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Tutorial Lecture

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Winds and Composition in the Thermosphere

Winds & Composition in the Thermosphere

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Mostly about large-scale structure
in the mid-latitude E & F layers

- 1 General Principles
- 2 Wind systems
- 3 Composition Effects — Quiet Days
- 4 — Storms
- 5 More general principles

FIRST PHYSICAL THEORY OF THE IONIC SPHERE ?

Re Mr Marconi's Results in Day and Night Wireless Telegraphy

CLIVER LODGE [1902] : "The observed effect, which if confirmed is very interesting, seems to me to be due to the conductivity ... of air, under the influence of ultra-violet solar radiation."

"No doubt electrons must be given off from matter ... in the solar beams; and the presence of these will convert the atmosphere into a feeble conductor."

THE KEY YEARS

- 1931 Regular Ionospheric Sounding began at Slough
- 1931 Chapman's Theory of Ionized Layers
- 1931 Chapman's Theory of Airglow
- 1931 Chapman-Ferraro: Solar Particles & the Earth's magnetic field
 - First idea of Magnetosphere?
- 1932 Appleton-Hartree Magneto-Ionic Equn
- 1932 Start of K_p Geomagnetic Index
- 1933 Slough Swept-Frequency Ionograms

NEW UNDERSTANDING OF THE IONOSPHERE

- ¶ GLOBAL CIRCULATION in the upper atmosphere—
driven by SUN and MAGNETOSPHERE/SOLAR WIND
- * Chemical changes → F layer seasonal effects [D layer]
- * Temp. variations → winds/waves → F layer effects
- * Auroral heating → winds → F layer storms [D layer]
- * Electrodynamics → equatorial E & F layers
- ¶ MAGNETOSPHERE — links with ionosphere and IMF
- * New focus on the high latitude ionosphere :
[polar caps, auroral zone, "troughs"]

Ideas :

- "Vertical structure ↔ solar radiations"
- "Latitude structure ↔ geomagnetic field"
- "Ionosphere as part of interactive solar-terrestrial system"

IONOSPHERIC CONSERVATION EQUATIONS

Continuity equation [mass]:

$$\{\text{Density change}\} = \{\text{Production}\} - \{\text{Loss}\} - \{\text{Transport}\}$$

Equation of motion [momentum]:

$$\{\text{Acceleration}\} = \{\text{Force}\} - \{\text{Drag}\} - \{\text{Transport}\}$$

Heat equation [energy]:

$$\{\text{Temp change}\} = \{\text{Heating}\} - \{\text{Cooling}\} - \{\text{Conduction}\}$$

Plus:

Equation of state [perfect gas equation]

F2 Layer continuity equation

Plasma diffusion
Neutral-air winds
Electric fields

$$\begin{aligned}\partial N/\partial t &= q - \beta N - \operatorname{div}(N\mathbf{V}) \\ &= q - \beta N - \mathbf{V} \cdot \nabla N - N \operatorname{div} \mathbf{V}^*\end{aligned}$$

"Diffusion controls height of peak"

"Main effect of vertical drift is to move the plasma to a different q/β "

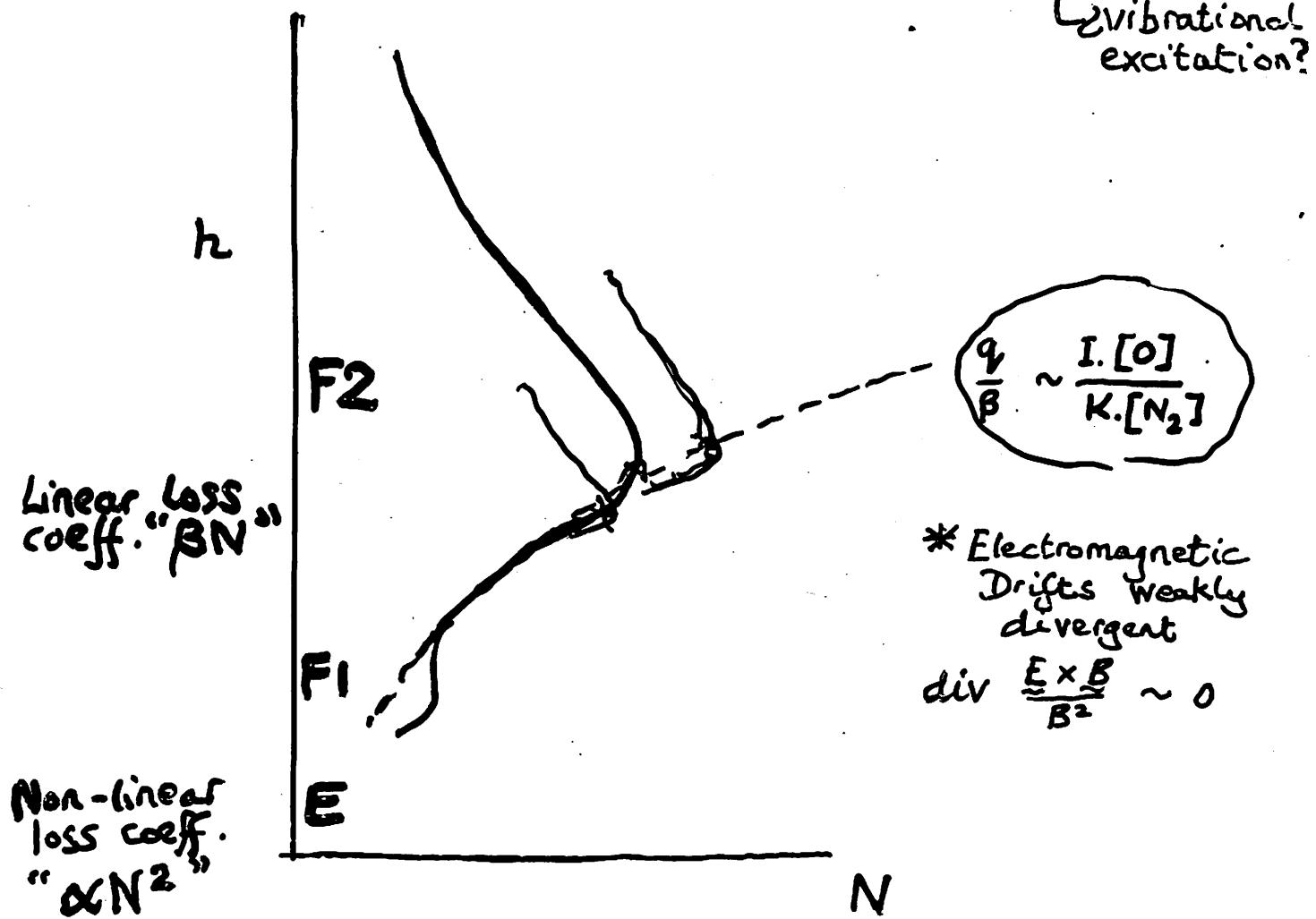
"Vertical drift \leftarrow Horiz Winds"

\uparrow
Motion of existing gradient
 \uparrow
Velocity divergence

$$q \propto [O]$$

$$\beta \propto [N_2, O_2]$$

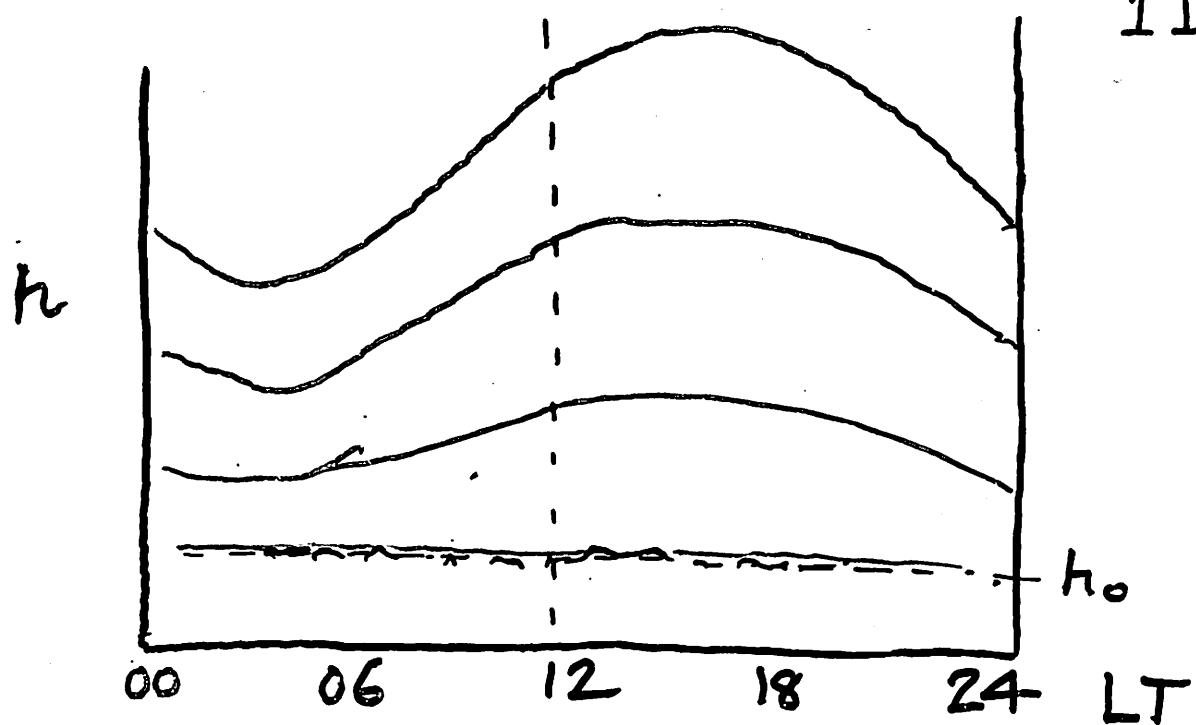
\uparrow
vibrational excitation?



Thermal Expansion

"Barometric Motion"

1D



$$Z = \frac{h - h_0}{H} = \int_{h_0}^h \frac{dh}{H} \leftarrow \frac{RT}{Mg}$$

[O/N₂]

If \downarrow Compo is fixed at lower boundary h_0
then it remains fixed at any fixed Z

Can be change by
 * Vertical circulation (horiz divergence) "up" (convergence) "down" welling

* Change of mixing at tropopause ?

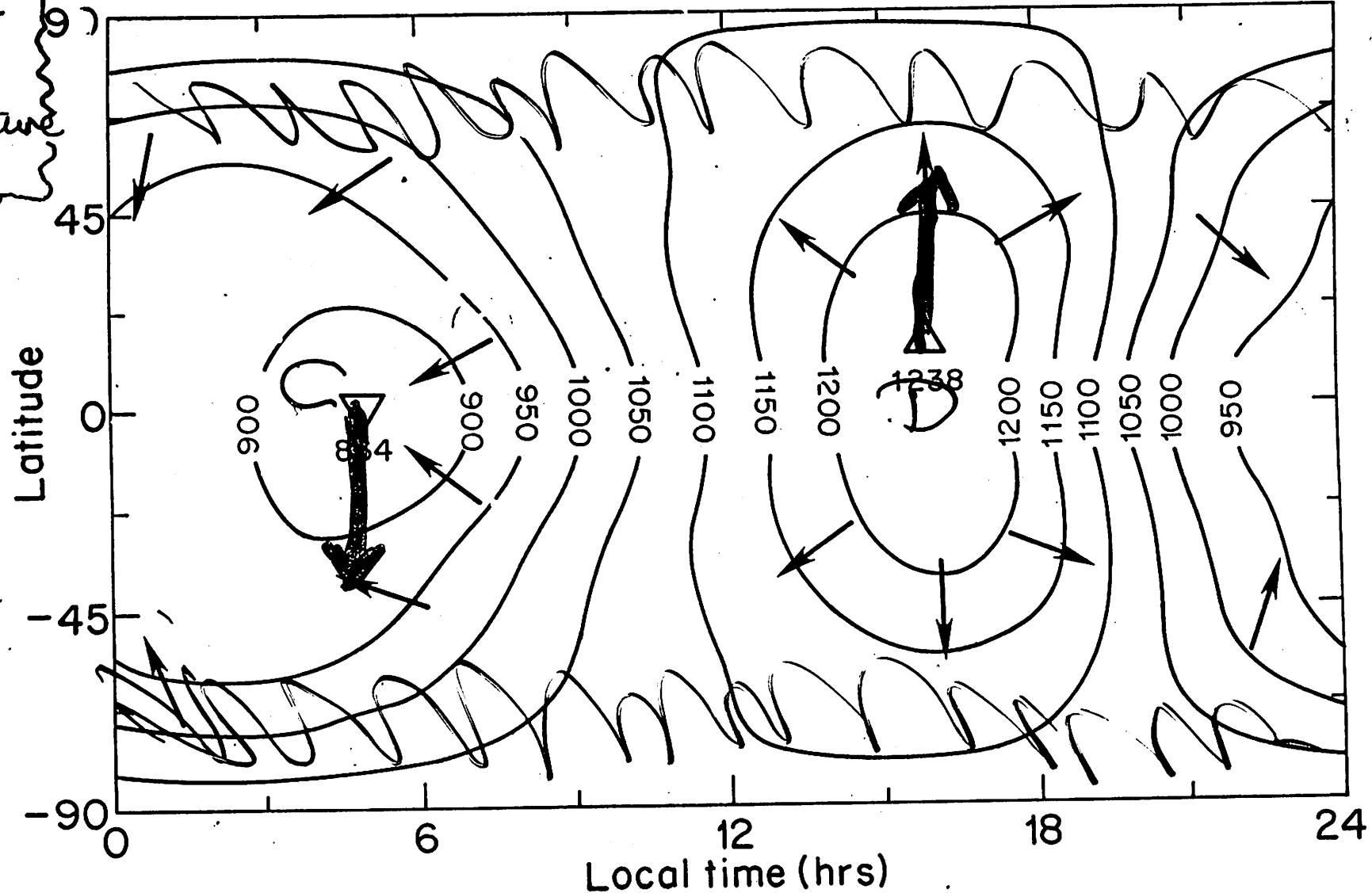
* Gravity waves ???

300 km
Wind &
Temperature
Map

MSIS 83 : MARCH

$S_{1C.7} = 150$

$A_p = 4$



112.11

Winds: due to : Solar Heating ("Thermal tide"?)
Gravity waves, tides from below Magnetospheric Heating (joule heating, particle ppt)

Longitude Averaged

10 I SOLAR VARIABILITY: A SUMMARY

EQUINOX

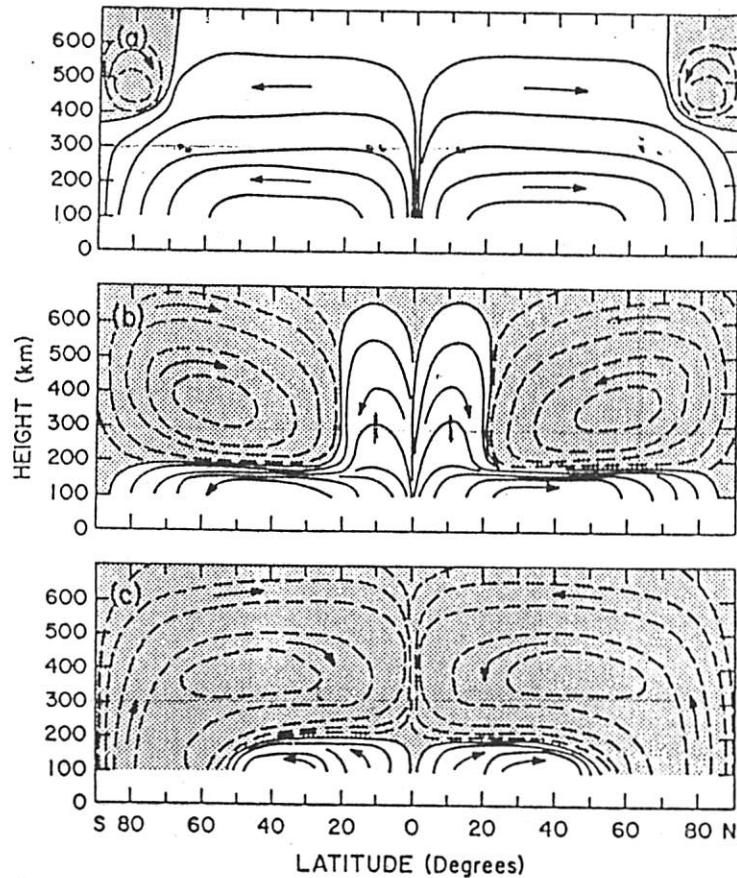


Fig. 1 Schematic diagram of the mean thermospheric circulation during equinox. The upper panel shows the circulation during quiet geomagnetic conditions, the middle panel is for average conditions, and the bottom panel is for geomagnetic storms.

Thermospheric Circulations

A Summary

11

SOLSTICE

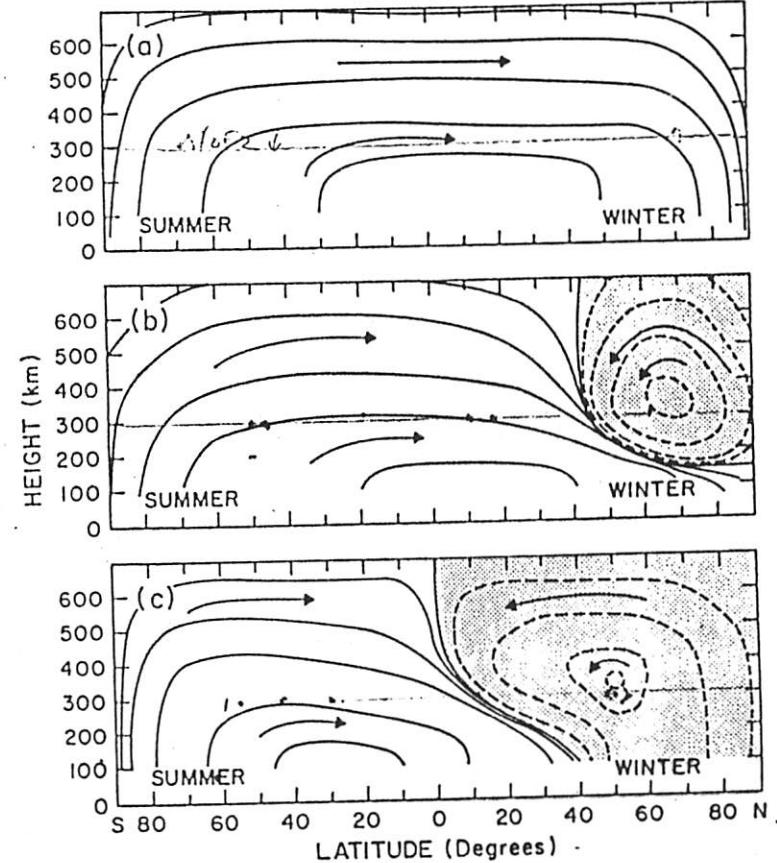


Fig. 2 Schematic diagram of the mean thermospheric circulation during solstice. The upper panel shows the circulation during quiet geomagnetic conditions, the middle panel is for average conditions, and the bottom panel is for geomagnetic storms.

Composition Changes & Vertical Winds

O/N₂ ratio is decreased by upwelling of air
increased by downwelling of air

Vertical motion \leftrightarrow energy input / release

Vertical
wind
(neutral air)

$$W = W_B + W_D$$

"BAROMETRIC" 1-dim¹ expansion/contraction

$$W_B = \left(\frac{\partial h}{\partial t} \right)_P \quad \begin{array}{l} \text{(Does not change composition)} \\ \text{at a fixed pressure-level} \end{array}$$

"DIVERGENCE" 3-dim¹ circulation (linked to horiz. wind
 (by continuity eqn)

$$W_D = - \frac{1}{\rho g} \cdot \frac{DP}{DT} \quad \begin{array}{l} \text{Motion of air through} \\ \text{pressure levels} \end{array}$$

To specify composition of thermosphere

$$P = 28 \ln [O] - 16 \ln [N_2] - 12 \ln T$$

* Independent of height if thermosph. is in diffusive equilibrium

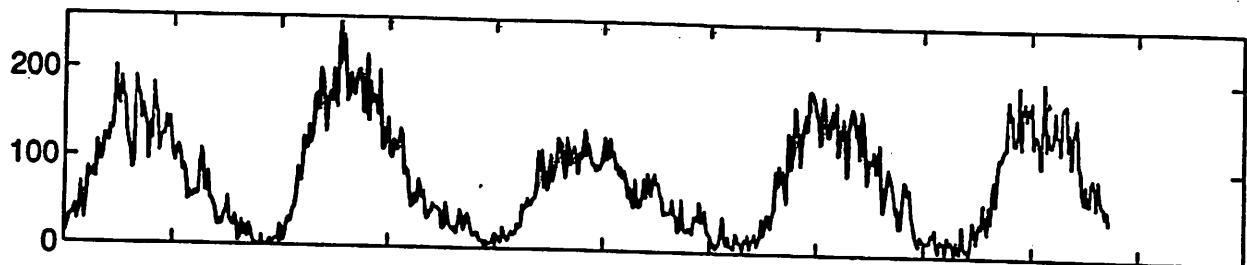
* $\partial P / \partial h \leftrightarrow W_D$

* Horiz. variation \leftrightarrow "wind induced diffusion" \rightarrow global circulation

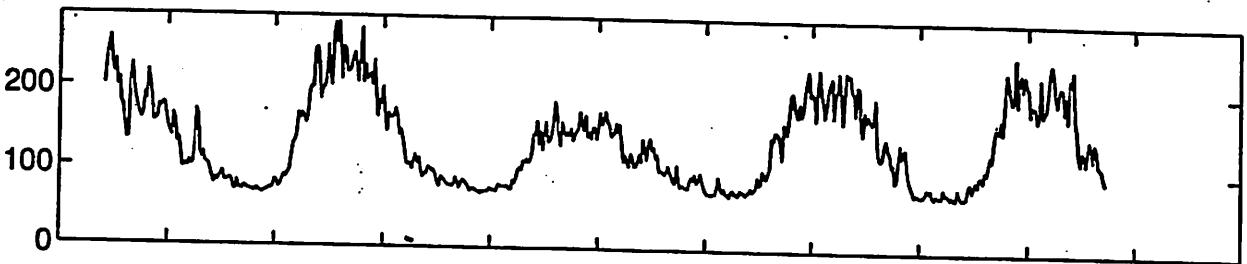
* "Think pressure-levels": $Z = \int_{h_0}^h \frac{dh}{F}$

A HALF-CENTURY OF SOLAR & IONOSPHERIC DATA

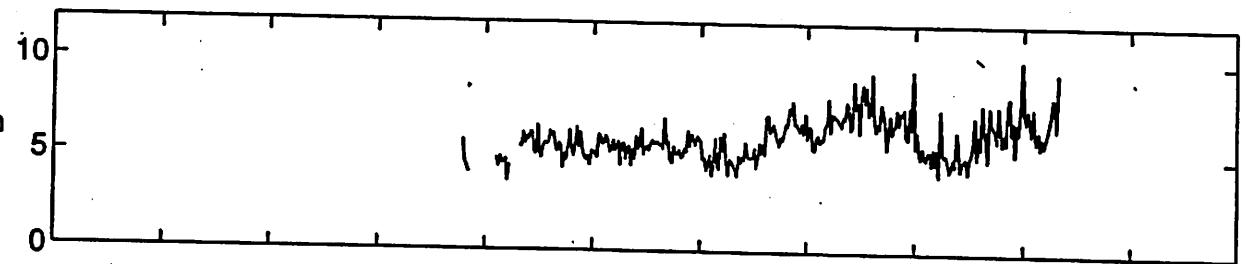
1749



F10.7

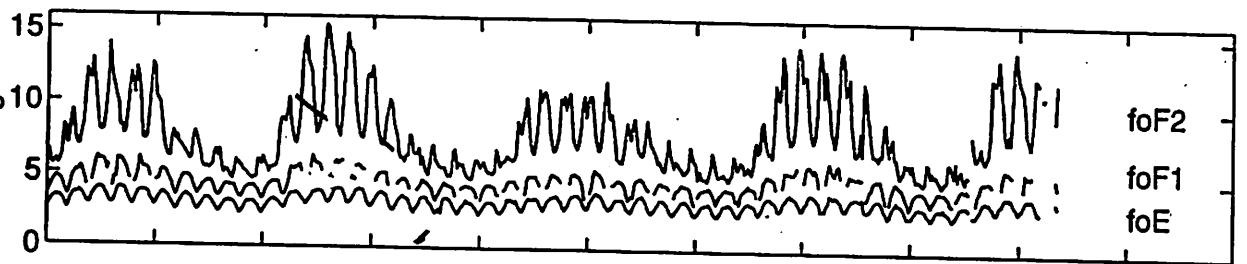


B

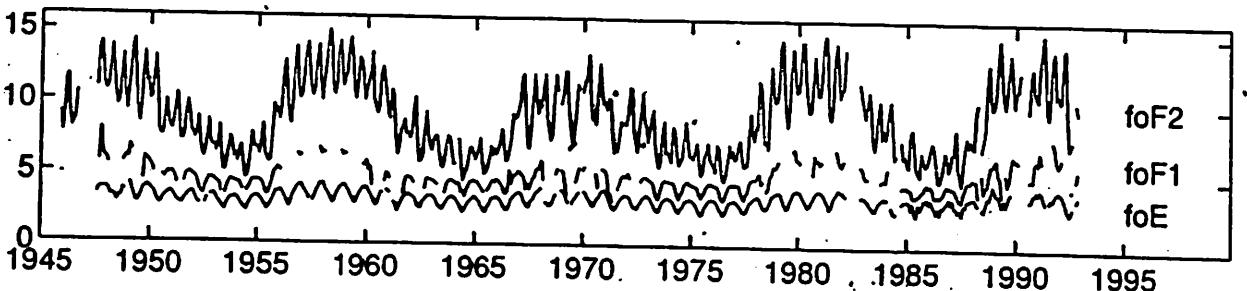


1931

Slough

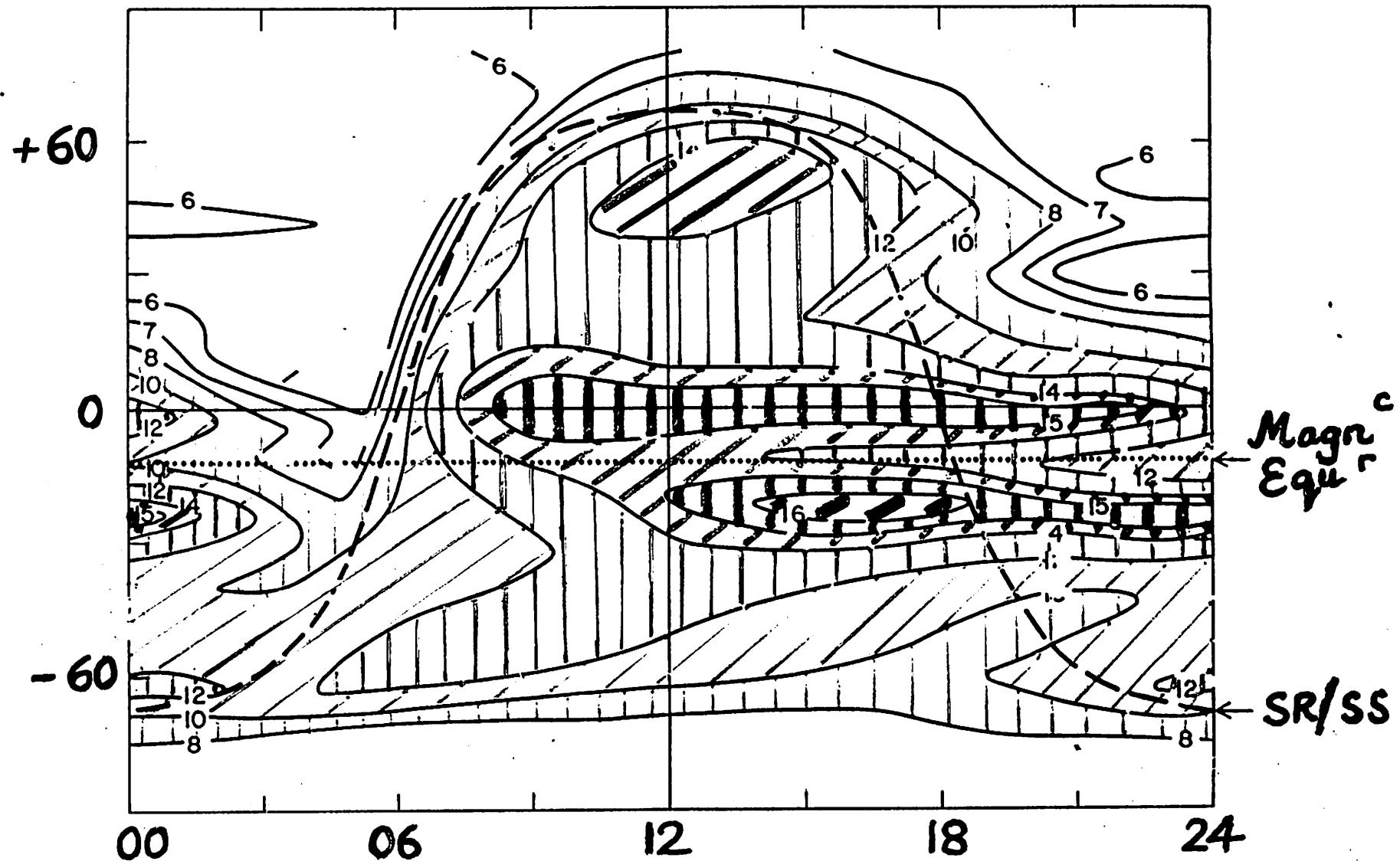


Port Stanley



IONOSPHERIC F2-LAYER MAP FOR AMERICAN SECTOR

*f*_xF2 in Latitude & Local Time : Dec 1957



F2-LAYER ANOMALIES at MID-LATITUDES

“Winter” or “Seasonal”

Greatest NmF2 in winter

“Annual”

Greatest NmF2 in December

“Semiannual”

Greatest NmF2 & hmF2 at equinox

PHYSICAL CAUSES

Neutral composition changes

[O/N₂] ratio or MMM

Horizontal neutral-air winds

Vertical ion drift $U \cos I \sin I$

Solar zenith angle

Ion production $\int q dh \propto \cos \chi$

NEUTRAL COMPOSITION is controlled by vertical air motion

“Upwelling” \Rightarrow Decreased [O/N₂] ratio, increased MMM

“Downwelling” \Rightarrow Increased [O/N₂] ratio, decreased MMM

*Greatest downwelling at $\sim 5^\circ$ equatorward of dayside auroral oval
(Millward et al.)*

EXPLANATION of Seasonal & Semiannual Anomaly :

Solar-driven circulation \Rightarrow seasonal changes of $[O/N_2]$ ratio

Summer midlatitudes : Upwelling \Rightarrow Decreased $[O/N_2]$

Winter midlatitudes : Downwelling \Rightarrow Increased $[O/N_2]$

Greatest downwelling at about 5° equatorward of auroral oval

Horiz wind U is weak near downwelling, strong at lower latitudes

Winter at $\sim 50^\circ$ geog: sectors near magnetic poles:

Atlantic
Australasian } sectors

Max downwelling & compo effect - close to auroral oval

Wind effect moderate : U small, I large, $\cos I \sin I \sim 0.25$

{Compo} $>$ {zenith angle} $\Rightarrow NmF2$ (winter) $>$ (equinox) $>$ (summer)

SEASONAL
↑

Winter at $\sim 50^\circ$ geog: sectors remote from magnetic poles:

E. Asian
S. American } sectors

Downwelling & compo effect smaller - far from auroral oval

Wind effect strong : U large, I moderate, $\cos I \sin I \sim 0.5$

{Compo} $<$ {zenith angle} $\Rightarrow NmF2$ (winter) $<$ $NmF2$ (equinox)

SEMIANNUAL
↑

Torr & Torr, JATP 1973

Max Noon foF2

1958 (a)

241

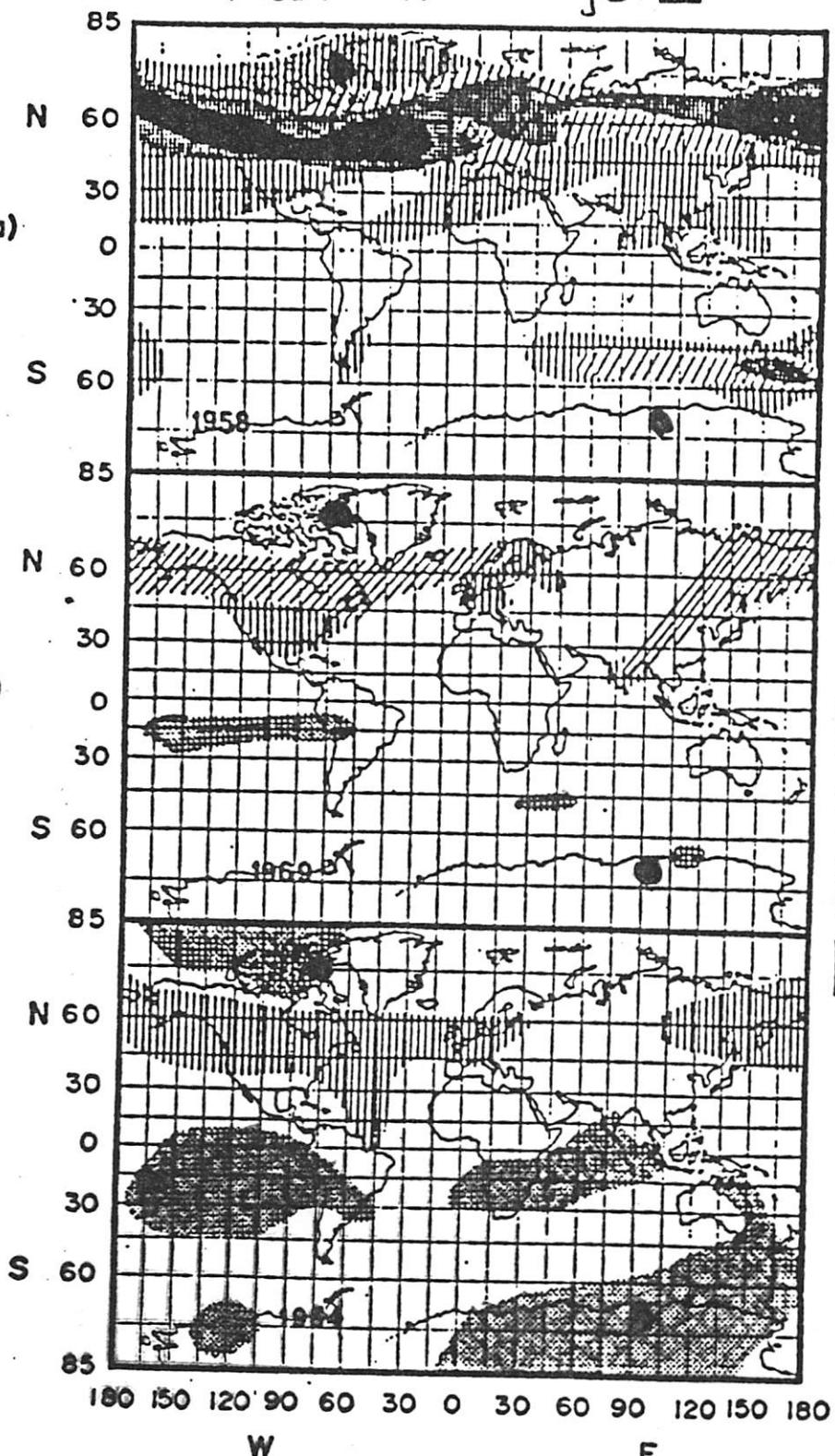
1969 (b)

158

1964(c)

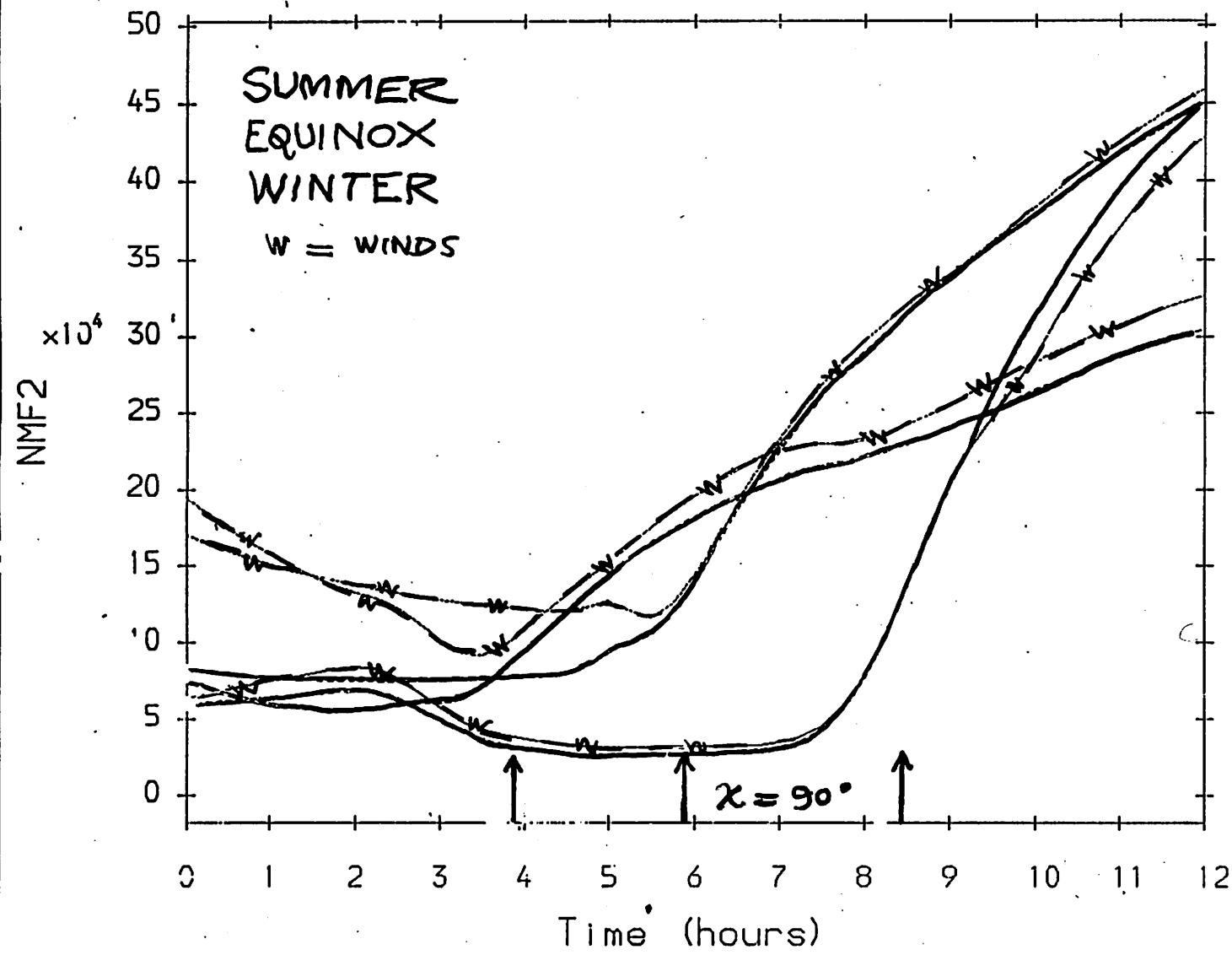
65

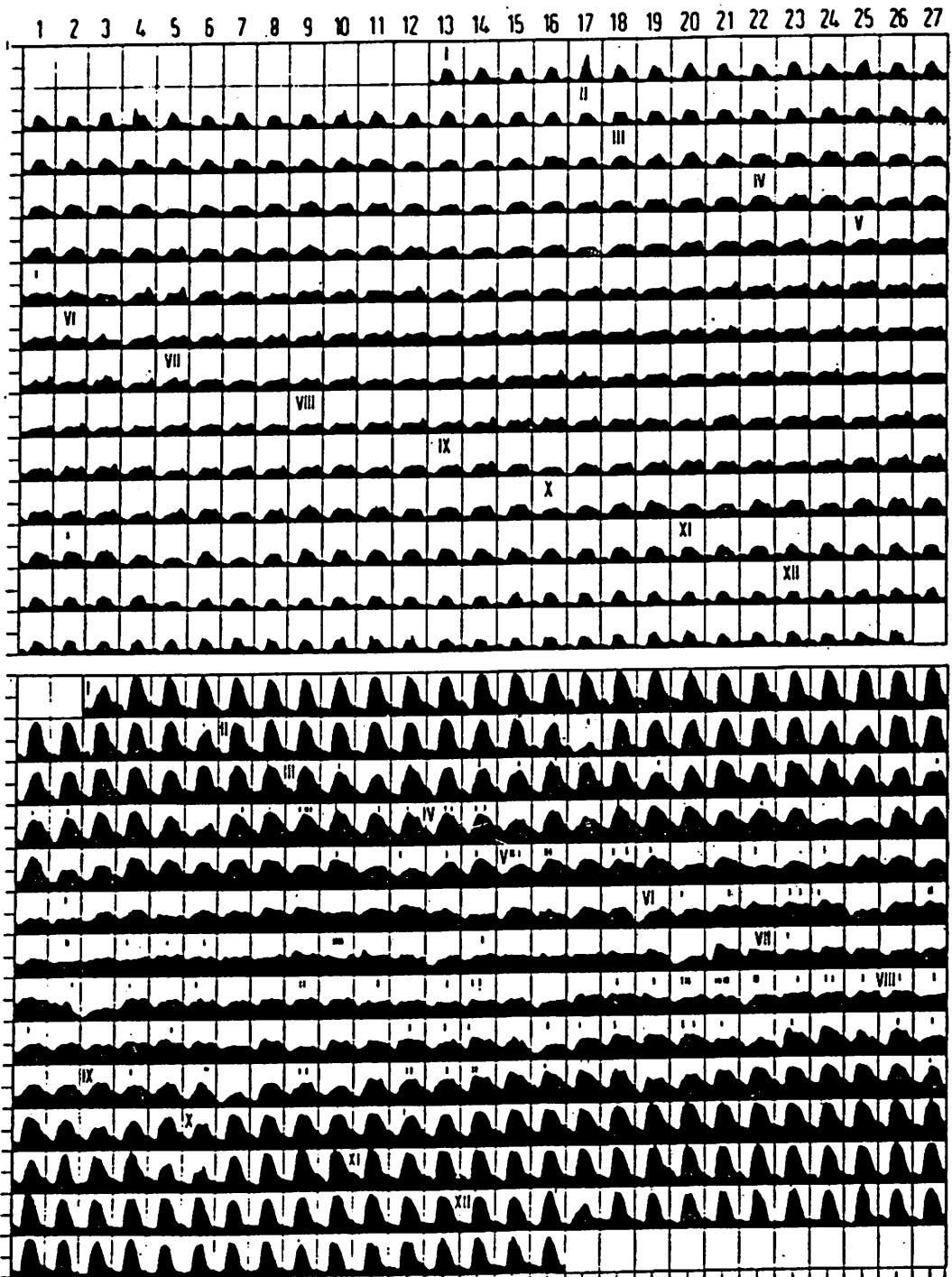
F10.7



SUPIM NmF2

SOLAR
MIN





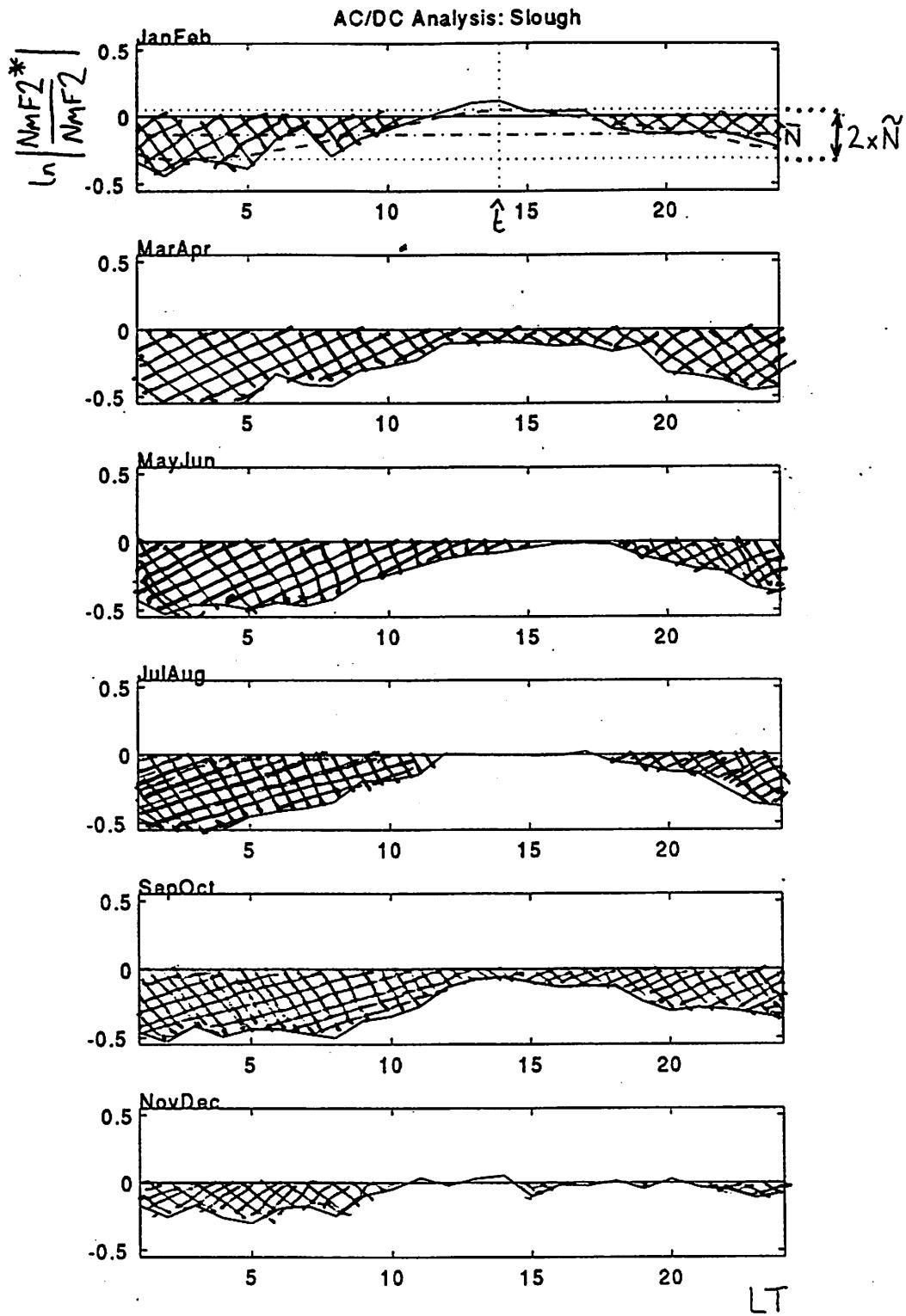
Critical Frequency f_{oF2} at Lindau $52^{\circ}N$

for sunspot minimum & maximum

The boxes are 1-16 MHz \int 00-24 LT
 \longleftrightarrow

$$\ln \left(\frac{NmF2^*}{NmF2} \right) = \bar{N}_{DC} + \tilde{N} \cdot f(t - \hat{t})$$

STORM QUIET AC LT phase

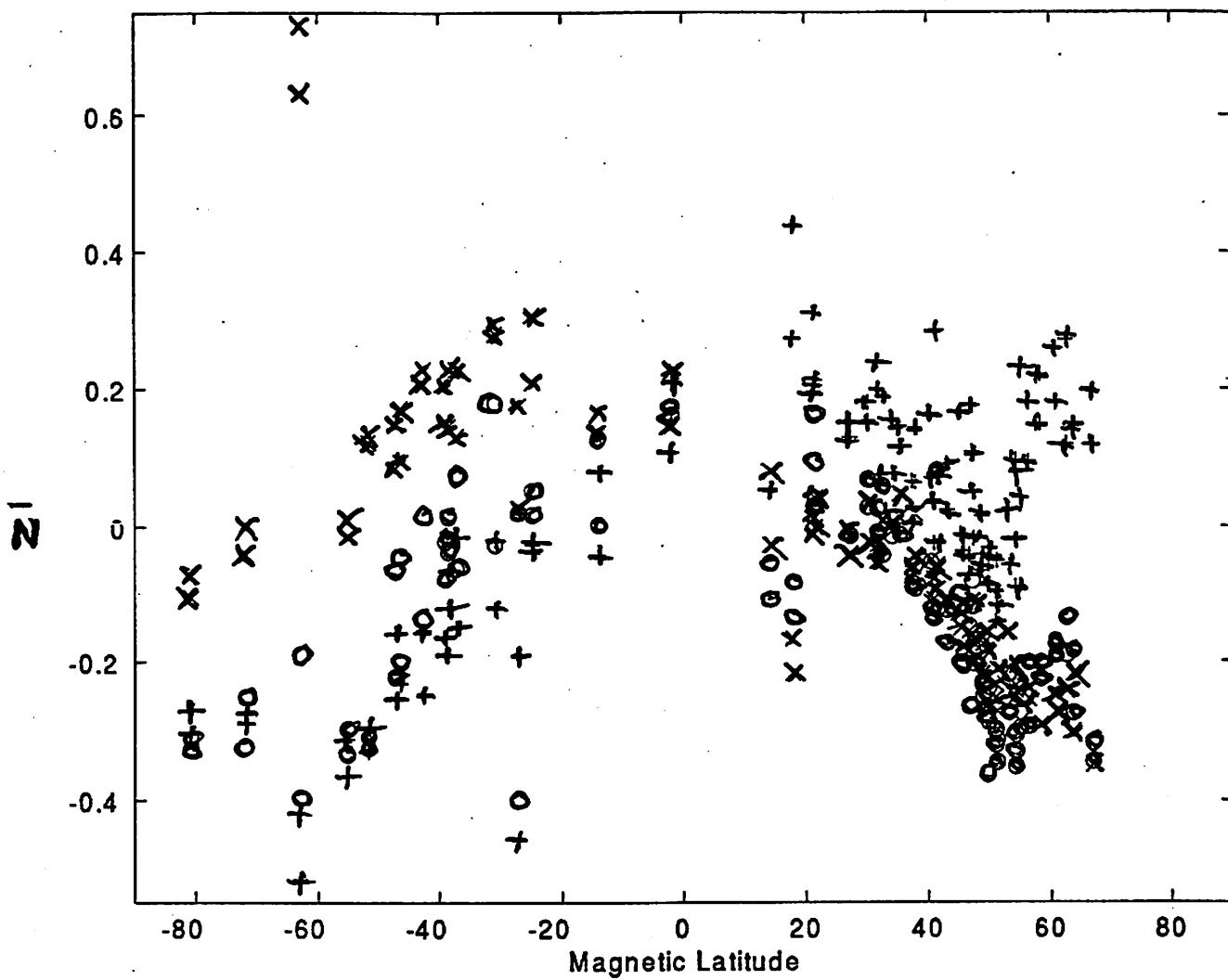


Rodger Wrenn Rishbeth JATP 1987/9
 Field & Rishbeth SASTP 1997

DC

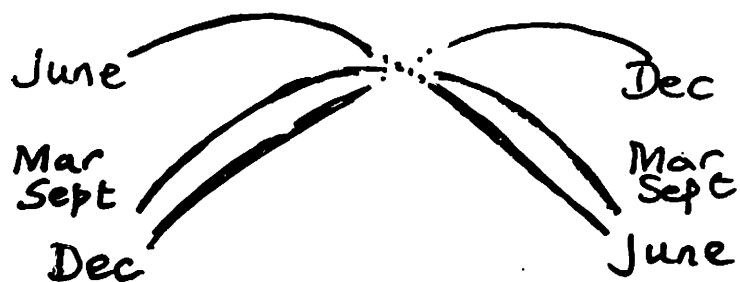
$\ln \frac{NmF2}{NnF2}^*$ · Storm
quiet

+ = Nov-Feb, x = May-Aug, o = Mar Apr Sep Oct



SUMMER
EQUINOX
WINTER

\bar{N} vs. lat :



60S 0 60N

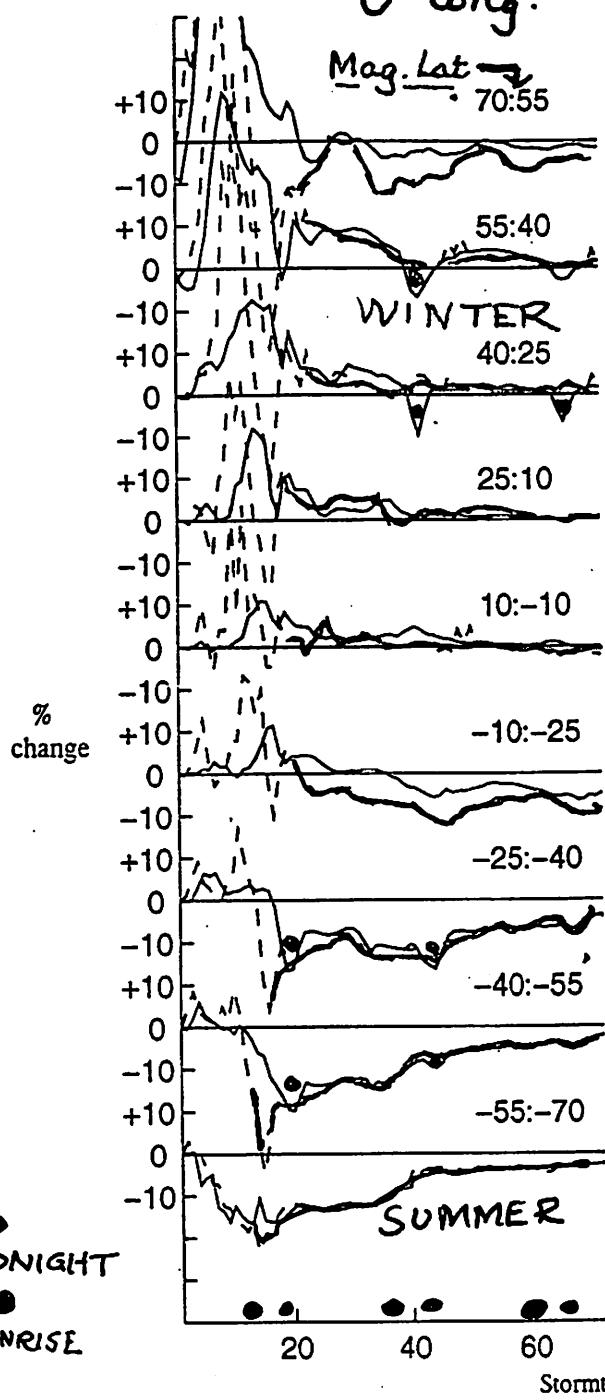
Storm - quiet
% change

$\overline{\text{NmF2}}$
 $\overline{[\text{O}/\text{N}_2]}$

December

0° long.

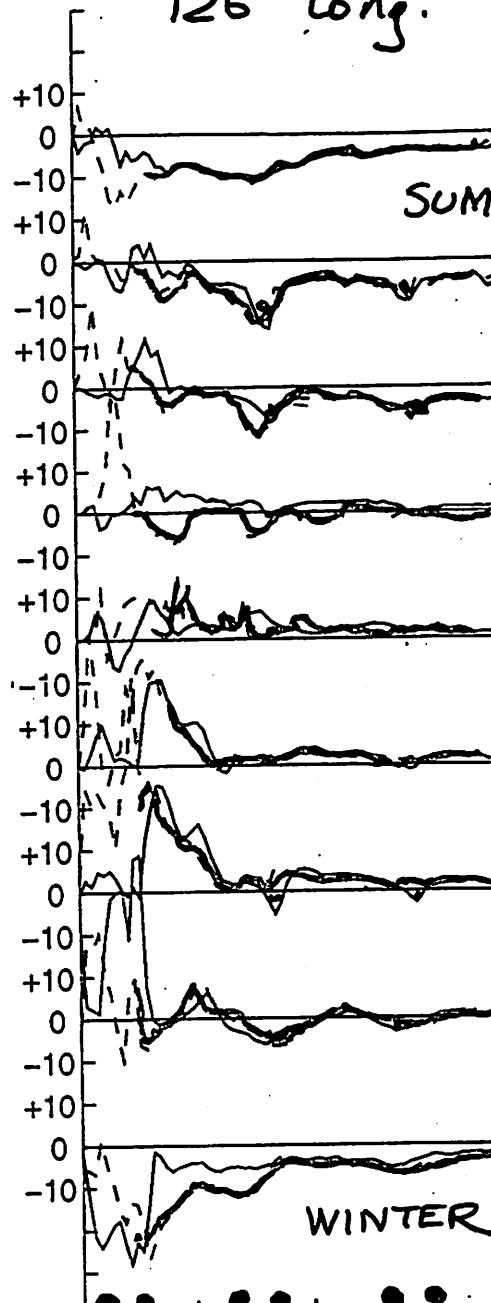
Mag. Lat. 70:55



June

126° long.

SUMMER



MIDNIGHT
SUNRISE

Stormtime ST (hours)

Fig. 1. Percentage changes in NmF2 (full curves) and $[\text{O}/\text{N}_2]$ ratio (dashed curves), computed from the CTIP coupled model for stormtime 0-72 hours, at nine zones of magnetic latitude as shown. Left: December solstice, geographic longitude 0° (European/African sector). Right: June solstice, geographic longitude 126° (Asian/Australian sector). Black dots show local midnight.

$$\left\{ \begin{array}{l} \text{EXCEPT IN} \\ \text{"INITIAL} \\ \text{PHASE"} \end{array} \right\} \frac{\text{NmF2}^d}{\text{NmF2}^q} = \frac{[\text{O}/\text{N}_2]^d}{[\text{O}/\text{N}_2]^q} \text{ AT F2 PEAK}$$

THINGS to THINK ABOUT ...

- * Always consider time & distance scales

"*How long does [a process] take?*" "*Over what distance does it operate ?*"

Vertical scales (tens of km) << Horizontal scales (100-1000 km)

Vertical speeds ($\sim \text{m s}^{-1}$) << Horizontal speeds ($\sim 100 \text{ m s}^{-1}$)

But – Vertical motions are very important (structure, energy)

- * Difficult to get *absolute* values of parameters from ionospheric observations
But – modelling is often insensitive to parameters

- * Better space/time resolution \Rightarrow new science

- * Importance of routine solar-terrestrial monitoring & WDCs

- * Use SI units !

HOW TO USE c.g.s. UNITS [gaussian, e.m.u., e.s.u. etc]

- 1 Work out numerical values, taking care with powers of 10
- 2 If it looks wrong, multiply by c
- 3 If it's still wrong, try c^2 , $1/c$, $1/c^2$, ... (etc)
- 4 If it's still wrong, try a different animal from the c.g.s. zoo
- 5 Repeat as necessary ...

HOW TO USE M.K.S.A. [S I] UNITS

- 1 Work out numerical values, taking care with powers of 10