ULTRAVIOLET SPECTROSCOPY AND REMOTE SENSING

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CEDAR WORKSHOP

Boulder CO

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OVERVIEW

- The Ultraviolet Earth: Images and Spectra
- Background and Early History
- Present Status of UV Remote Sensing
 - Global Imaging from High Altitude: DE
 - Limb Imaging from Low Altitude
- Inversion Techniques
- Global Change and Climatological Databases
- Summary

NRL FUV CAMERA ON APOLLO-16

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- Observations: 21-23 April 1972
- Location: Lunar Surface
- Sun-Earth-Moon Angle: 105 deg
- Broadband Imagery:
 - ▶ 1250 1600 Å
 - ▶ 1050 1600 Å
- Spectroscopy: 500 1600 Å





U. IOWA EXPERIMENT ON DYNAMICS EXPLORER (DE) - 1

- Launch: 3 August 1981
- Initial orbit: 570 km x 3.65 R_e
- Initial latitude of apogee: 78.2°
- Spin-scan imaging with about 12 min per image at apogee
 - Photometer A: visible
 - ► Photometer B: visible
 - ► Photometer C: Far UV

with 12 filters each

• Photometer instantaneous field-of-view 0.32°



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UVLIM

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EXCITATION MECHANISMS

• RESONANT SCATTERING OF SUNLIGHT

- ▶ SOLAR LINES: OI 1304, 1026; HI 1216, 1026; HeI 584, 540; HeII 304
- ► SOLAR CONTINUUM: NO_{γ} ; N_2^+ 1NG
- ENERGETIC PARTICLE IMPACT
 - ▶ PHOTO ELECTRONS: DAYGLOW
 - ► PRECIPITATING ELECTRONS (AND PROTONS): AURORA
 - ► HEAVY ATOM PRECIPITATION: NIGHTGLOW
- PHOTO/ION CHEMISTRY
 - ▶ PHOTO-IONIZATION: $O + h\nu \rightarrow O^{*+} + e \rightarrow 834 \text{ Å}$
 - ► RADIATIVE RECOMBINATION: $O^+ + e \rightarrow O$ (1356, 1304, 989, 911)
 - ► ION-ION RECOMBINATION: $O^+ + O^- \rightarrow O^* + O^*$
 - ► CHEMILUMINESCENCE: $N(^{4}S) + O(^{3}P) \Leftrightarrow NO(C)$

 $O + O + M \rightarrow O_2 (A)$

► DISSOCIATION: $O_2 + [h\nu \text{ or } e] \rightarrow O + O^*$

 $N_2 + [h\nu \text{ or } e] \rightarrow N + N^* \text{ or } N + N^{+*} + e$

EARLY HISTORY OF UV SPECTROSCOPY

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1946 1949 1955	Sun observed to 2100 Å, O_3 observed to 55 km	Baum et al. [1946]
	O_2 extinction observed H Lyman α discovered	Friedman et al. [1951] Byram et al. [1961]

Other early contributors: T.M. Donahue, R. Tousey, J.W. Chamberlain, J. Brandt, F.S. Johnson, I.S. Shklovsky, C. Barth, V.G.Kurt, S.A. Kaplan

VERY LONG MATURATION TIME FOR ROUTINE UV REMOTE SENSING MISSIONS

because:

- Reliable cross sections, oscillator strengths... required
- High sensitivity imagers and spectrographs needed
- Science models had to be developed

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• Retrieval algorithms necessary for parameter extraction

GLOBAL IMAGING MODEL

• AIRGLOW MODEL

- ► MSIS-86 SPECIFIES GLOBAL CONCENTRATIONS
- ► HINTEREGGER SOLAR EUV FLUX ALGORITHM
- ► STRICKLAND AND MEIER PHOTOELECTRON MODEL
- ► RADIATIVE TRANSPORT MODELS
 - OI 1356: Strickland and Anderson
 - OI 1304: Gladstone or Meier
- IMAGING MODEL
 - MODEL DESCRIBED BY: Strickland, Cox, Barnes, Paxton, Meier, Thonnard -accepted in Applied Optics [1993]
 - USES GEOGRAPHIC GRID WITH VARIABLE SPACING IN SOLAR ZENITH AND AZIMUTH ANGLES
 - PRE-COMPUTES INTENSITIES VS LOOK ANGLE AT EACH GRID POINT AND INTERPOLATES TO DYNAMICS EXPLORER LOOK ANGLES



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CONCLUSIONS: 1230 - 1800 Å DATA

- MODEL REPRODUCES DATA TO WITHIN \pm 17%
 - ▶ model scaling for individual images: 0.92 0.99
 - deviations due to statistics or aurora
- SPATIAL VARIABILITY WELL-REPRODUCED BY MODEL
- OI 1304 Å DOMINATES (~ 85 %) DISK PORTION OF (FILTER 2) IMAGES
 - ▶ 60 % (photoelec.) + 25 % (solar res.scat.) = 85%
 - ▶ OI 1356 Å ~ 7 %
 - ▶ N_2 LBH bands ~ 7 %
- OI 1356 Å AND N₂ LBH BANDS DOMINATE AT LIMB
- CAN PROCEED WITH CONFIDENCE TO DISTURBED-TIME IMAGES



- EVALUATE CONCEPT OF GLOBAL UV REMOTE IONOSPHERIC SENSING
- OBTAIN ELECTRON DISTRIBUTION FROM AIRGLOW MEASUREMENTS
- MEASURE THE ALTITUDE DISTRIBUTION OF:
 O⁺, O, N₂, O₂, N, NO, Na (75 750 km; 6 km RESOLUTION)
- FLY STATE OF THE ART IMAGING SPECTROSCOPIC INSTRUMENTATION (500 8700 \)



O, O₂ AND N₂ CONCENTRATIONS FROM OI 1356 AND LBH

• OI 1356 Å

$$4\pi I_{1356} = \int n_O g_{1356} e^{-\tau_{O_2}} ds$$

• N₂ LBH (2,0) 1383 Å

$$4\pi I_{1383} = \int n_{N_2} g_{1383} e^{-\tau_{O_2}} ds$$

• photoelectron g-factors are (non-linearly) proportional to total column density to sun, N through the photoelectron flux, Φ :

$$g[z] = \int \sigma(E) \Phi(E, N[z]) dE$$

RAIDS ALGORITHM APPROACH:

DISCRETE INVERSE THEORY

- Fast, rigorous, systematic inversion technique for obtaining environmental data records (EDR's) from UV limb scans
- Applicable to all RAIDS science data records (SDR's)
- Straightforward approach for obtaining uncertainties in EDR's
- Uncertainties from statistics in retrieved neutral densities of about 3-10% between 150 and 350 km, and 2% for exospheric temperature
- Validation will ultimately be with observational data and ground truth

NONLINEAR DISCRETE INVERSE THEORY

• Model Equation:

$$G(m) = I^{pre}$$

 P^{re} = vector of length N, representing predictions of quantities which have been observed

G = vector function representing model

m = vector of length M, corresponding to the model parameters

• Simple Iterative Approach; Use Taylor's Theorem:

$$\nabla G[m_{i+1}^e - m_i^e] = I^{obs} - G(m_i^e)$$

• Taking inverse (overdetermined case):

$$\boldsymbol{m}_{i+1}^{e} = \boldsymbol{m}_{i}^{e} + \nabla \boldsymbol{G}_{inv} [\boldsymbol{I}^{obs} - \boldsymbol{G}(\boldsymbol{m}_{i}^{e})]$$

• Generalized inverse:

$$\nabla G_{inv} = \{ \nabla G^T [cov I^{obs}]^{-1} \nabla G \}^{-1} \nabla G^T [cov I^{obs}]^{-1}$$

• Covariance of Model Parameters:

$$[cov m] = \nabla G_{inv}[cov I^{obs}] \nabla G_{inv}^{T}$$

Equivalent to Maximum Likelihood Method

DISCRETE INVERSE THEORY ALGORITHM FOR OI 1356 Å AND N₂ LBH

EXAMPLES: Limb scans from 65 to 325 km tangent altitude

- 65 measurement for each line
- Synthetic data with Gaussian noise (1 ct/s/R)
- Solar min, max; 0, 70 deg solar zenith angle, 850 km satellite

CASE 1: *m* represents 5 scaling parameters describing the MSIS model:

$$m = [f_{N2}, f_{O}, f_{O2}, f_{10}, f_{euv}]$$

CASE 2: *m* represents individual densities:

$$\boldsymbol{m} = [n_{N_2}(110), \dots, n_{N_2}(500), n_0(110), \dots, n_0(500), \\ n_{O_2}(110), \dots, n_{O_2}(500)]$$





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Roble and Dickinson: Trace Gases and the Upper Atmosphere

NRL UPPER ATMOSPHERE/IONOSPHERE MISSIONS



SUMMARY

- UV remote sensing has evolved sufficiently that accurate, routine observations of the upper atmosphere (and ionosphere) are possible on a global scale
- Non-linear inversion methods show great promise for retrieval of concentrations from airglow observations
- Validation is underway with various data sets
- DoD missions should provide continuous observations over 1-2 solar cycles, beginning next year with RAIDS