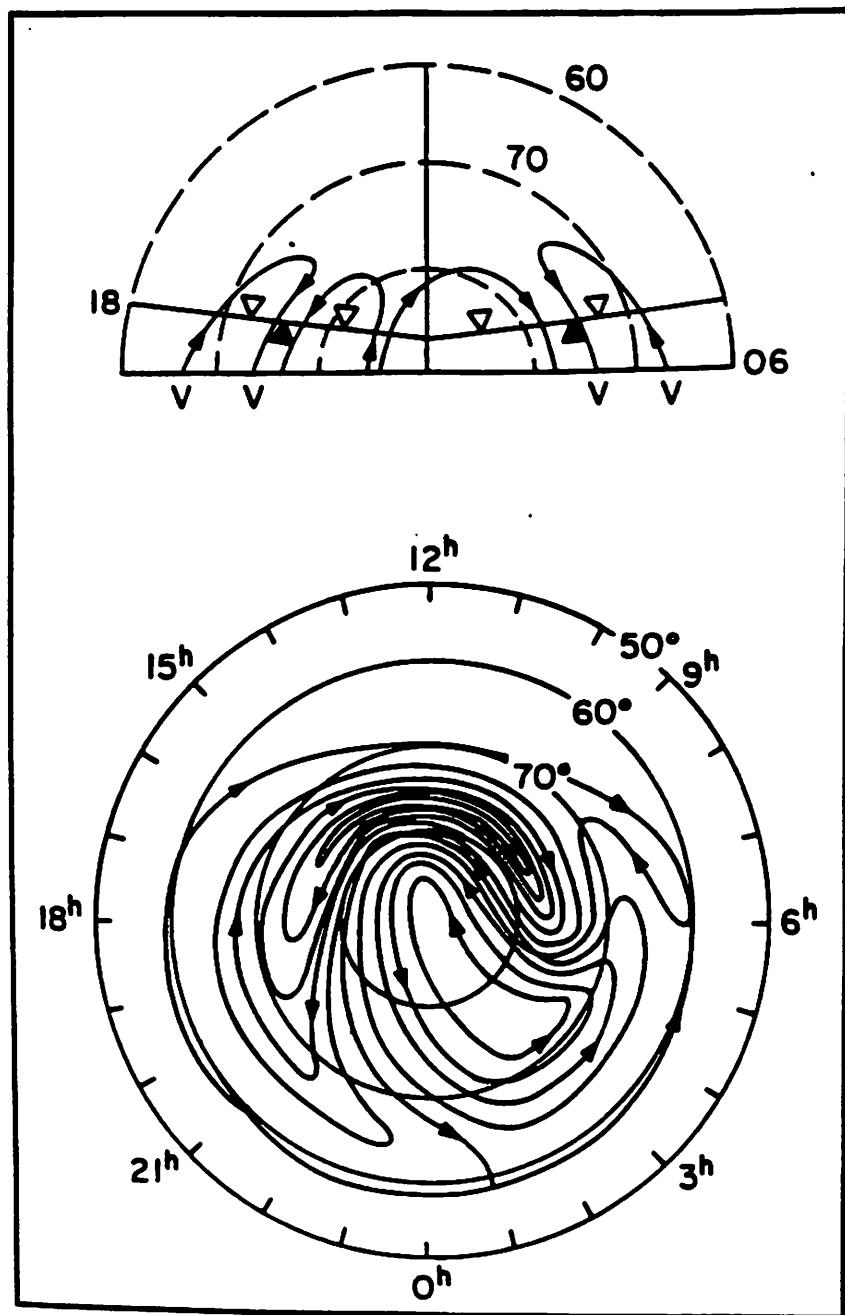


Figure 2.4.

Two different hypothetical polar cap convection patterns based on the same experimental data.

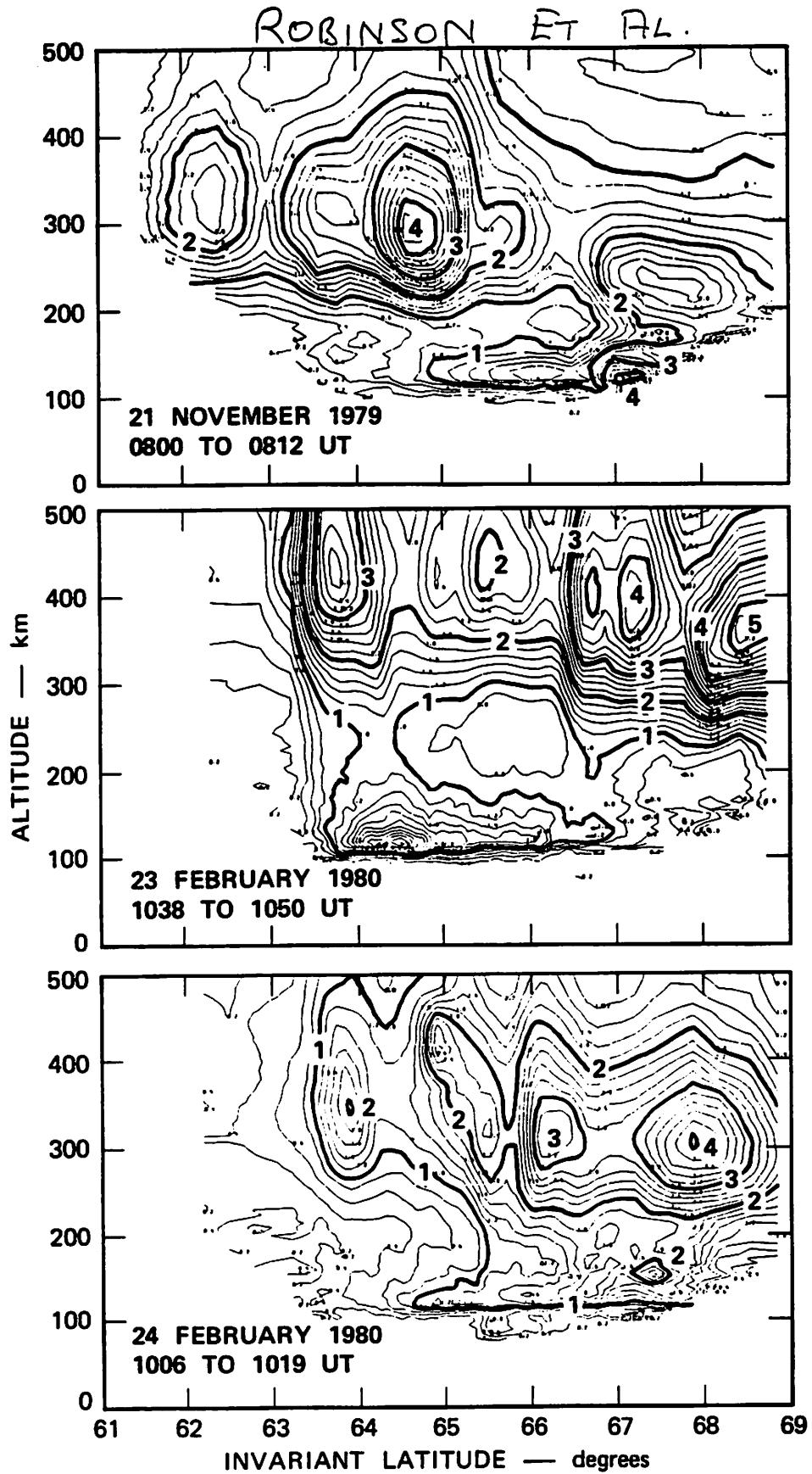
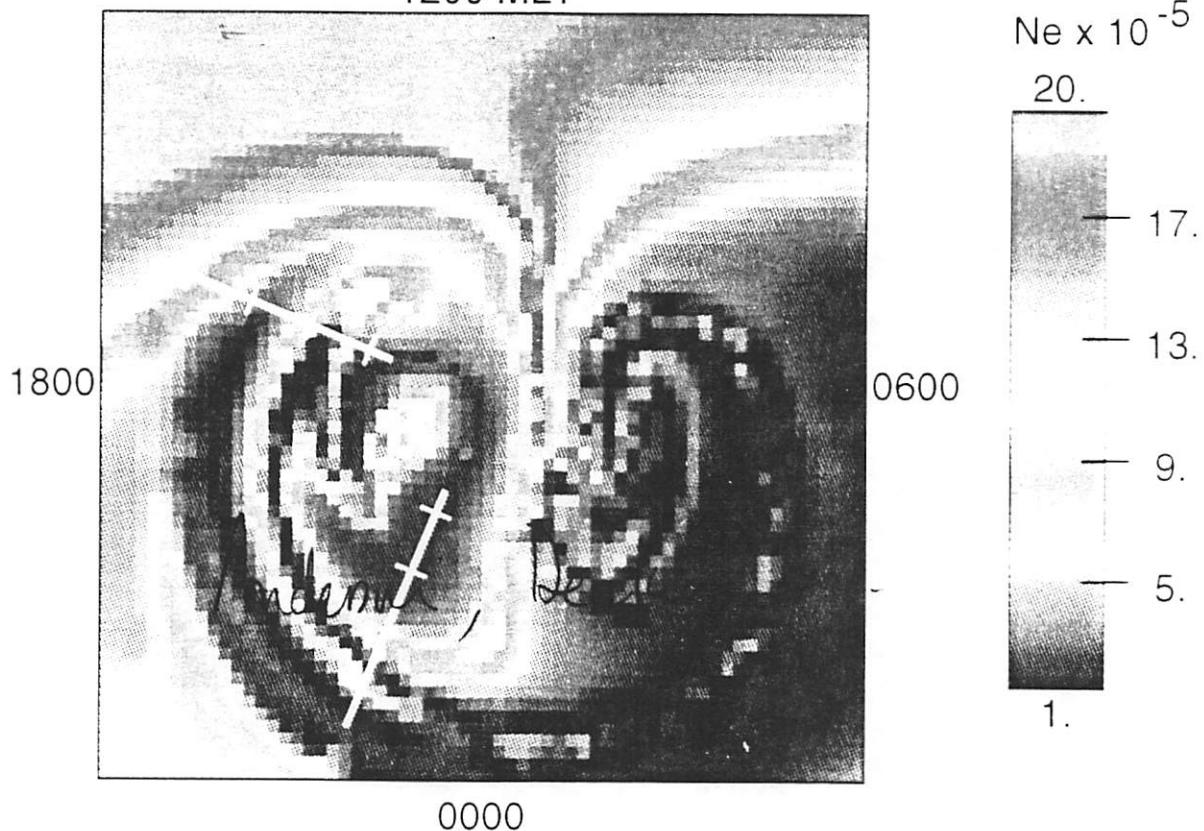


Fig. 1. Examples of elevation scan measurements of electron density obtained on three different evenings.

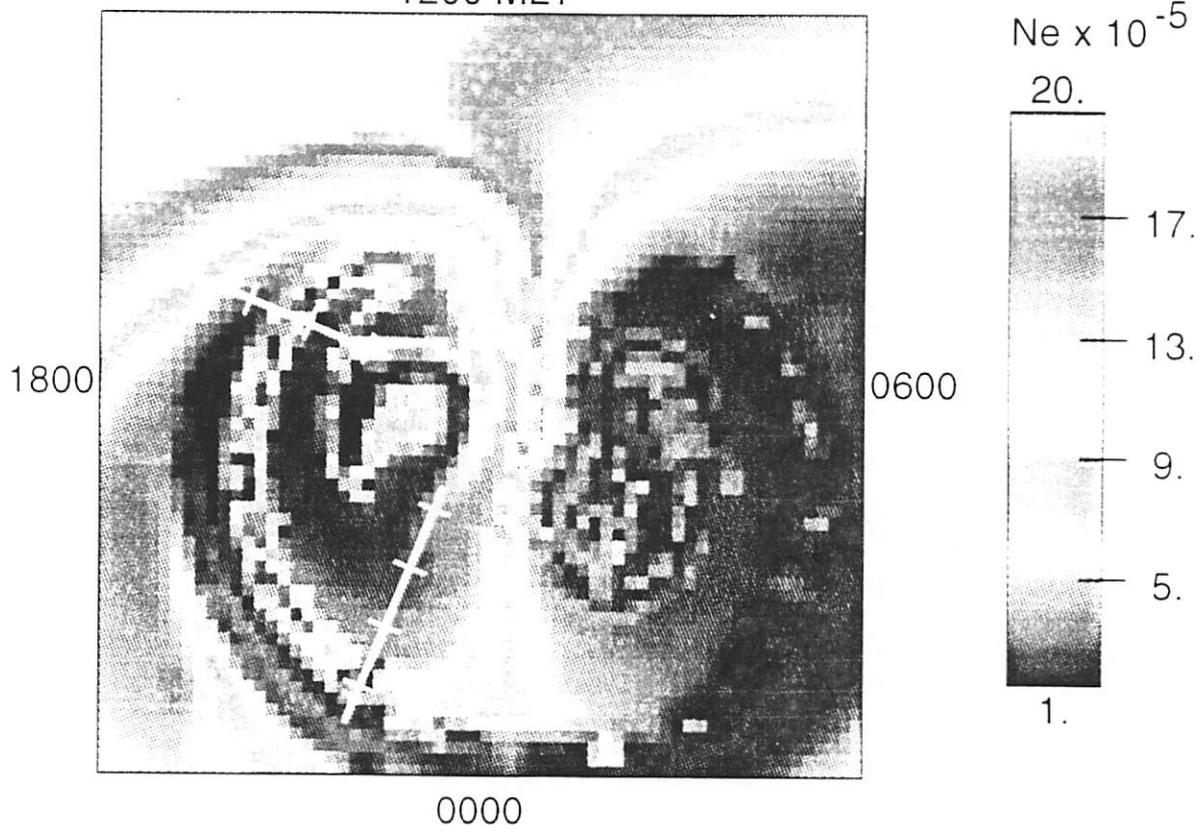
NmF2, 2030 UT

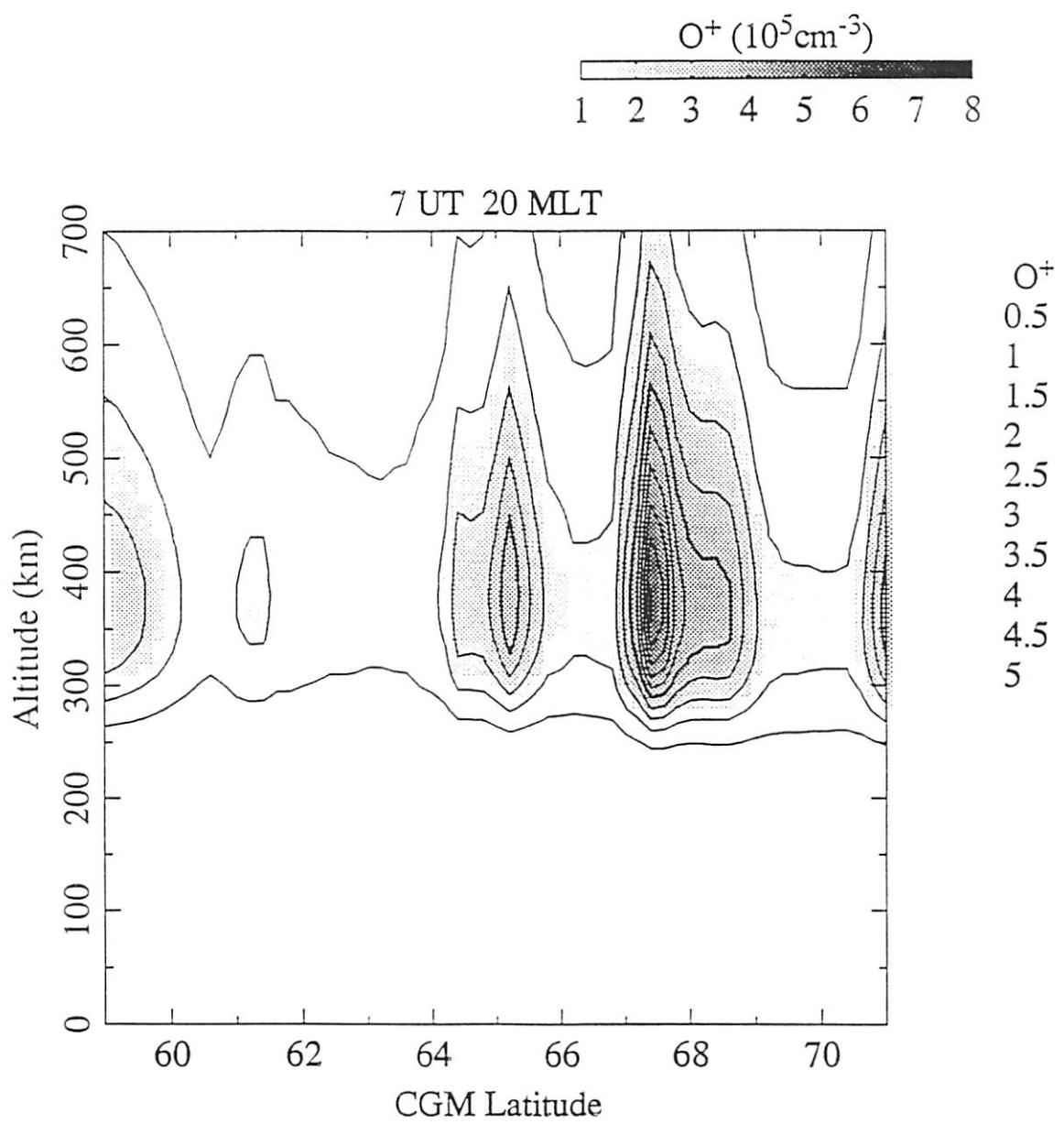
1200 MLT



NmF2, 2130 UT

1200 MLT





SOTKA, SCHUNK, ...

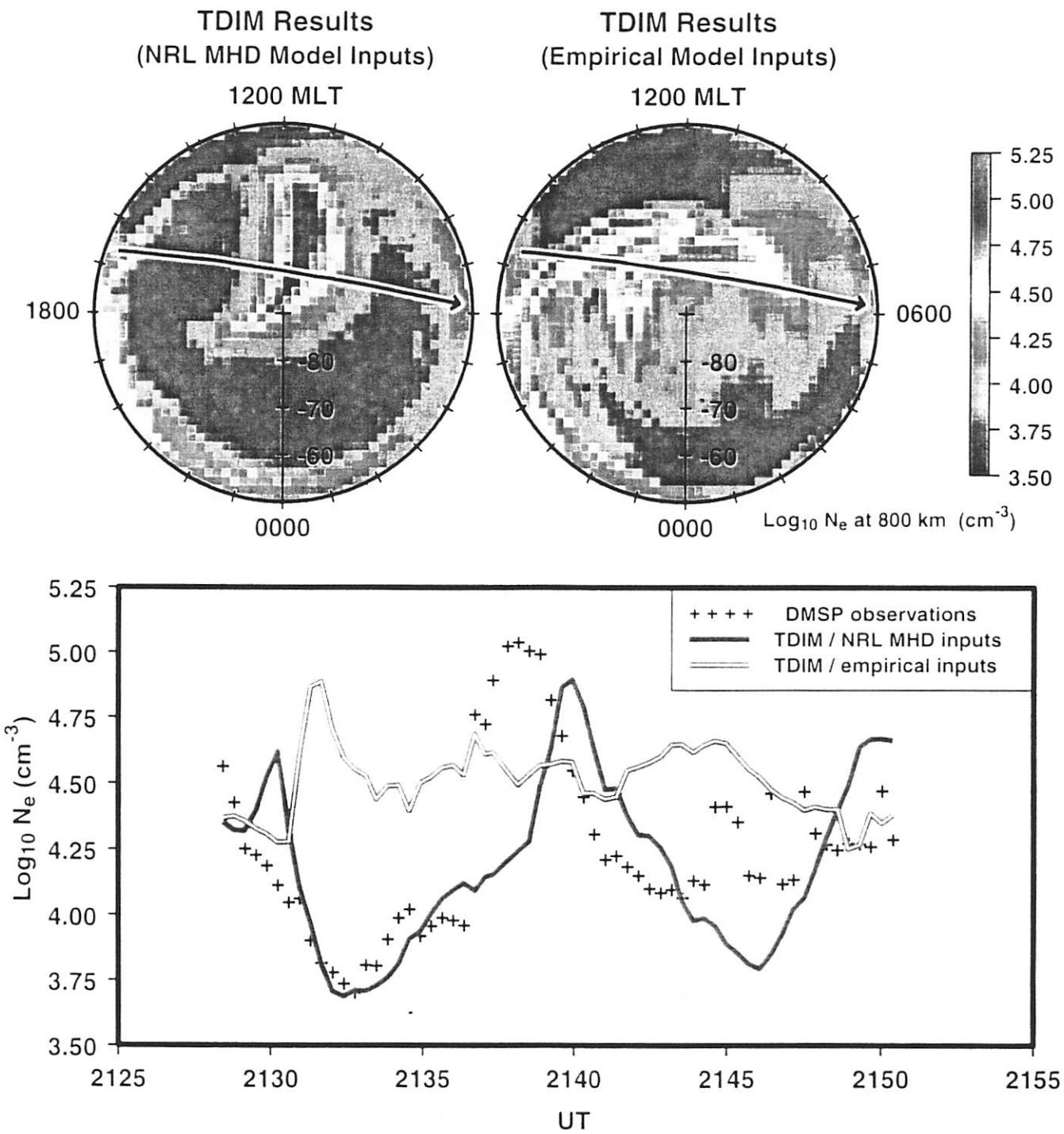
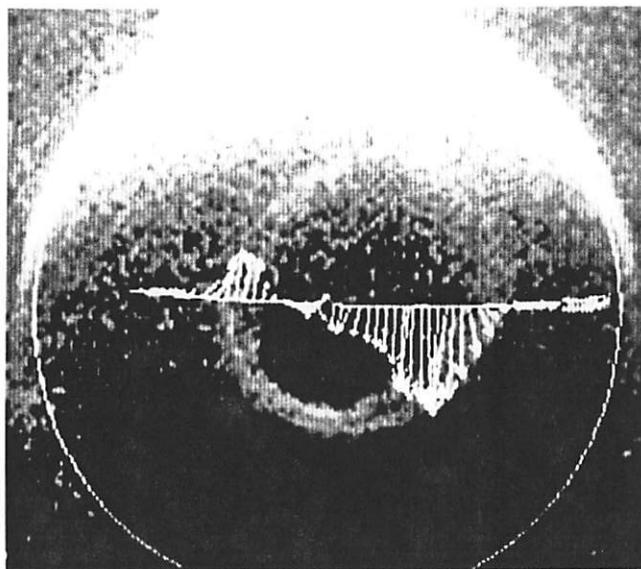


Plate 4

1853 1505 UT

6 DEC 1981

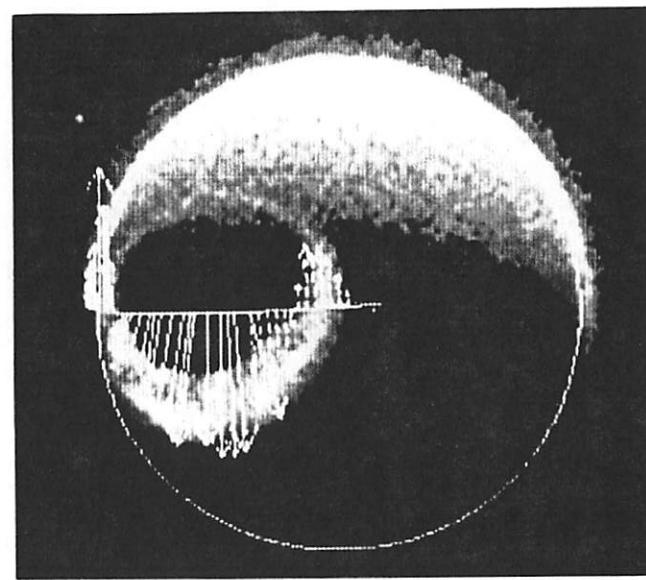


A

500 m/s

1941 1244 UT

12 DEC 1981

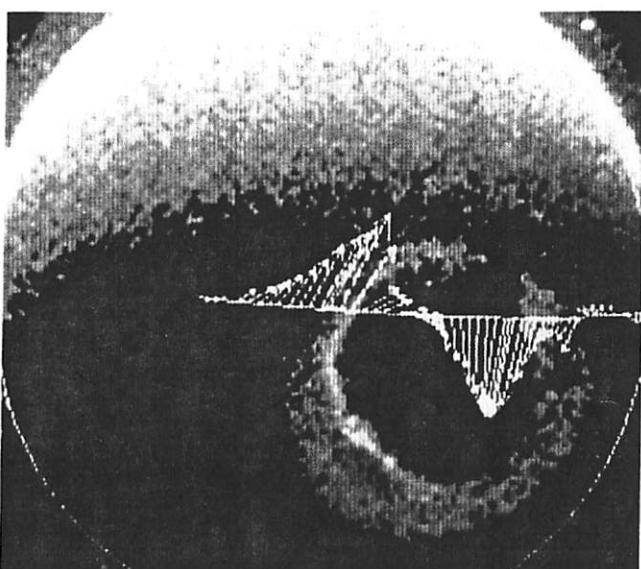


B

500 m/s

1849 0838 UT

6 DEC 1981

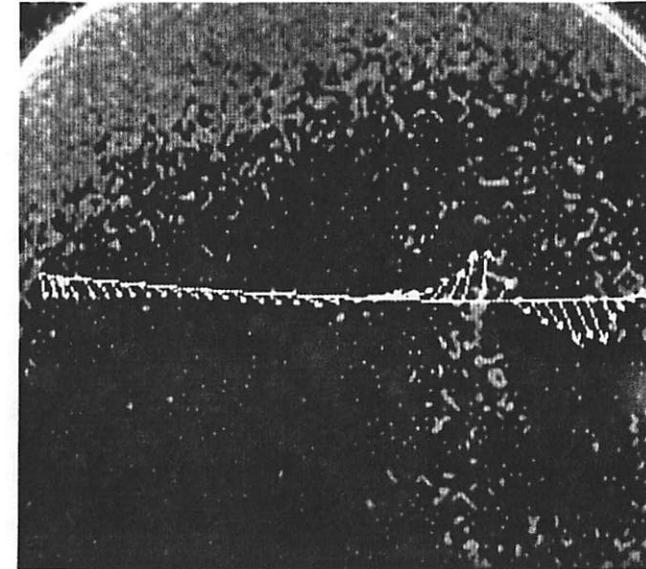


C

500 m/s

1981 0505 UT

15 DEC 1981



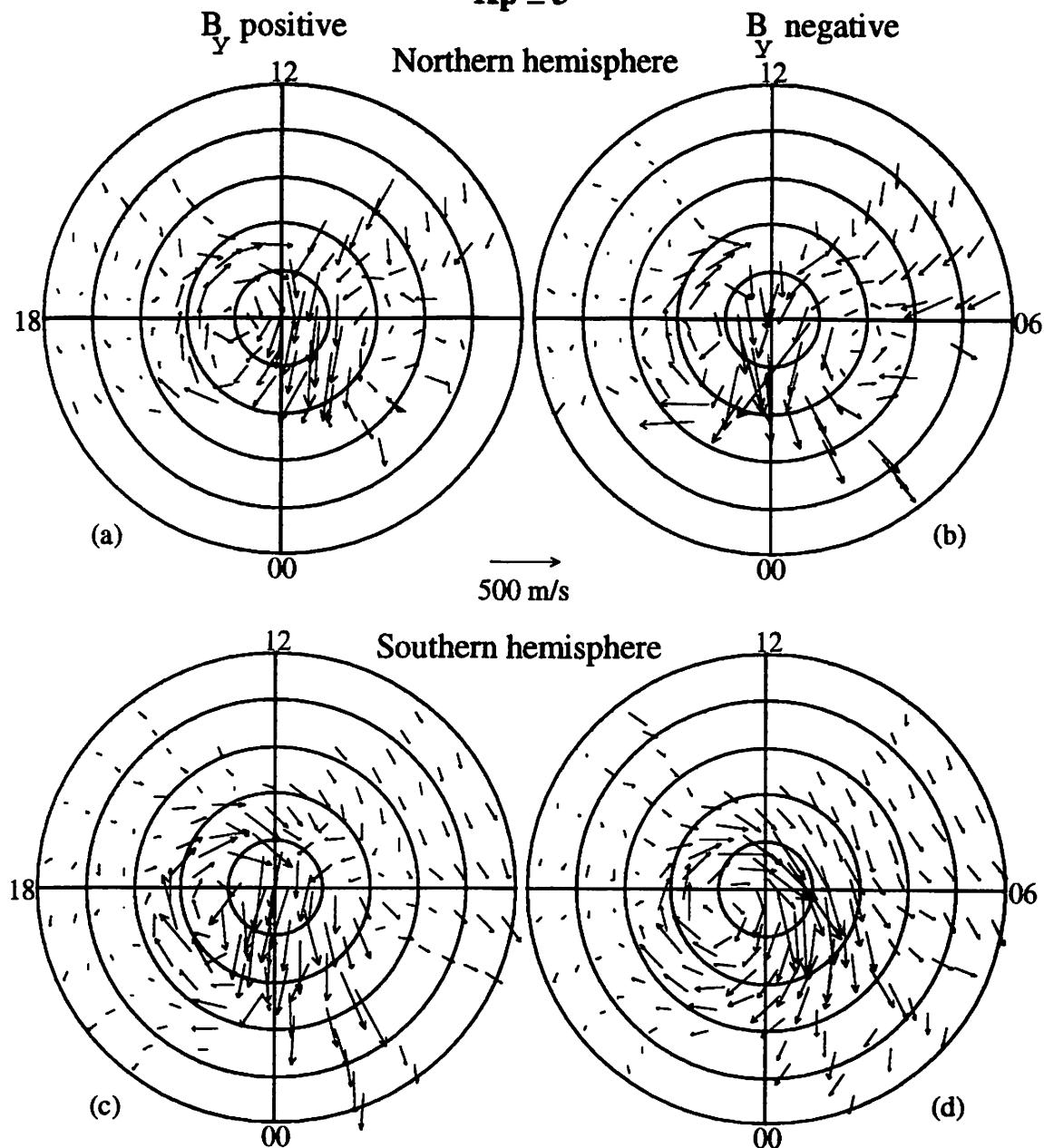
D

500 m/s

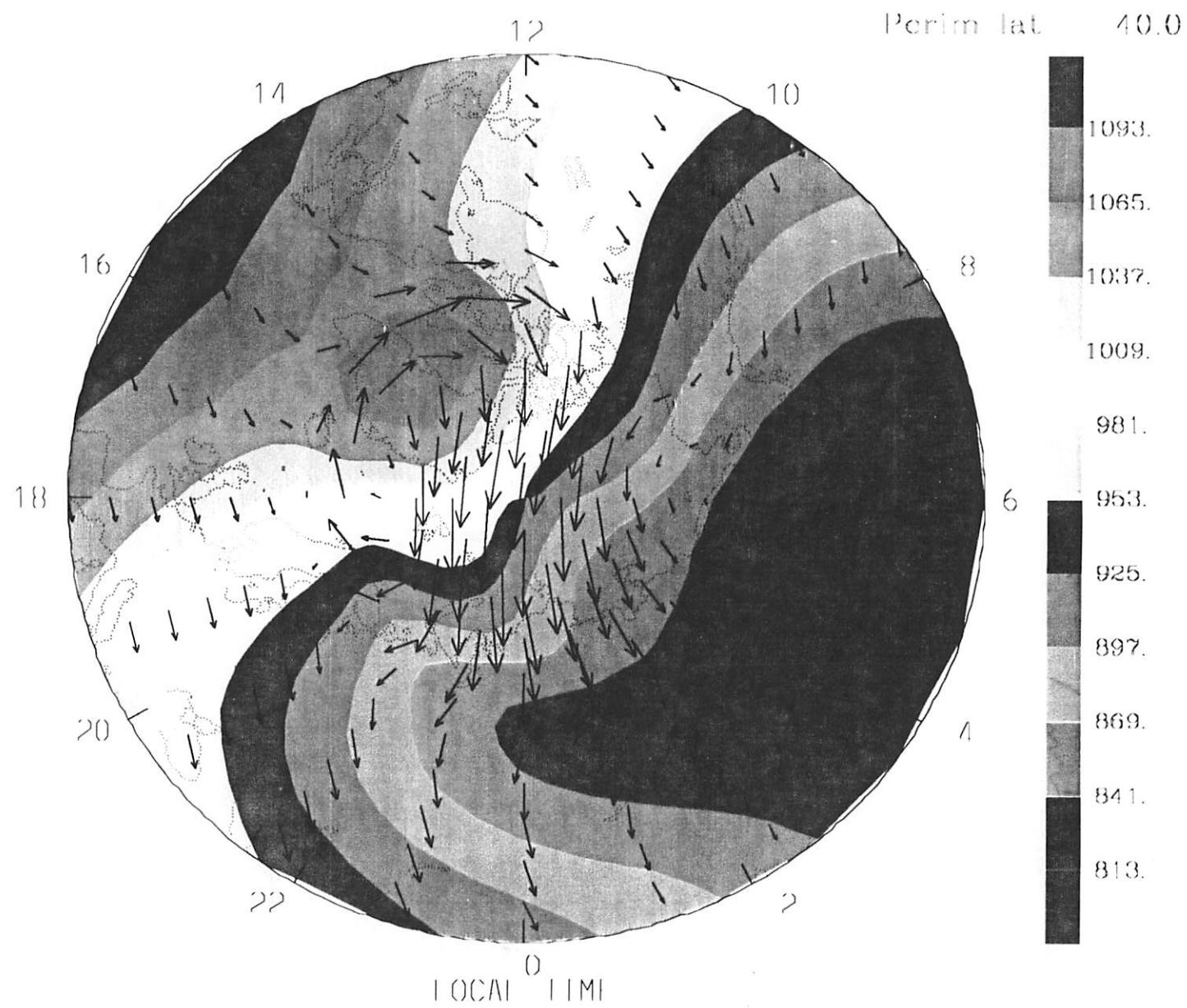
Killeen et al.

Thayer et al., JGR, 1990

B_y -DEPENDENCE OF DE-2 NEUTRAL WINDS
(in geomagnetic coordinates)
 $K_p \leq 3$



CTIM NEUTRAL TEMPERATURE (DEG. K) 855



UT 18.0 Pressure level 12
MINIMUM 784.7, MAXIMUM 1121.2, CONTOUR INTERVAL 28.0

405 v/s
>

EQUATION OF MOTION

$$\frac{D}{Dt} \vec{V} = -\frac{1}{\rho} \nabla P - 2\vec{\Omega} \times \vec{V} - \nu_{ni}(\vec{V} - \vec{U}) + \frac{1}{\rho} \nabla(\mu \nabla \vec{V}) + \vec{g}$$

Pressure Coriolis Ion drag Viscosity Gravity
 gradient

$$\frac{D}{Dt} X = \frac{\partial}{\partial t} X + (\vec{V} \cdot \nabla) X$$

$$\frac{\partial}{\partial t} V_\theta = -\frac{V_\theta}{r} \frac{\partial}{\partial \theta} V_\theta - \frac{V_\phi}{r \sin \theta} \frac{\partial}{\partial \phi} V_\theta - \omega \frac{\partial}{\partial p} V_\theta - \frac{g}{r} \frac{\partial}{\partial \theta} h$$

$$+ \left(2\Omega + \frac{V_\phi}{r \sin \theta} \right) V_\phi \cos \theta + g \frac{\partial}{\partial p} \left[(\mu_m + \mu_T) \frac{p}{H} \frac{\partial}{\partial p} V_\theta \right] - \nu_{ni} (V_\theta - U_\theta)$$

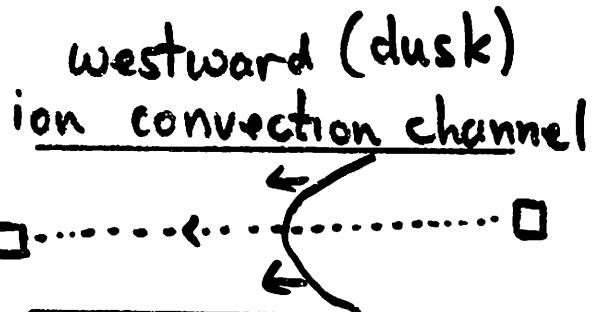
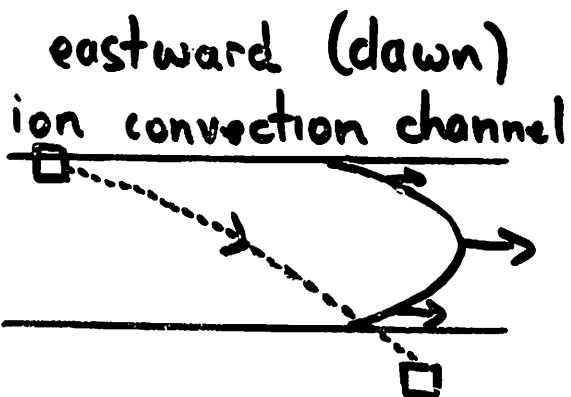
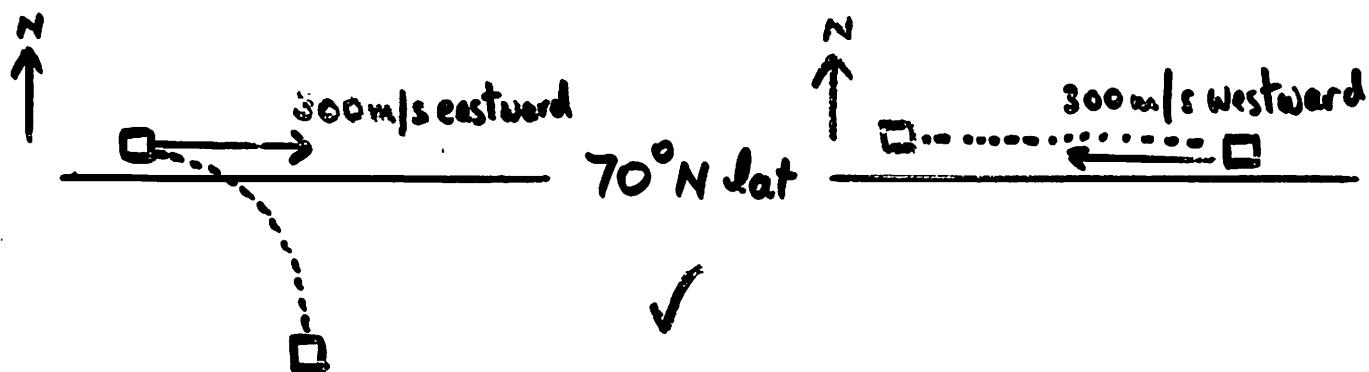
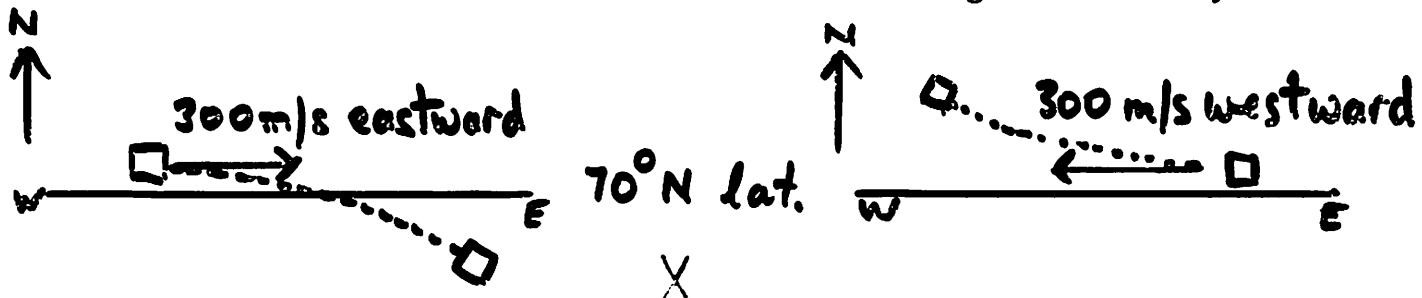
$$\frac{\partial}{\partial t} V_\phi = -\frac{V_\theta}{r} \frac{\partial}{\partial \theta} V_\phi - \frac{V_\phi}{r \sin \theta} \frac{\partial}{\partial \phi} V_\phi - \omega \frac{\partial}{\partial p} V_\phi - \frac{g}{r \sin \theta} \frac{\partial}{\partial \phi} h$$

$$+ \left(2\Omega + \frac{V_\phi}{r \sin \theta} \right) V_\theta \cos \theta + g \frac{\partial}{\partial p} \left[(\mu_m + \mu_T) \frac{p}{H} \frac{\partial}{\partial p} V_\phi \right] - \nu_{ni} (V_\theta - U_\theta)$$

EQUATION OF MOTION

$$\frac{D}{Dt} \mathbf{V} = -\frac{1}{\rho} \nabla P - 2\Omega \wedge \mathbf{V} - \mathbf{v}_{ni} (\mathbf{V} - \mathbf{U}) + \frac{1}{\rho} \nabla (\mu \nabla \mathbf{V})$$

Pressure Coriolis Ion Drag Viscosity



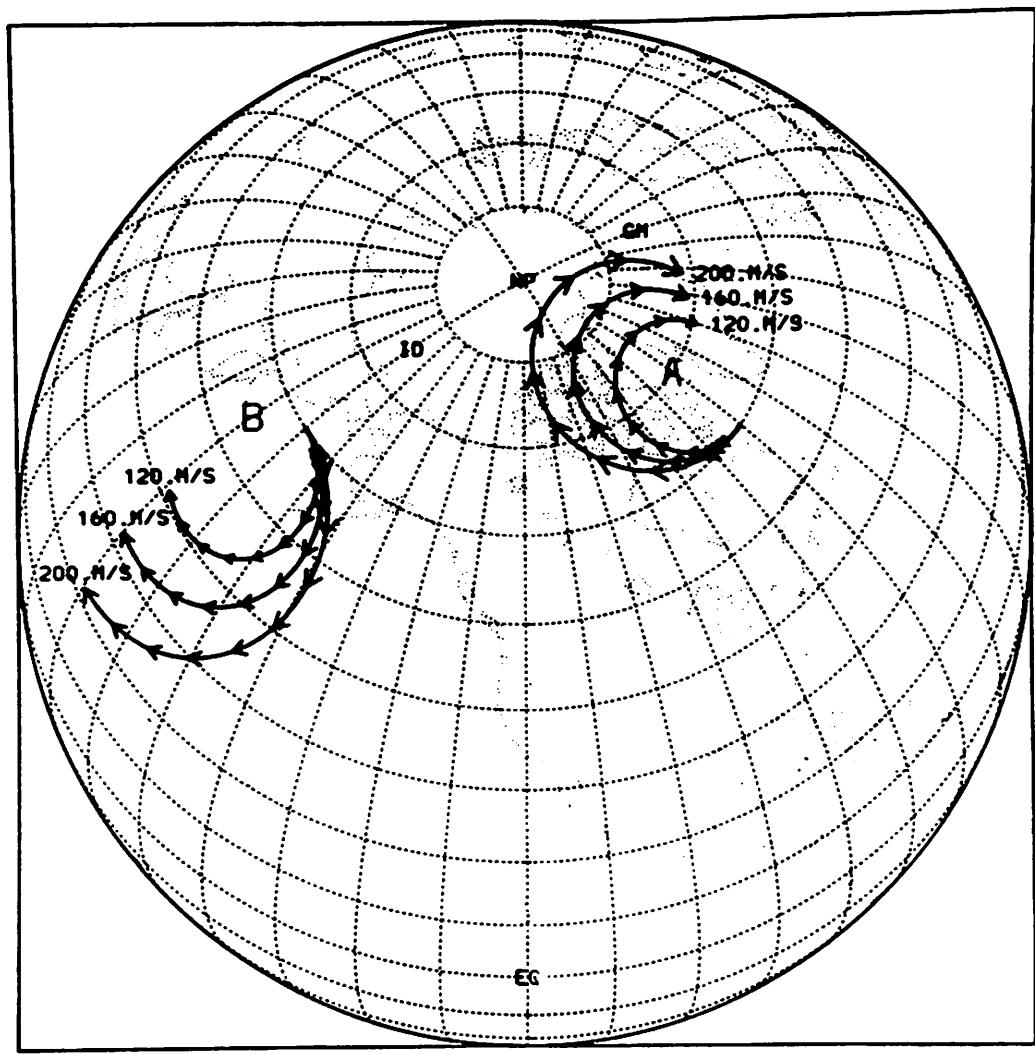


FIG. 2. THE PROGRESS OF PARCELS OF AIR OVER A 9 h PERIOD INITIALLY DIRECTED A—WESTWARD, B—EASTWARD.
(NP = North Pole, GM = Greenwich Meridian, ID = International Date Line.)

Inertial Oscillation: natural motion of
parcels of air

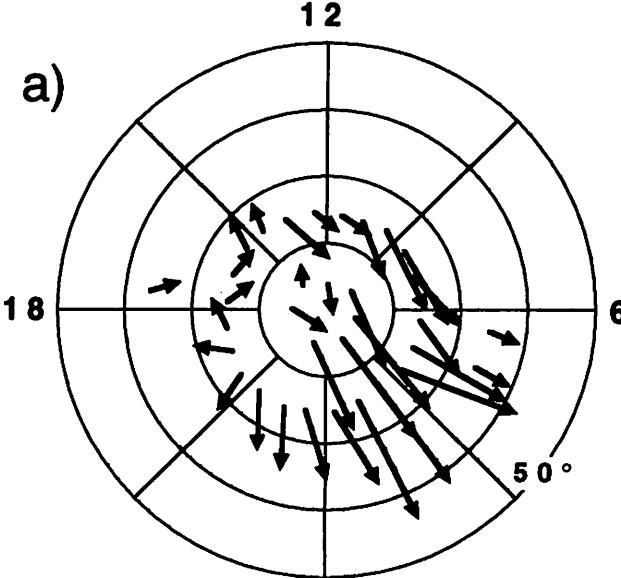
$$\frac{V^2}{r} = f V$$

$$\frac{V}{r} \approx \text{const}$$

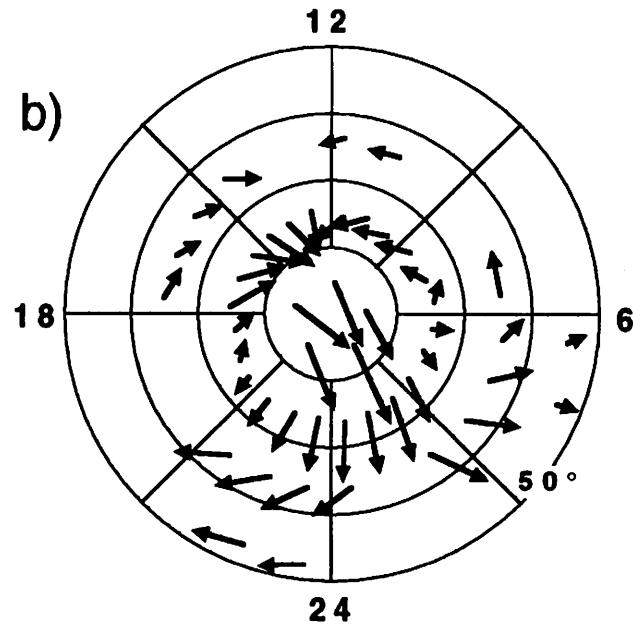
clockwise vortex
in N.H.

Binned winds from DE-2
at ~120 km altitude
compared with model
predictions.
Southern hemisphere
in geomagnetic coordinates

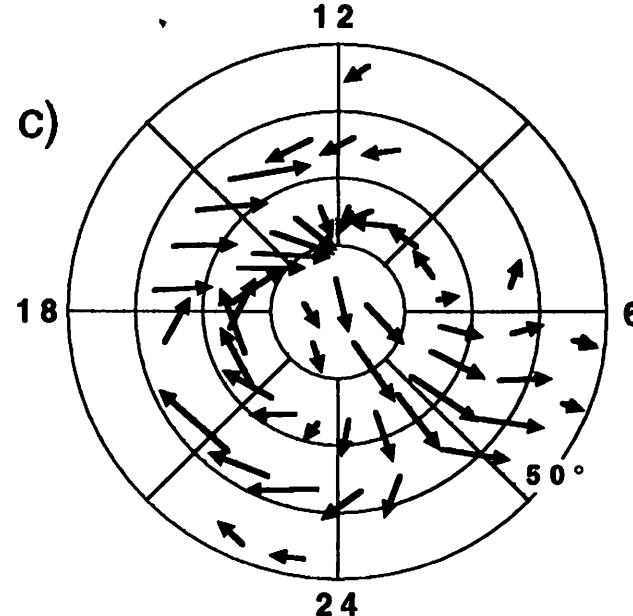
DE-2



NCAR-TGCM



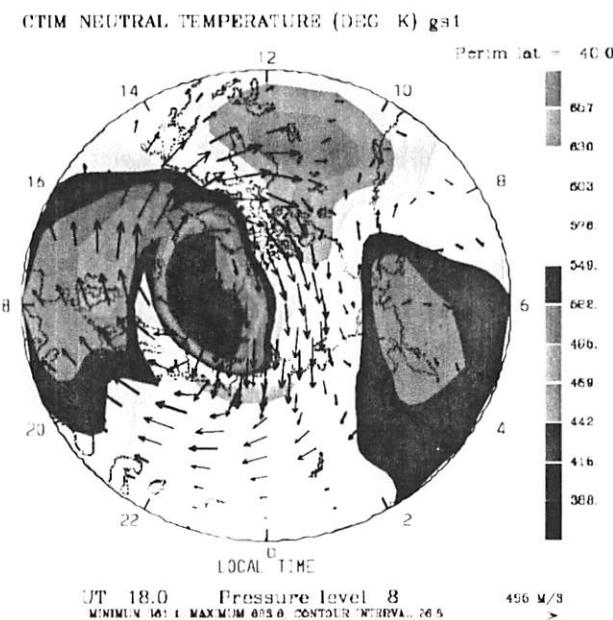
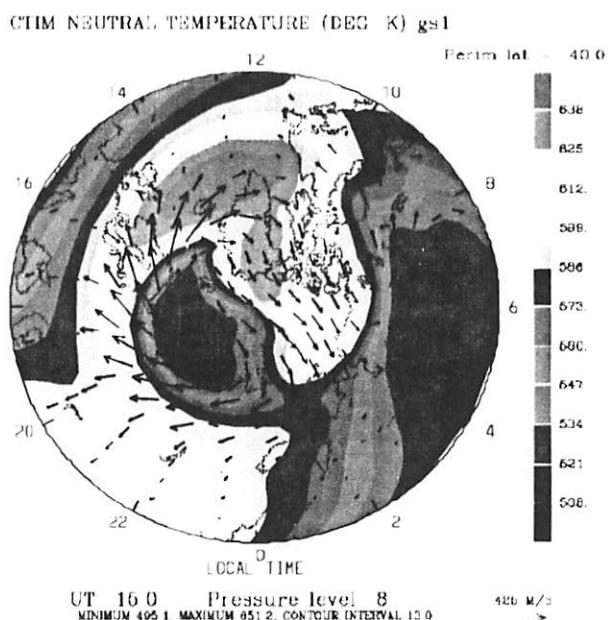
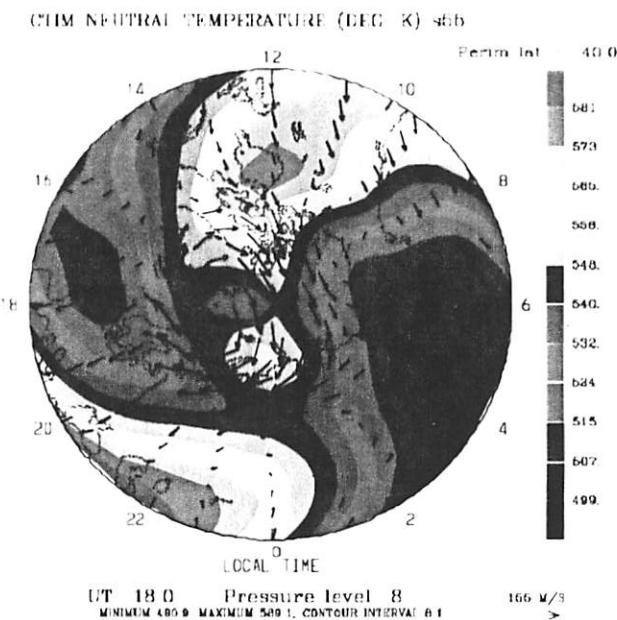
UCL-TGCM



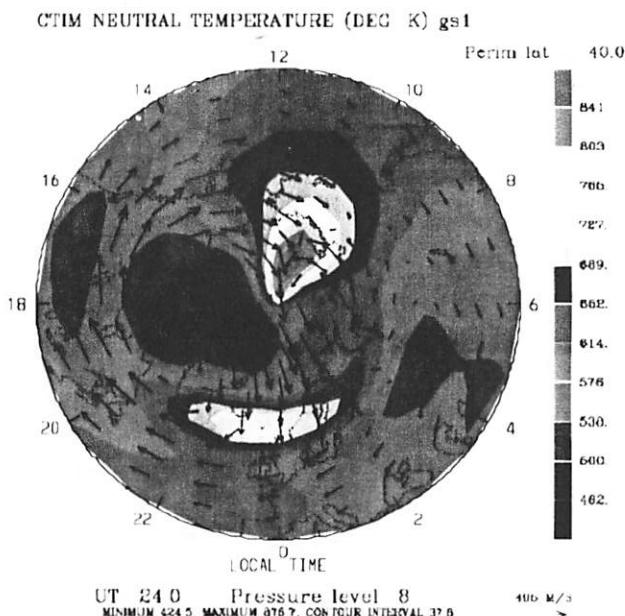
500 m/sec

Killeen et al., GRL, 1992

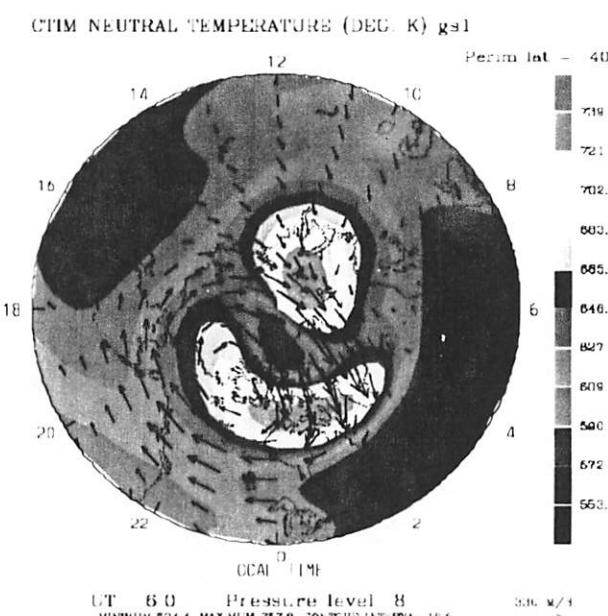
Quick



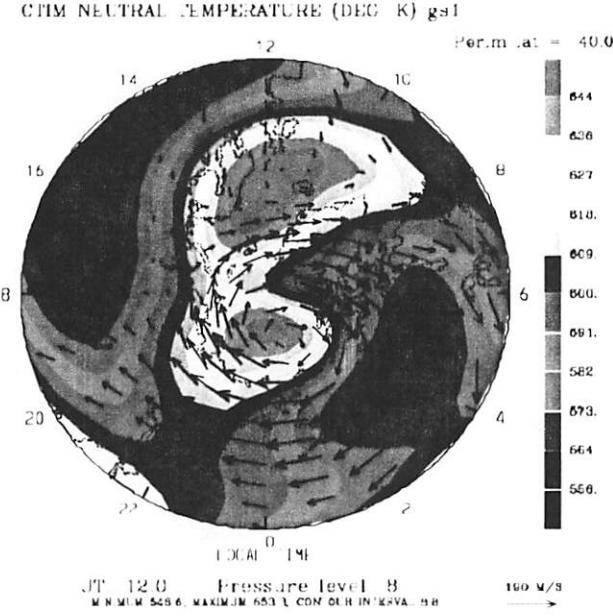
ST 12 hrs



ST 18 hrs



ST 24 hrs.



RECOVERY

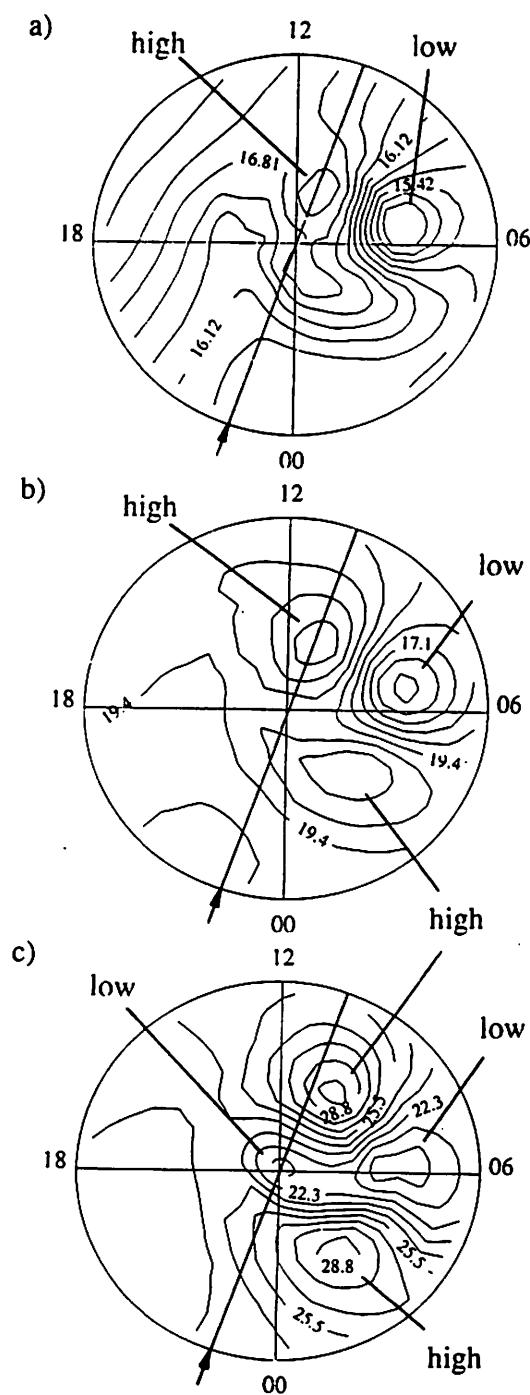
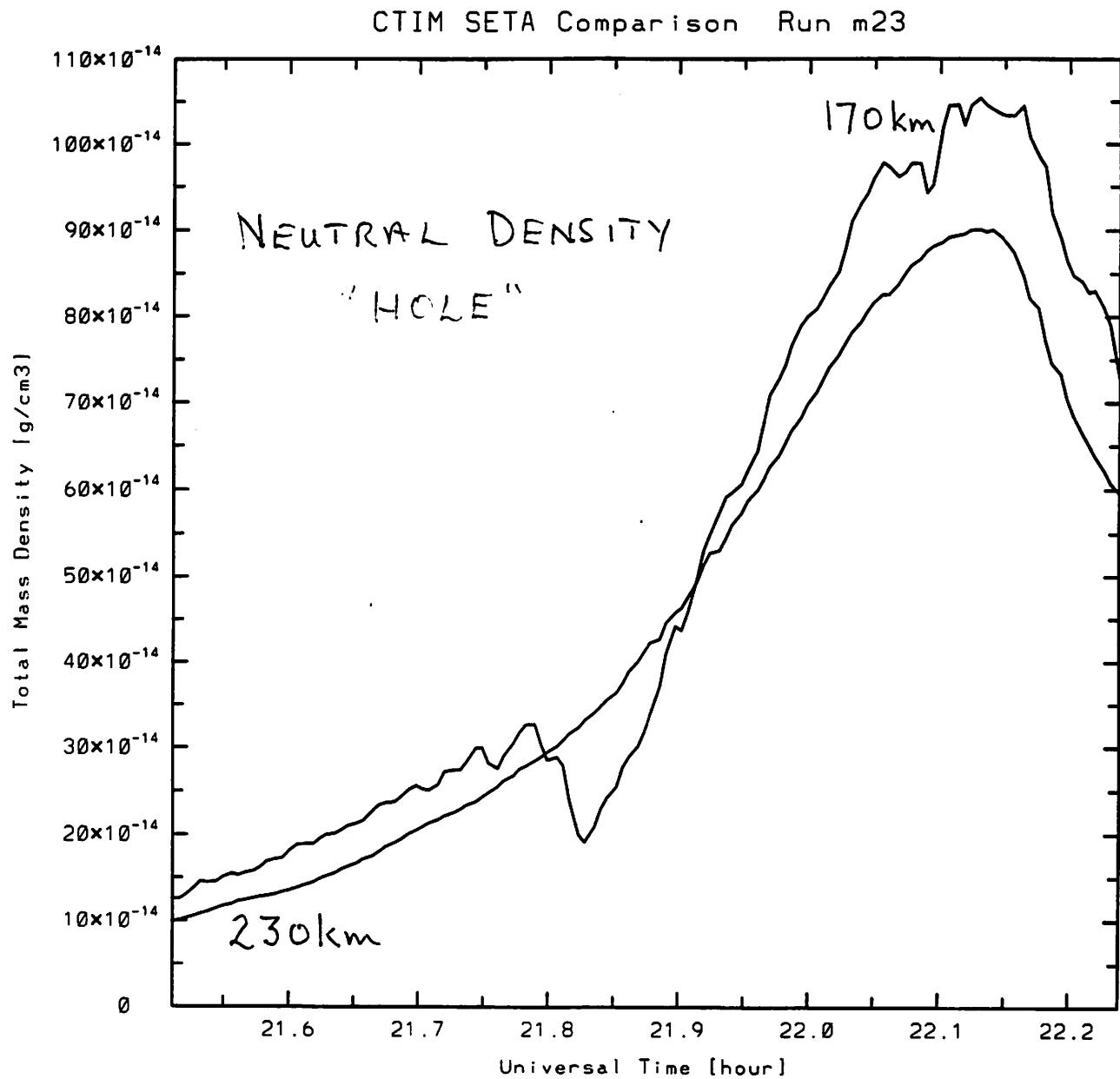
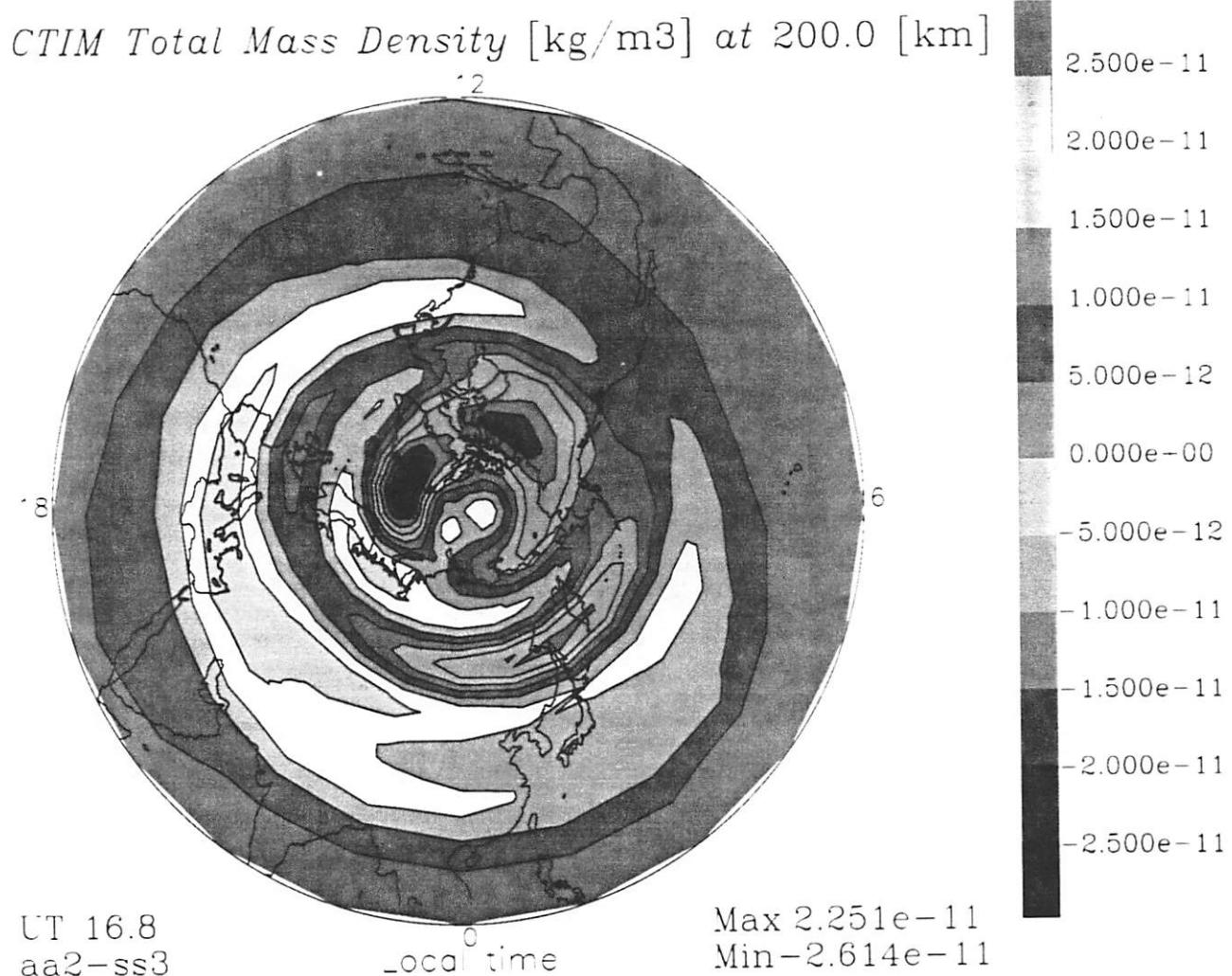


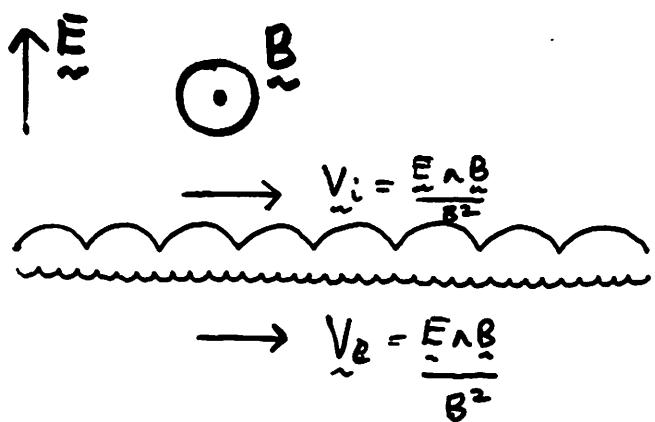
Fig. 1. TIGCM predictions of neutral mass density showing 2-, 3-, and 4-cell patterns in geographic coordinates at 200-km altitude and 12 UT from (a) quiet time (30 kV), (b) moderate (60 kV), and (c) active (90 kV) conditions, respectively. A satellite trajectory in the 2240–1040 Local Time plane is superimposed on each figure. The outer latitude circle corresponds to 45°N. Intermediate latitude circles are omitted for clarity.

MARCOS,

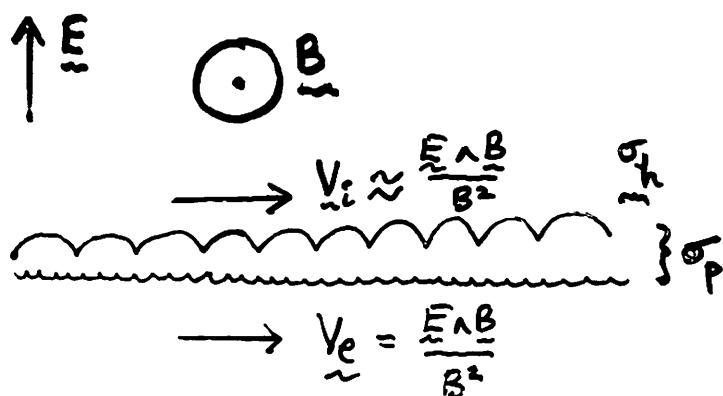


MATSUO ET. AL.

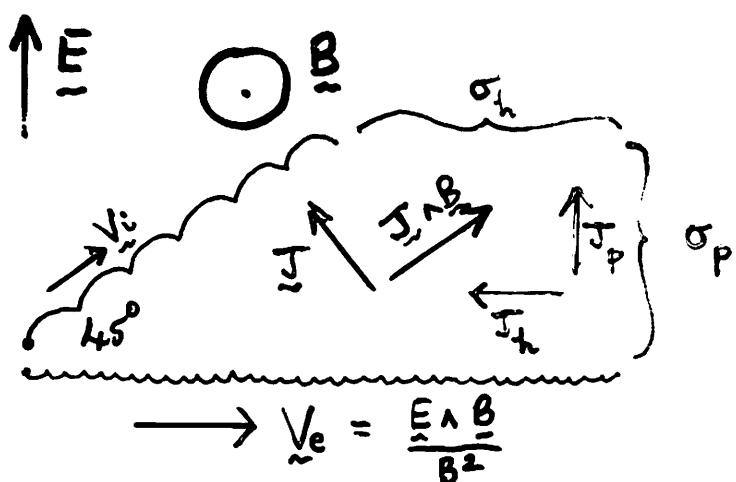




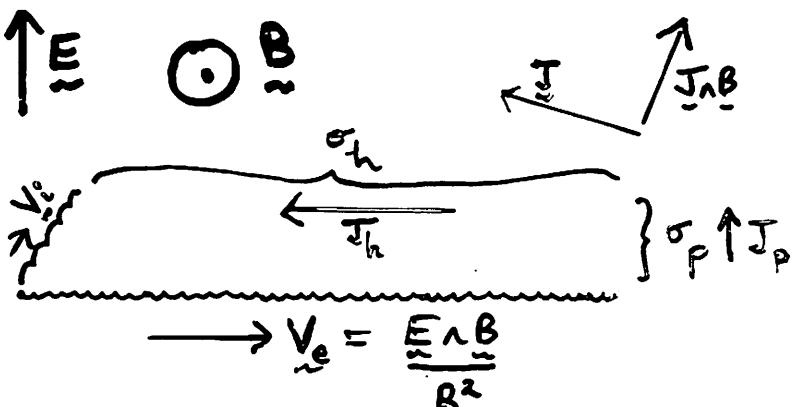
$$\begin{aligned}\underline{V}_{in} &= 0 \\ \underline{V}_i &= \underline{V}_e \\ \underline{J} &= 0 \\ \underline{J} \wedge \underline{B} &= 0 \\ \underline{J} \cdot \underline{E} &= 0\end{aligned}$$



$$\begin{aligned}300 \text{ km } (> 150 \text{ km}) \\ \underline{V}_{in} &\ll \omega_i \\ \underline{V}_i &< \underline{V}_e \\ \sigma_p &> \sigma_h\end{aligned}$$

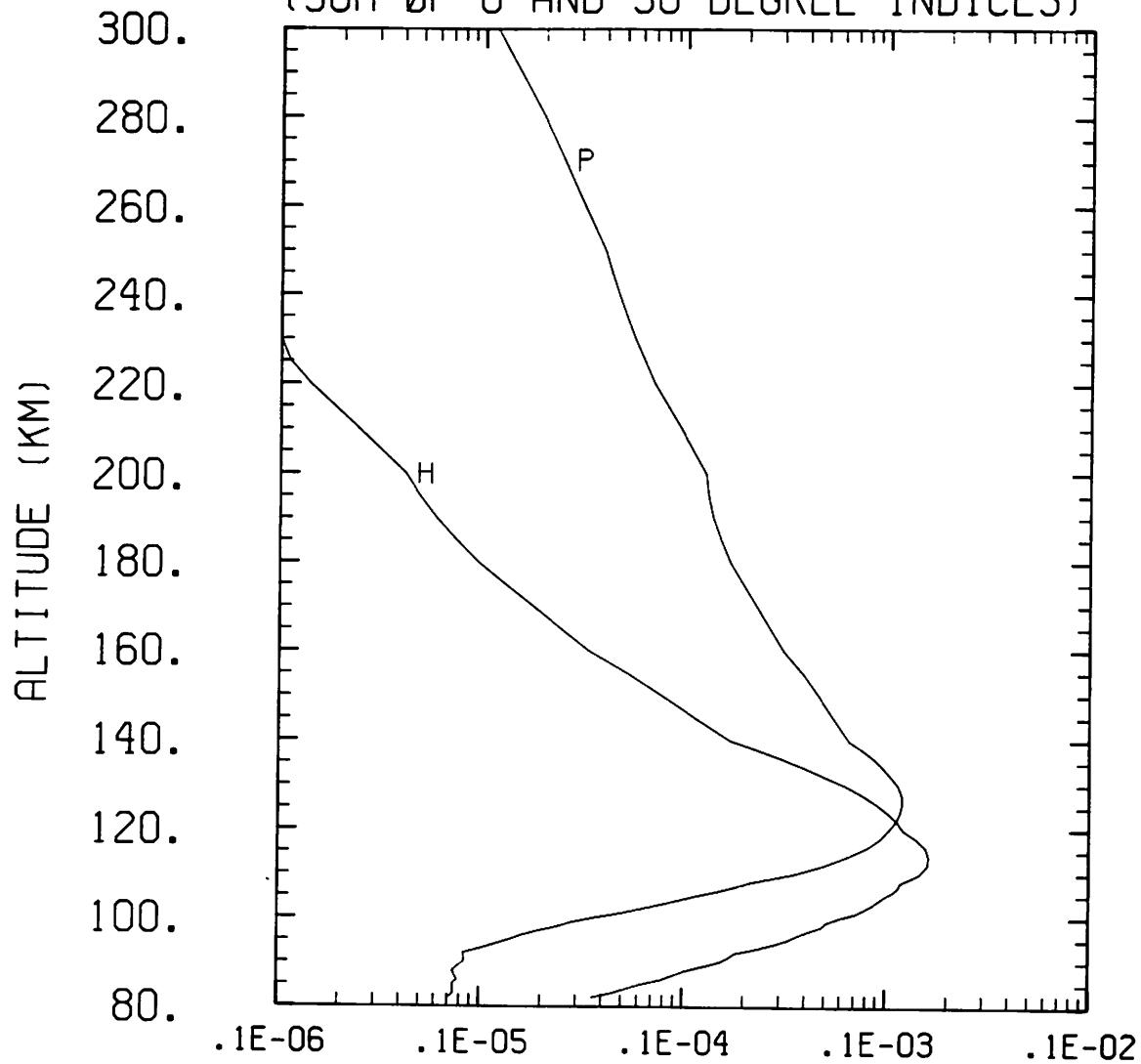


$$\begin{aligned}125 \text{ km} \\ \underline{V}_{in} = \omega_i \\ \underline{V}_i < \underline{V}_e \\ \sigma_p = \sigma_h\end{aligned}$$



$$\begin{aligned}110 \text{ km} \\ \underline{V}_{in} > \omega_i \\ \underline{V}_i \ll \underline{V}_e \\ \sigma_p < \sigma_h\end{aligned}$$

PEDERSEN(P) AND HALL(H) CONDUCTIVITY
 PROFILES PER UNIT FLUX FOR SPECTRA
 WITH PEAK IN BAND 7
 (SUM OF 0 AND 30 DEGREE INDICES)



CONDUCTIVITY (MHOS/M)
 HEIGHT INTEGRATED CONDUCTIVITIES
 HALL = 4.439 MHOS
 PEDERSEN = 5.094 MHOS

$$\begin{aligned}\tilde{\mathcal{T}} &= \tilde{\sigma} \cdot (\tilde{E} + \tilde{V}_n \wedge \tilde{B}) \\ &= \tilde{\sigma} \cdot (E')\end{aligned}$$

Joule Heating and Mechanical Energy Transfer Rates

$$\mathbf{j}(z) \bullet \mathbf{E} = \mathbf{j}(z) \bullet \mathbf{E}' + \mathbf{u}_n(z) \bullet [\mathbf{j}(z) \times \mathbf{B}]$$

$$\left\{ \begin{array}{l} q(z) = q_j(z) + q_m(z) \end{array} \right\}$$

$$q_j(z) = \mathbf{j}(z) \bullet \mathbf{E}' \Rightarrow \text{Joule heating rate} \quad \mathcal{T} \cdot (\underline{\mathbf{E}} + \underline{\mathbf{u}}_n \wedge \underline{\mathbf{B}})$$

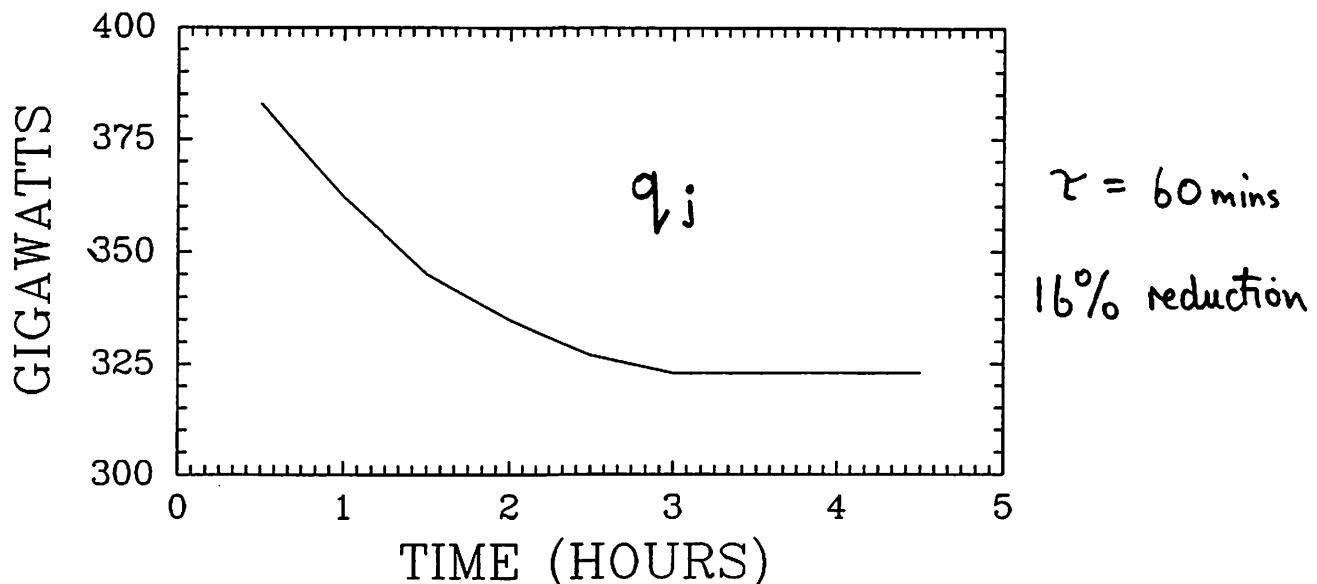
Always positive, q_j serves as a sink of electromagnetic energy

$$q_m(z) = \mathbf{u}_n(z) \bullet [\mathbf{j}(z) \times \mathbf{B}] \Rightarrow \text{Mechanical energy transfer rate}$$

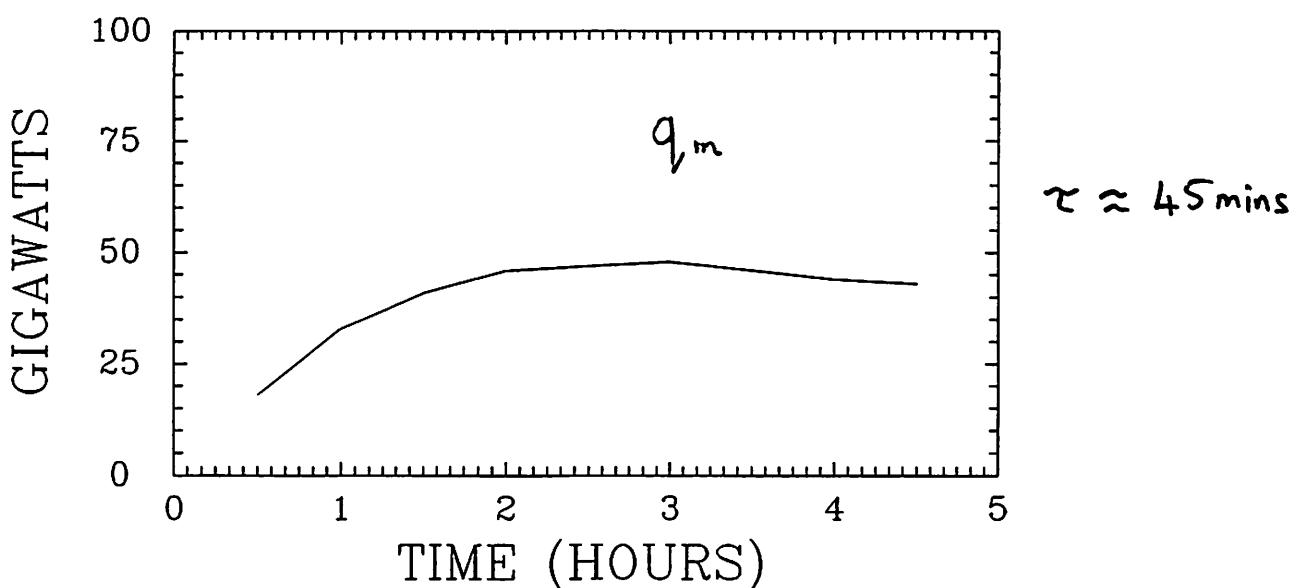
If negative, q_m serves as a source of electromagnetic energy

STORM JOULE HEAT DISSIPATION
DECEMBER SOLSTICE

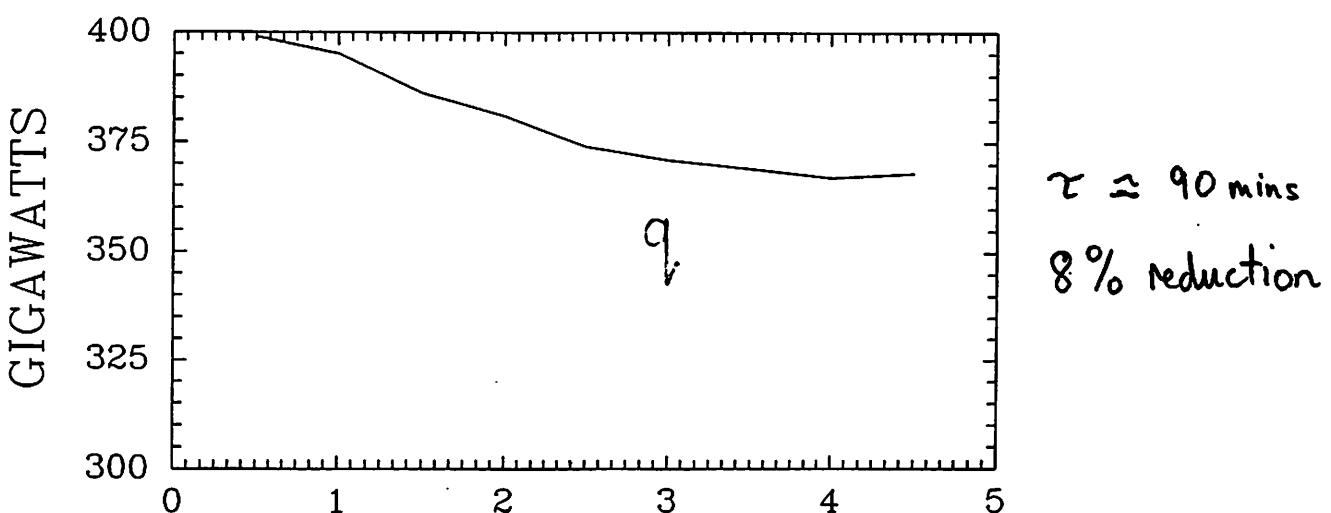
(one hemisphere)



STORM KINETIC DISSIPATION
DECEMBER SOLSTICE



TOTAL STORM DISSIPATION
DECEMBER SOLSTICE

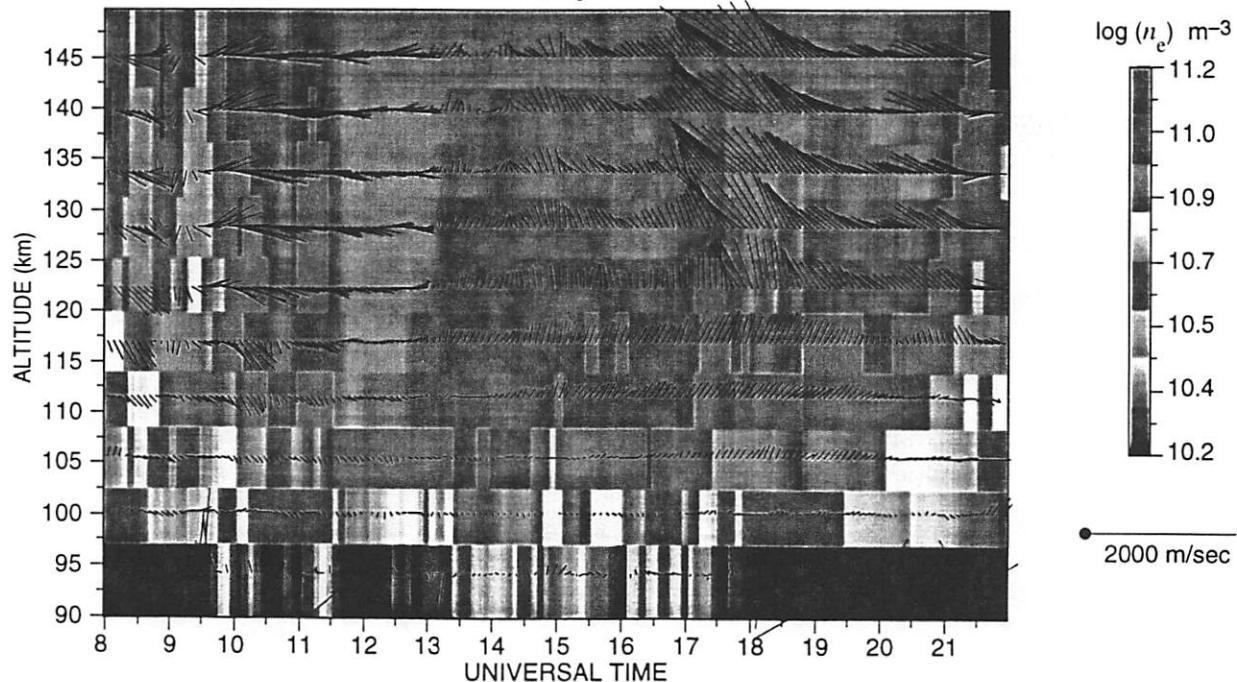




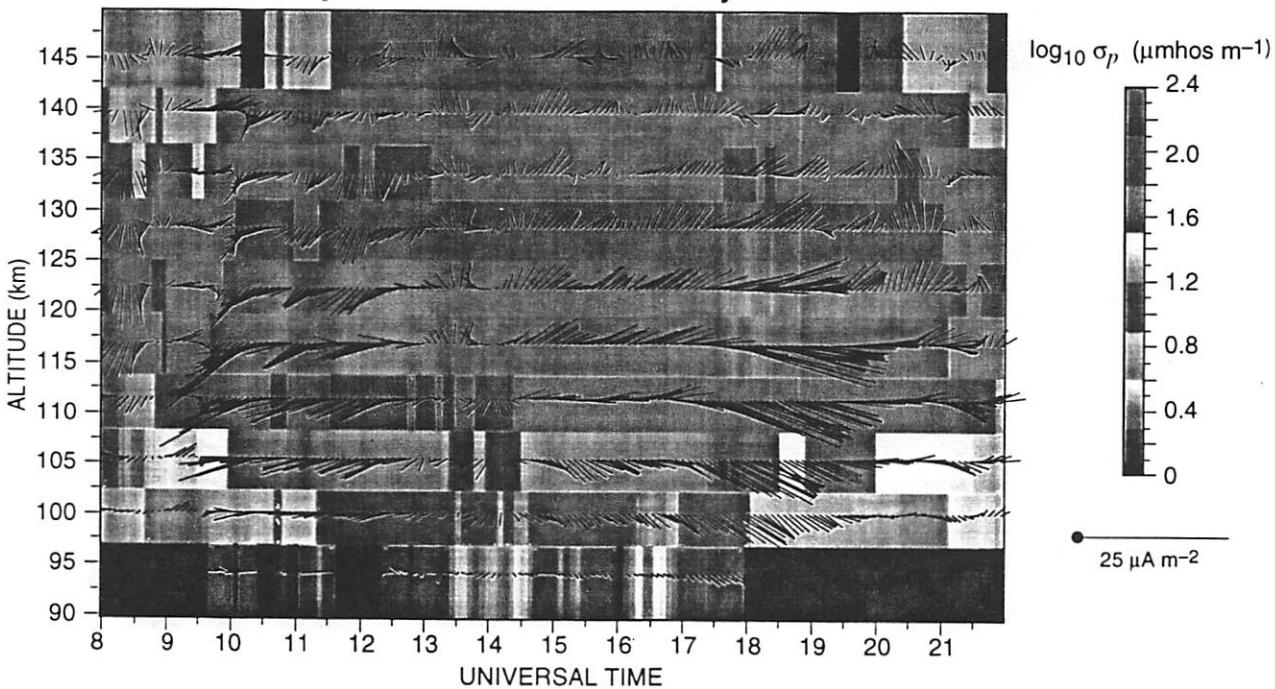
Thayer and Heelis

E-REGION ELECTRODYNAMIC PARAMETERS FOR AUGUST 5, 1993

Ion Velocity and Electron Density



Current Density and Pedersen Conductivity

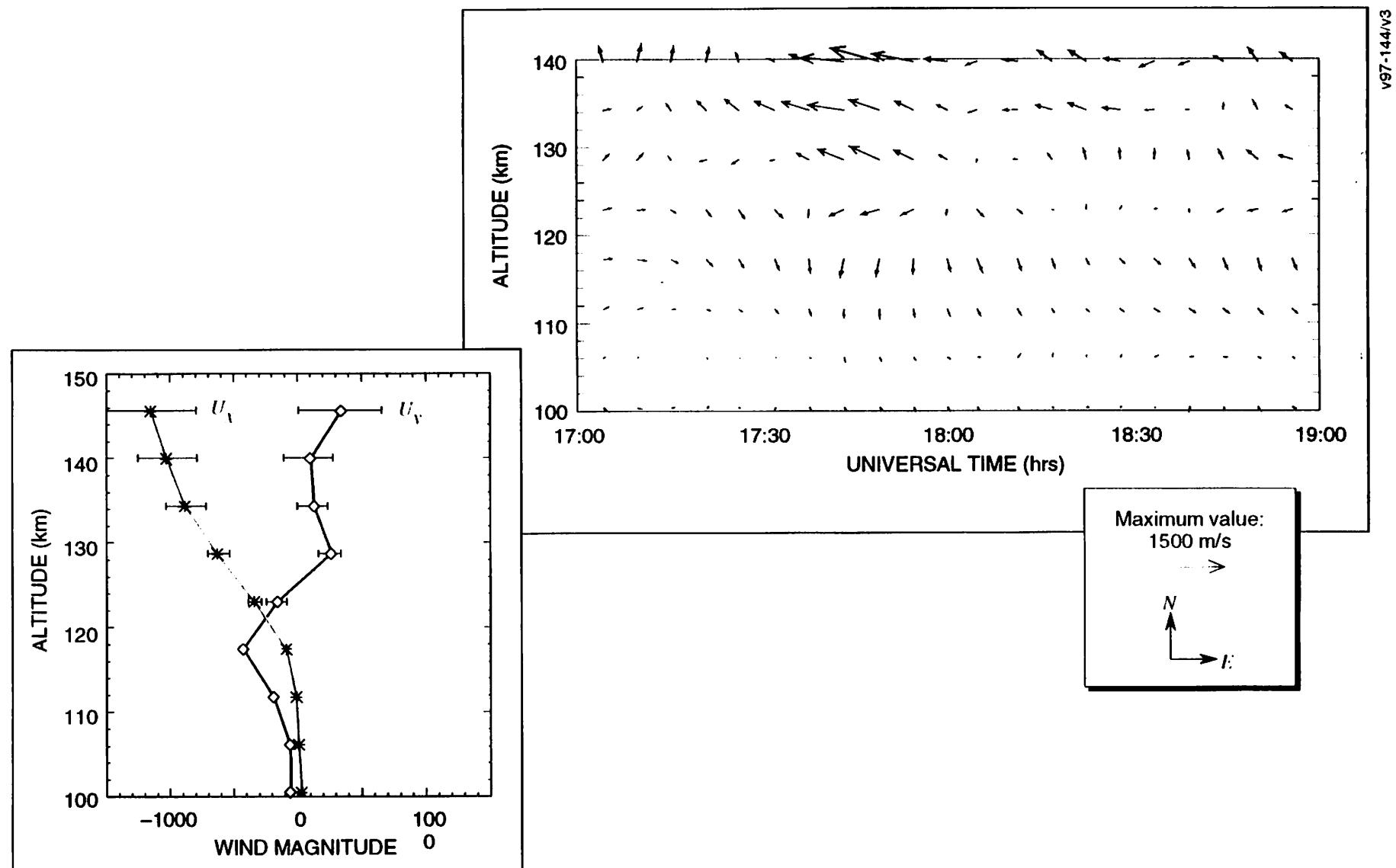


$$\sigma_p = \frac{en_e}{|B|} \left[\frac{\Omega_i v_{in}}{\Omega_i^2 + v_{in}^2} \right]$$

$$j_{\perp} = en_e [V_{i\perp} - V_{e\perp}]$$

Thayer & Heelis

Radar-Derived Neutral Winds (August 5, 1993)





Thayer & Heelis

E-REGION ELECTRODYNAMICS FOR AUGUST 5, 1993

