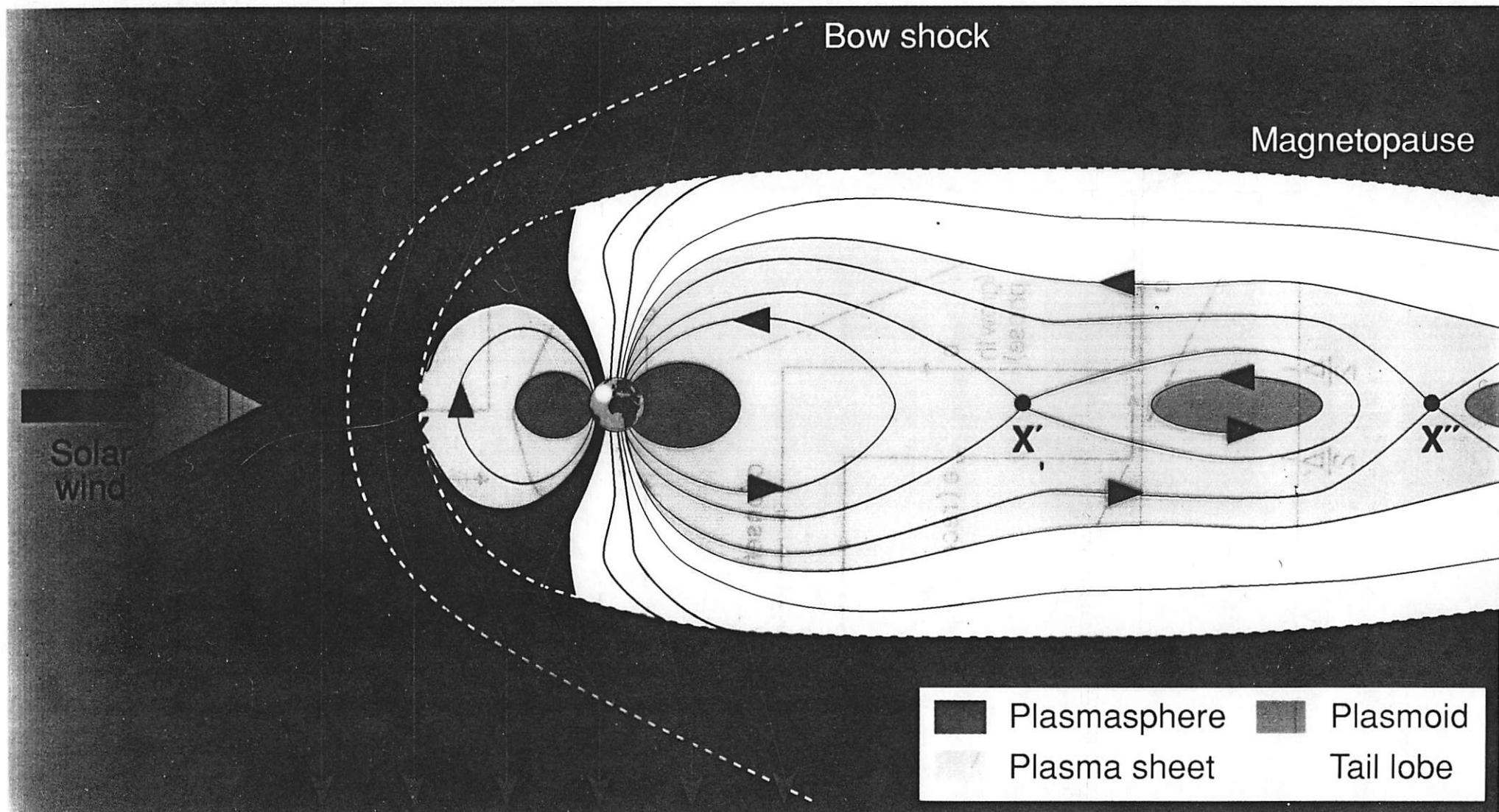


1992 CEDAR Workshop
Boulder, CO
June 21-26, 1992

Tutorial Lecture

by Michael Lockwood
Rutherford-Appleton Laboratory

Time-Varying Convection



Magnetohydrodynamics (MHD)

From 3 of Maxwell's equations:

$$\nabla \cdot \bar{B} = 0$$

$$\bar{J} = (\nabla \times \bar{B}) / \mu_0$$

$$\partial \bar{B} / \partial t = -(\nabla \times \bar{E})$$

(Ampère's Law with $\partial D / \partial t = 0$)
 (Faraday's Law)

Plus a generalised Ohm's Law

$$\bar{J} = \sigma (\bar{E}') = \sigma (\bar{E} + \bar{v} \times \bar{B})$$

We get the "INDUCTION EQUATION"

$$\partial \bar{B} / \partial t = \nabla \times (\bar{v} \times \bar{B}) + \nabla^2 \bar{B} / \mu_0 \sigma$$

convective term diffusive term

Magnetic Reynolds Number,

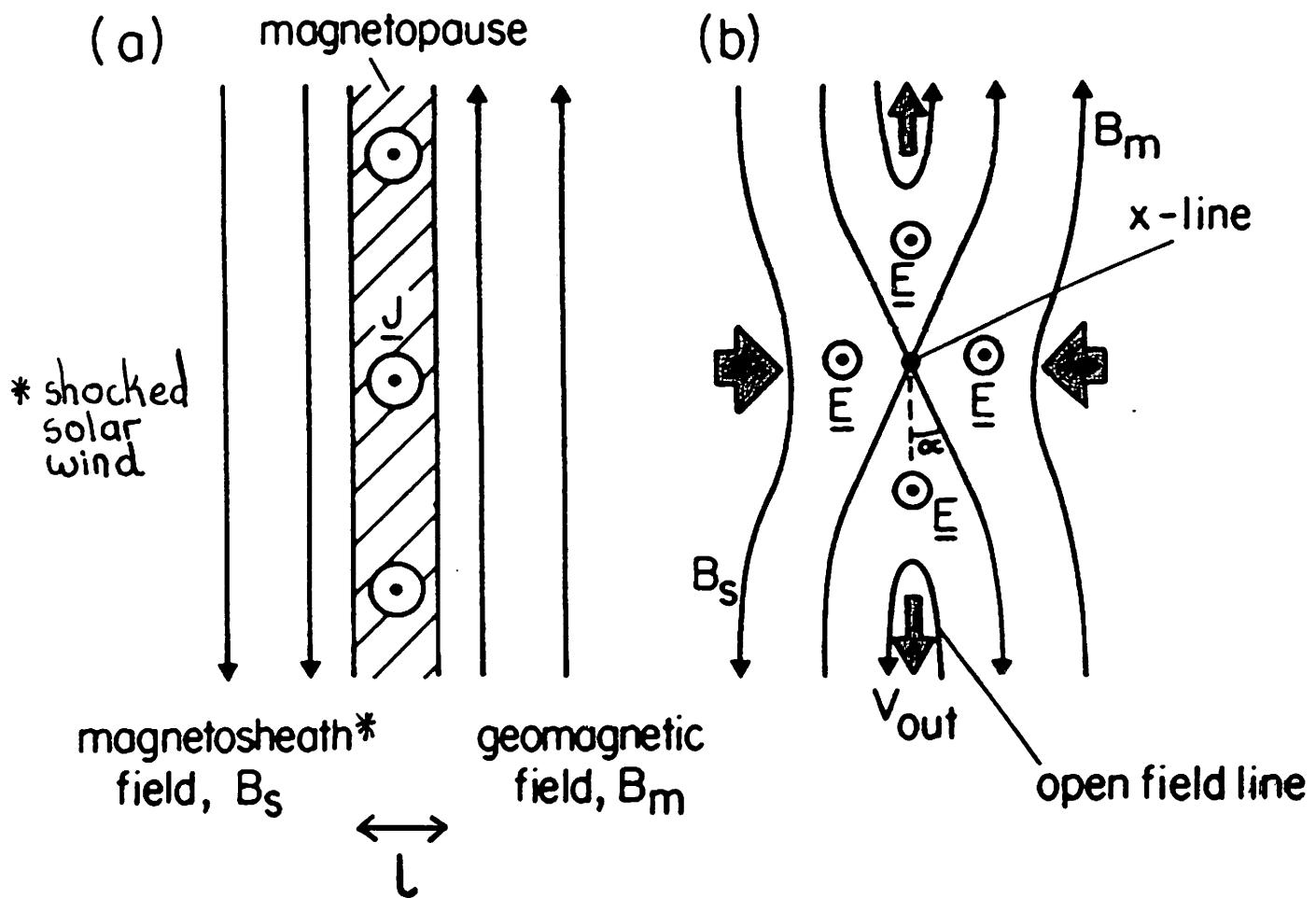
$$R_m = \frac{\text{convective term}}{\text{diffusive term}} \sim \frac{VB/L}{B/\mu_0 \sigma L^2} = \mu_0 \sigma V L$$

- If conductivity σ , and spatial scale, L , very large, R_m v. large
 $\partial \bar{B} / \partial t = \nabla \times (\bar{v} \times \bar{B})$: in fact means \bar{B} moves with \bar{v} - we say \bar{B} is FROZEN-IN to plasma flow, \bar{v} (applies in most of interplanetary space, magnetosphere + F-region ionosphere)
- If L small, so $R_m \ll 1$, $\partial \bar{B} / \partial t = \nabla^2 \bar{B} / \mu_0 \sigma$ and field DIFFUSES through plasma. Actually, high σ ensures that this only occurs in highly localised (small L) regions

MAGNETIC RECONNECTION

(Dungey, 1953; 1961)

- a breakdown of ideal MHD and 'frozen-in flux'



For ideal MHD, B_s and B_m cannot mix (both frozen-in)

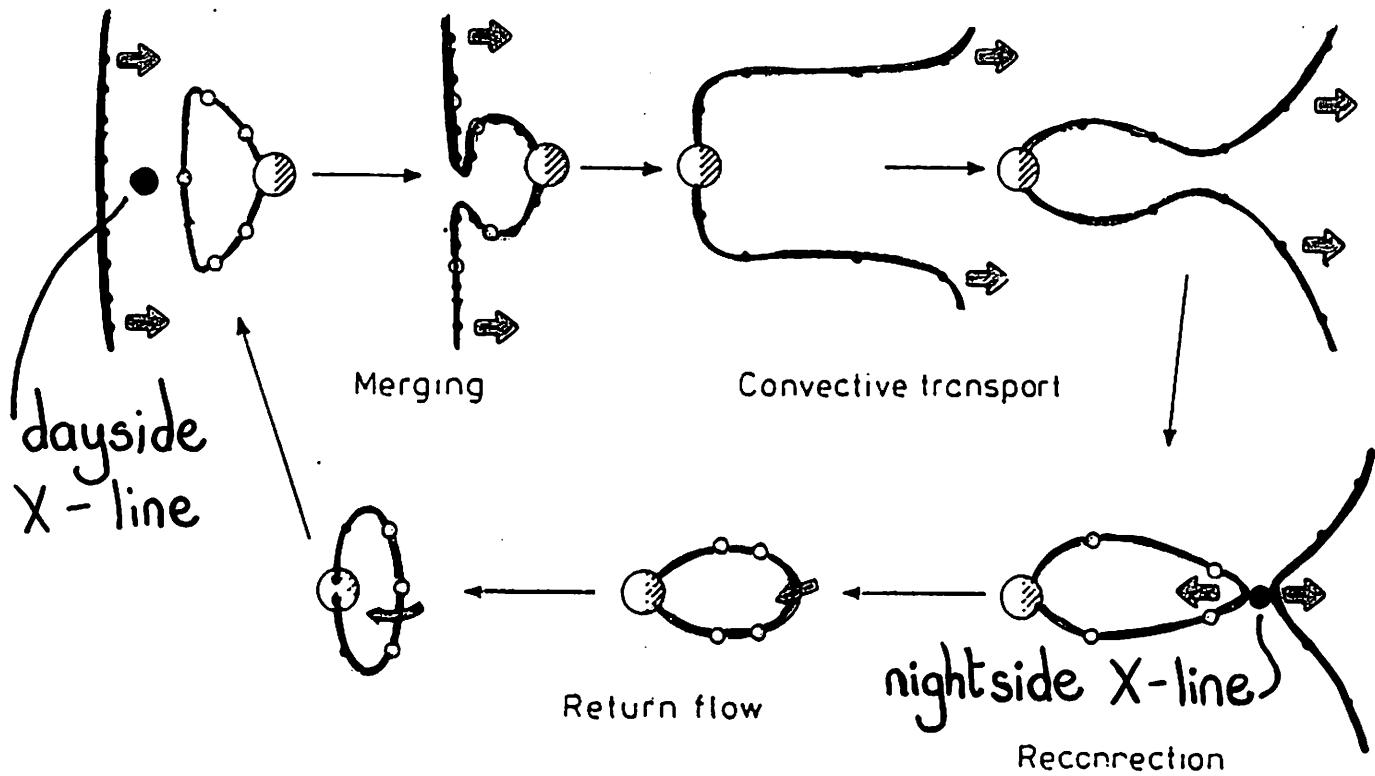
If $\bar{B}_s \neq \bar{B}_m$ current must flow in magnetopause

Solar wind compresses current layer so L small
and $R_m \rightarrow 1$. Hence diffusion of B through plasma becomes important.

Motion of \bar{B}_s and \bar{B}_m into X -line and of open field lines away from X -line, corresponds to an electric field, \bar{E} , along X -line, called the reconnection rate

Field Line Convection due to Magnetic Reconnection

cycle time $\gtrsim 6$ hours



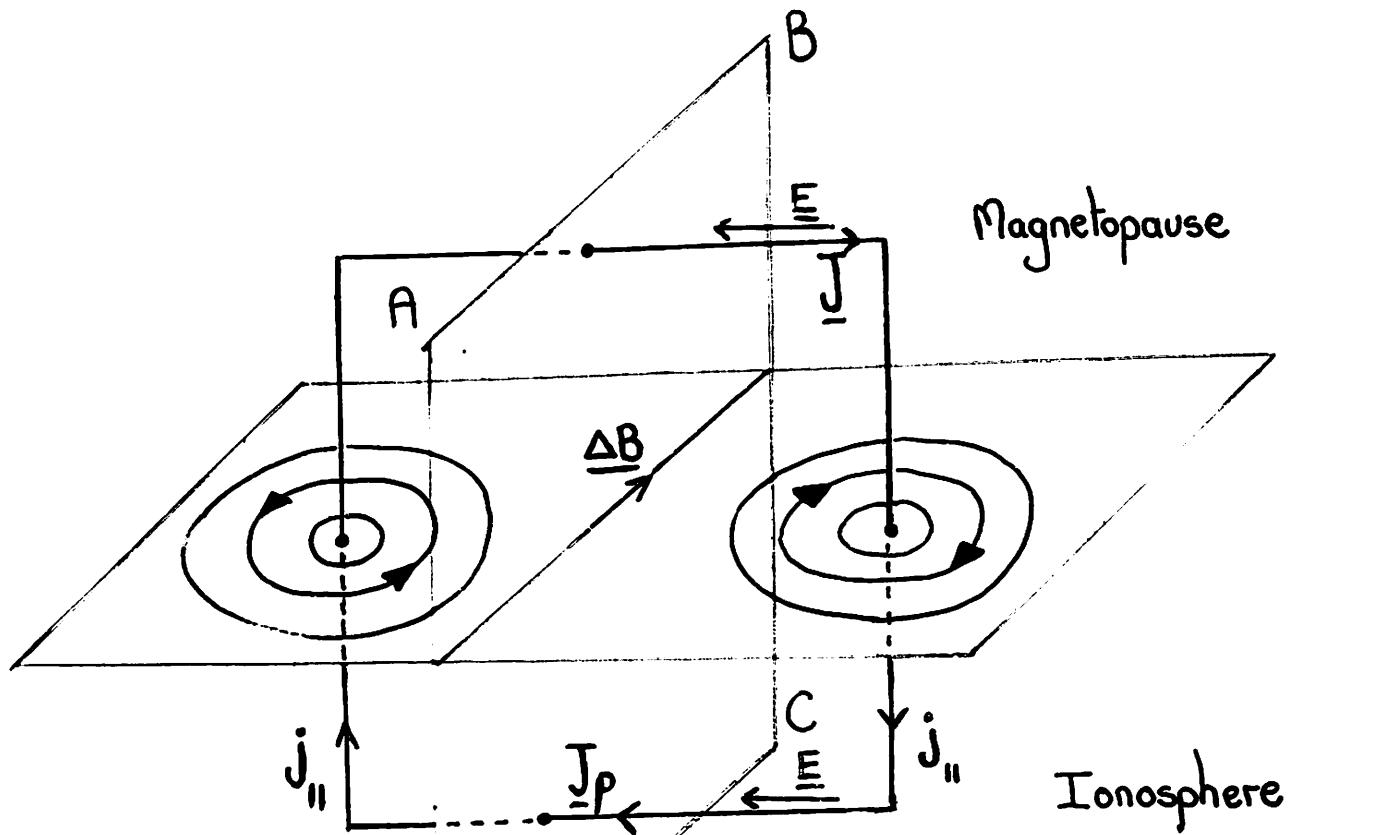
- Solar wind particle
- Plasma sheet

Figure 8.27 The history of a field line. (After S.-I. Akasofu, *Chapman Memorial Lecture*, 1973.)

- Interplanetary Field Line ($B_z < 0$)
- Closed Geomagnetic Field Line
- Open Field Line

Pedersen currents very important because:

- They close magnetospheric current systems
 - Allow momentum transfer from the solar wind into the ionosphere
 - Deposit solar wind energy in the upper atmosphere



At magnetopause, $\underline{J} \cdot \underline{E} < 0$
(energy extracted from plasma)

At mid-altitudes, downward
Poynting Flux, S_z —

In ionosphere, $\underline{J} \cdot \underline{E} > 0$
(energy deposited in plasma)

A-2 MODEL.

(IMF $B_y < 0$
For Northern
hemisphere)

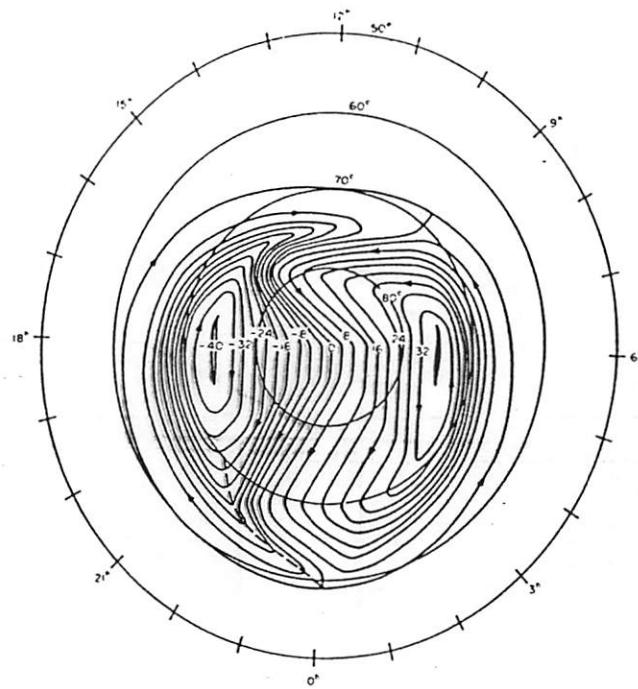
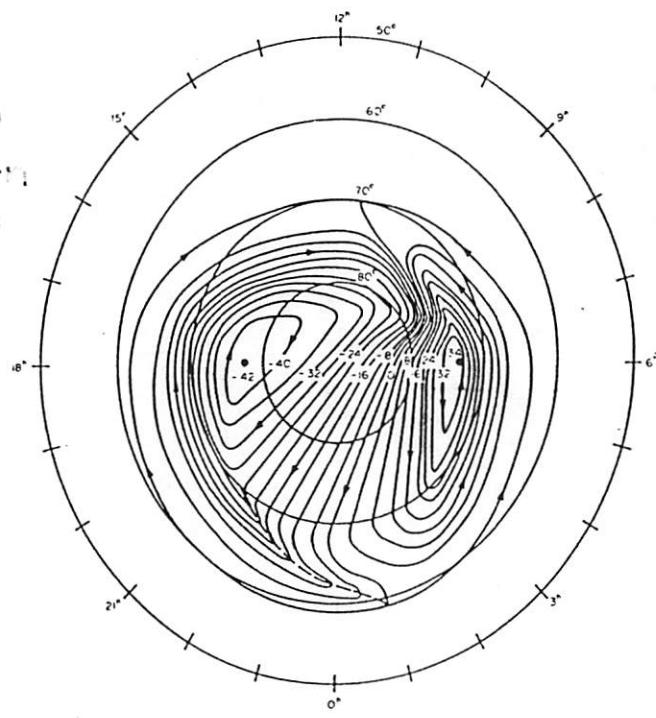


FIG. 7(a)

B-2 MODEL.

(IMF $B_y < 0$
For southern
hemisphere)



(b)

FIG. 7. POLAR CONVECTION ELECTRIC POTENTIAL DISTRIBUTIONS FROM HEPPNER (1983), WITH DEPENDENCE ON THE "Y" COMPONENT OF THE IMF.

- (a) The A-2 field—applied to the Northern Hemisphere with $B_y - ve$. Applied to the Southern Hemisphere with $B_y + ve$.
- (b) The B-2 field—applied to the Southern Hemisphere with $B_y - ve$. Applied to the Northern Hemisphere with $B_y + ve$.

HEPPNER MODEL OF CONVECTION ELECTRIC POTENTIAL

Origins of magnetospheric and ionospheric convection

• 1. Reconnection

Transfer of open flux from dayside to nightside
by solar wind

Return flow of closed field lines

• 2. Viscous-like interactions

any process which acts on closed field lines

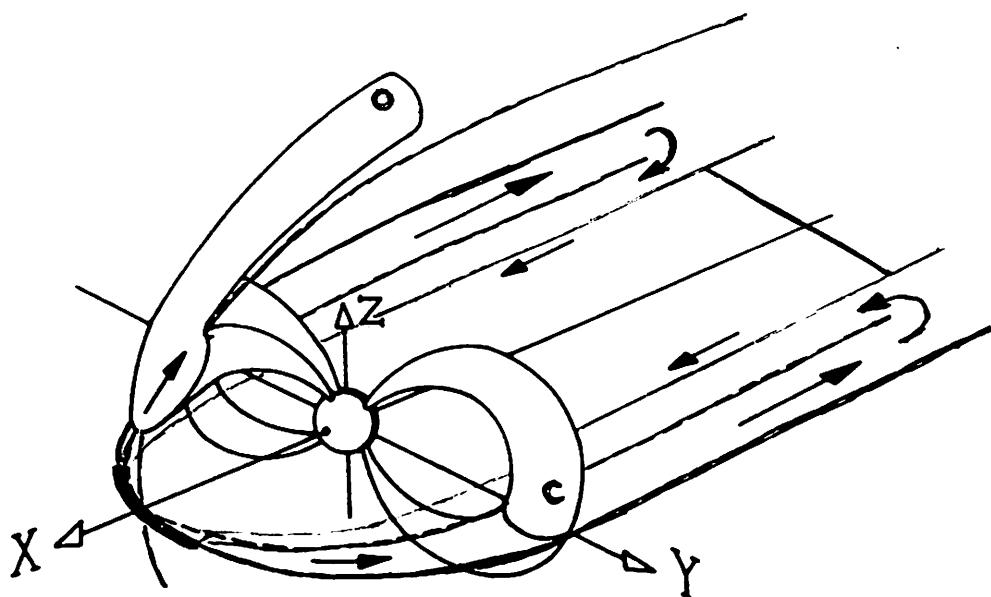
"anything but reconnection"

includes:

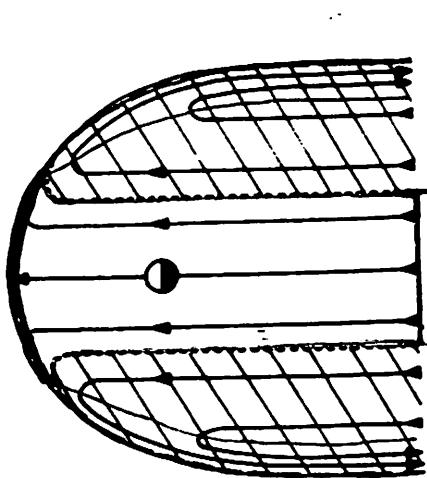
"impulsive penetration"; wave-driven diffusion;
gradient drift entry; Kelvin-Helmholtz waves;
solar wind dynamic pressure changes; etc.

BUT how much of the apparently viscous effect is due to
continuing tail reconnection of residual lobe open flux

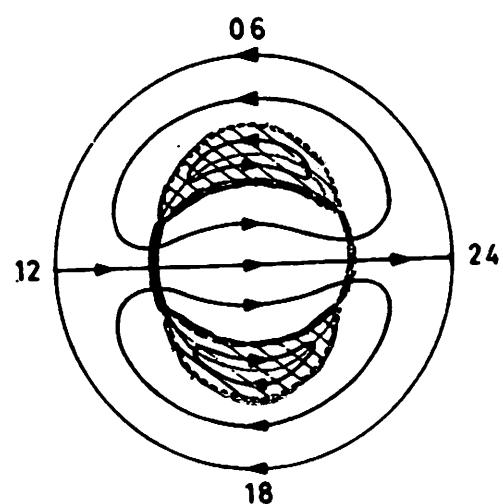
Cowley, 1984



(a)



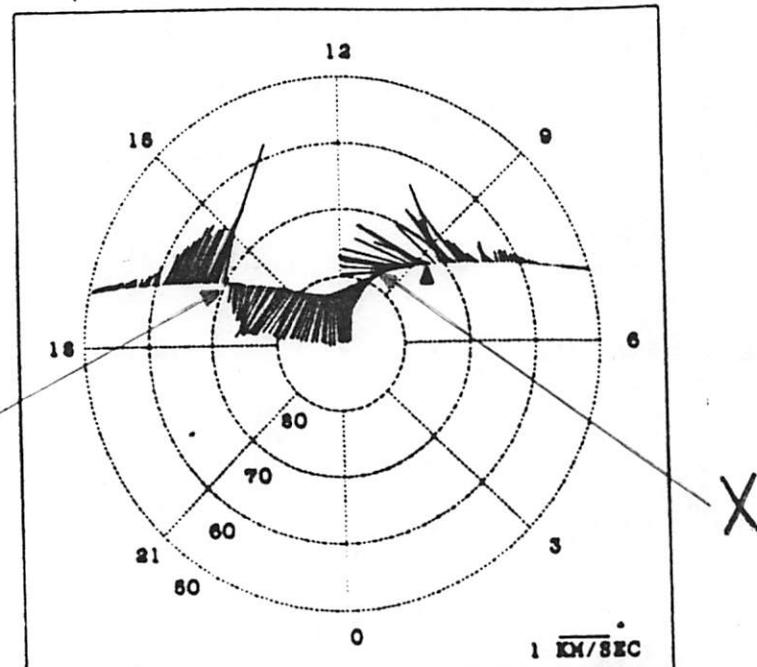
(b)



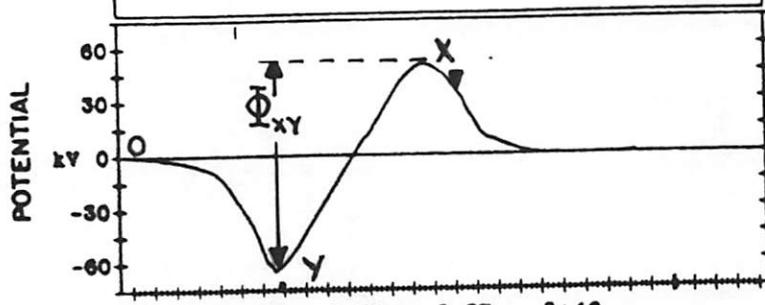
(c)

Satellite Flow Data (Coley et al., 1987)

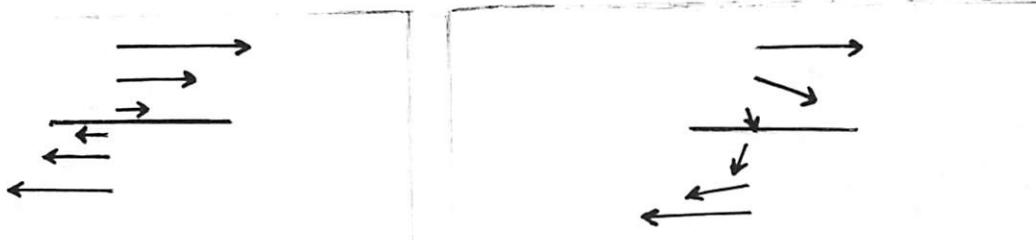
DE-8 ION DRIFT VELOCITIES
 MLT V ILAT NORTHERN HEMISPHERE
 DAY 62181 UT 8:31 ORBIT 4498



$$\Phi_x = \int_{-\infty}^x E_l dl$$



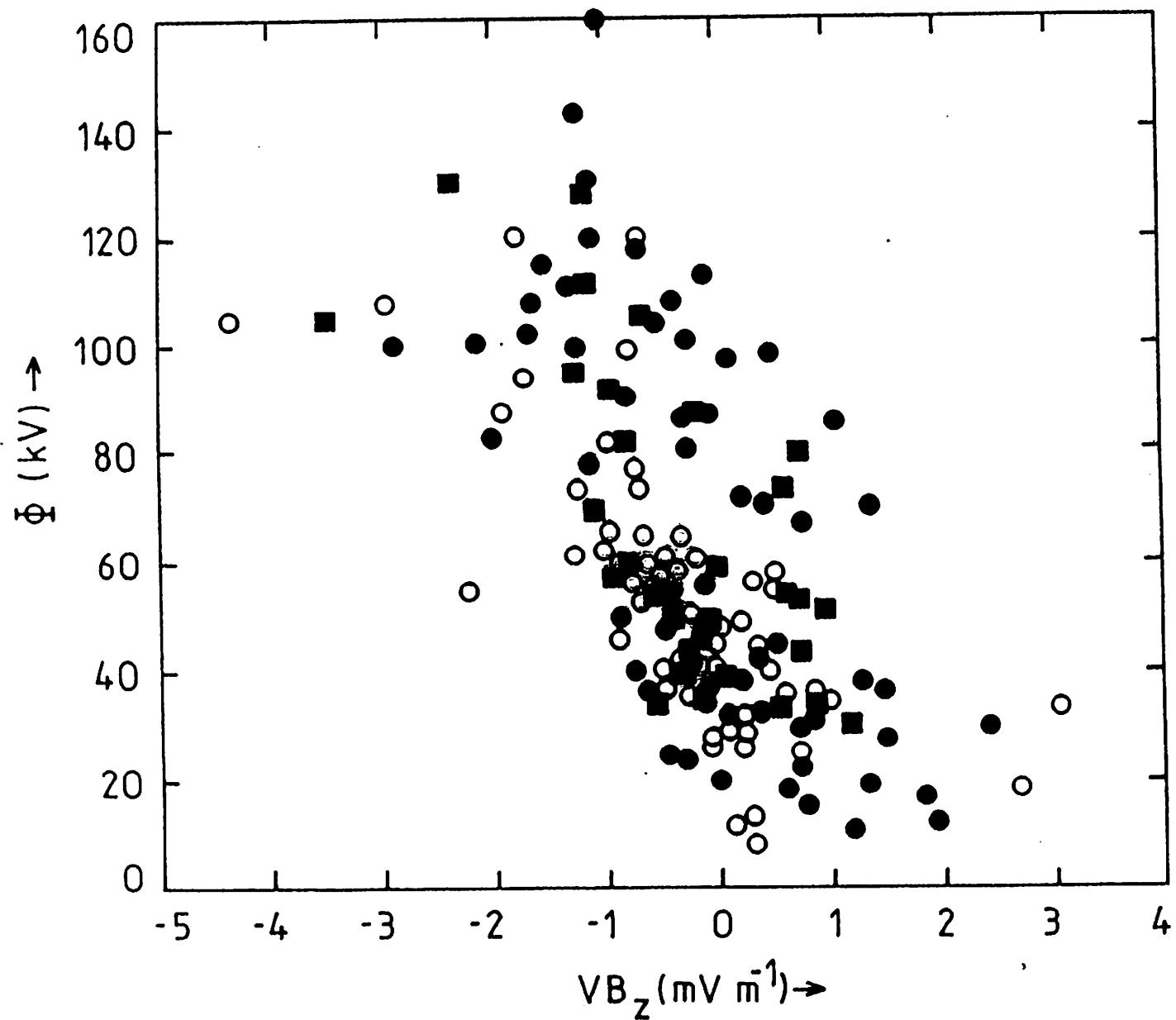
UT(HR:MIN)	8:27	8:32	8:37	8:42
ALT(KM)	692	622	447	378
GLAT(DEG)	69.46	67.14	73.77	64.24
GLNG(DEG)	-170.2	-170.8	7.1	8.9
MLT(HRS)	16.49	13.80	8.66	7.17
ILAT(DEG)	66.86	81.48	72.04	63.40



Shear reversal at Y Rotational reversal at X

If X is near dawn and Y near dusk, Φ_{xy} is called the transmeridional voltage. Φ_{xy}

■ AE-C & D ○ S 3-2 ● S 3-3



Reiff et al. (1981)
Doyle and Burke (1981)
Wygant et al. (1983)
Cowley (1984)

PLASMA FLOWS IN THE EQUATORIAL MAGNETOSPHERE

ELECTRIC EQUIPOTENTIALS OR FLOW LINES

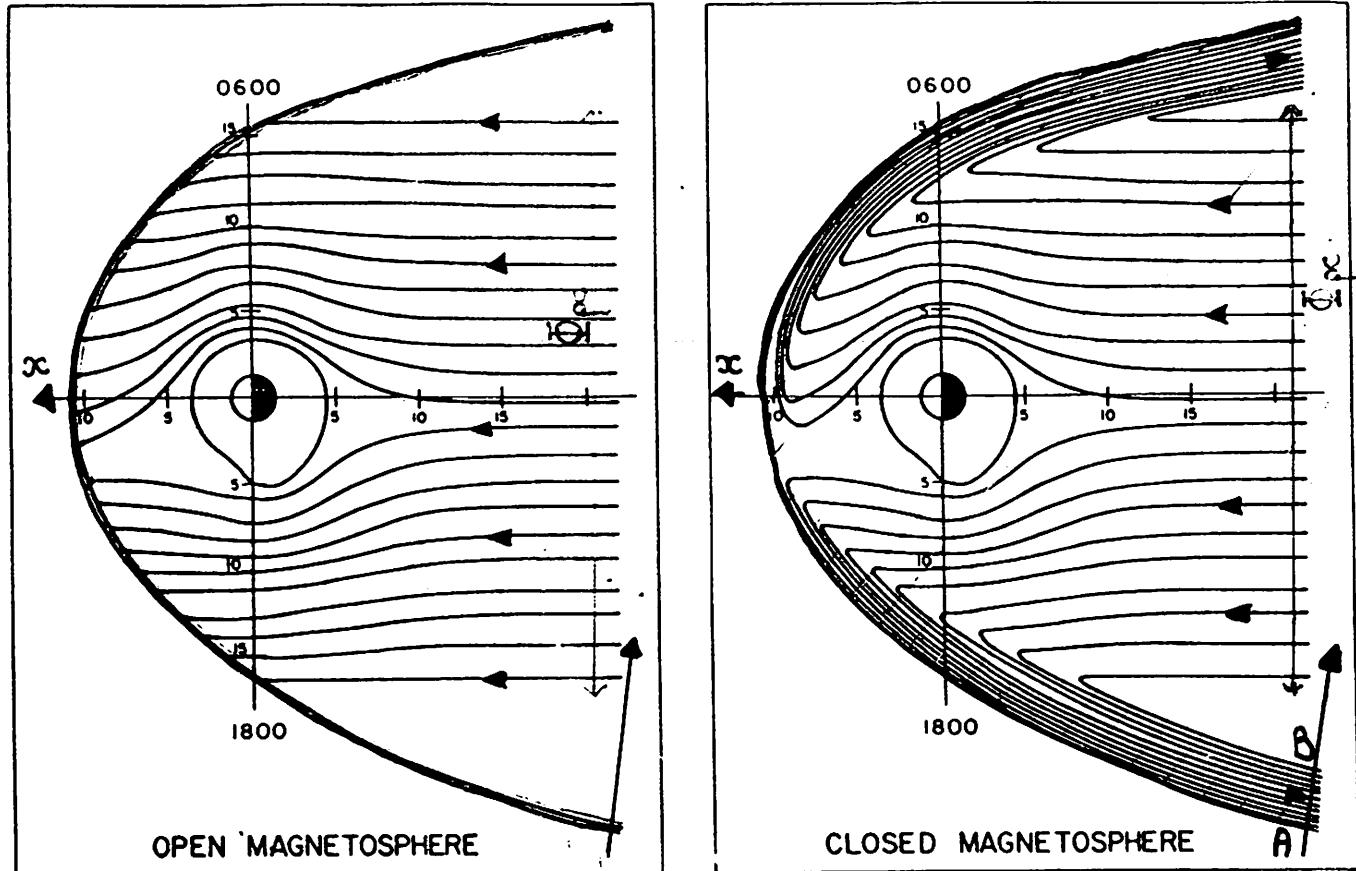


Fig. 1. Electric equipotentials (or $E \times B$ flow lines) in the equatorial plane of idealized completely open and completely closed magnetospheres.

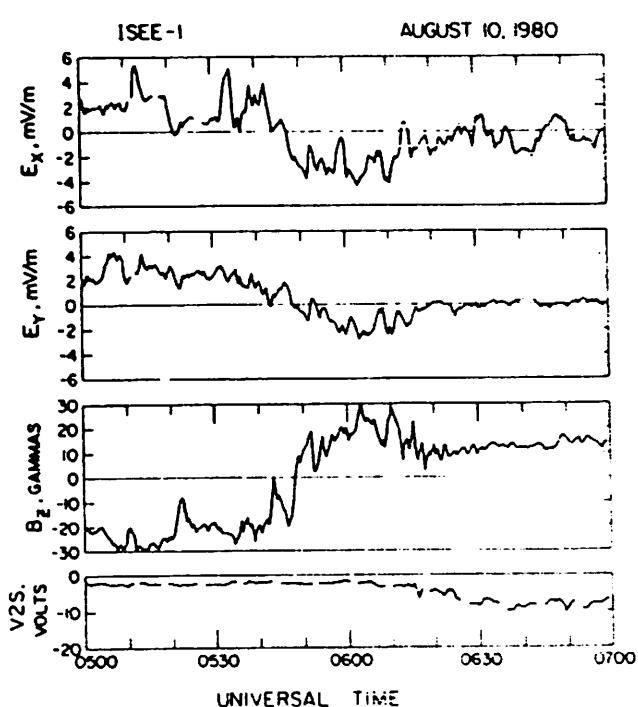


Fig. 2. Electric and magnetic fields measured before and after a dusk magnetopause crossing.

Consider electric field observations by a satellite moving into magnetosphere at low magnetic latitudes, as shown

$$\text{Plasma drift, } \bar{v} = \bar{E} \times \bar{B} / B^2$$

\therefore drift normal to path \equiv electric field component along path, E_0

$$\text{Voltage, } \Phi_v = \int_A^B E_0 dl$$

where A and B are the edges of the anti-sunward moving LLBL

For fully-open magnetosphere

$$\Phi_v = 0$$

Mozer, Geophys. Res. Lett., 11, 135, 1984

[NB. In comment, Heikkila (Geophys. Res. Lett., 9, 877, 1986) claims LBL was not crossed far enough down tail to see full Φ_v].

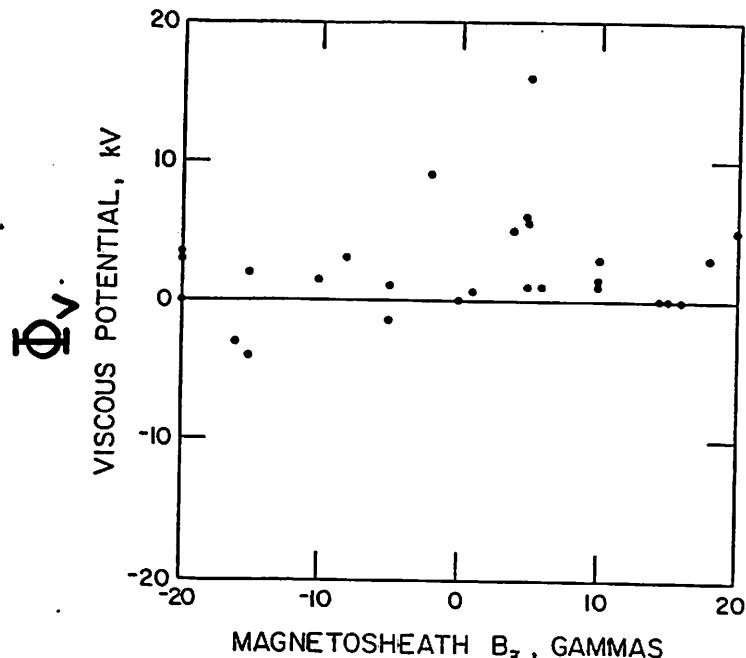


Fig. 4. Scatter plot of the viscous interaction potential versus B_z for 28 magnetopause crossings within 2 hours of local dusk.

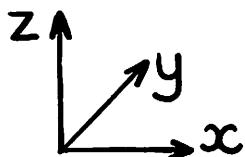
Φ_v is less than 15 kV in all cases
and typically < 5 kV.

If Φ_v is the same on both flanks

total contribution to Φ_{pc} of viscous-like interactions
is typically 10 kV and less than 30 kV in all cases
(consistent with Φ_{pc} observed during northward IMF)

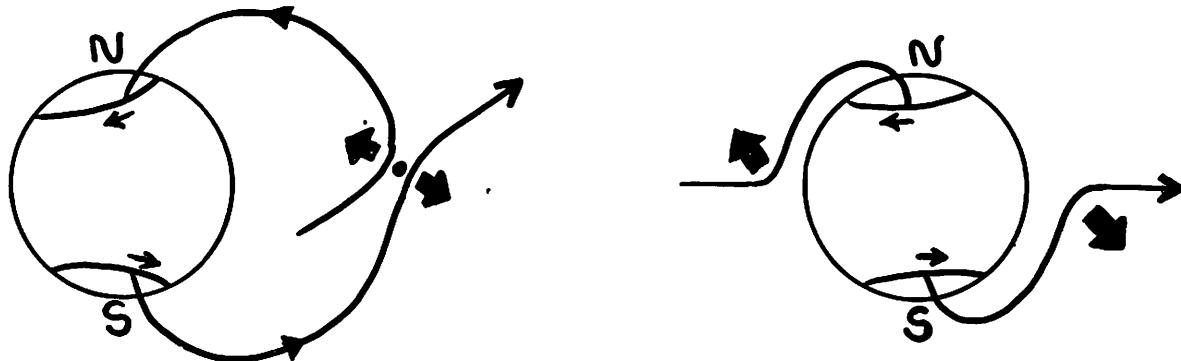
We Conclude: Reconnection contributes up to ~100 kV to Φ_{pc}
Viscous-like interaction contribute $\lesssim 20$ kV

Svalgaard - Mansurov Effect



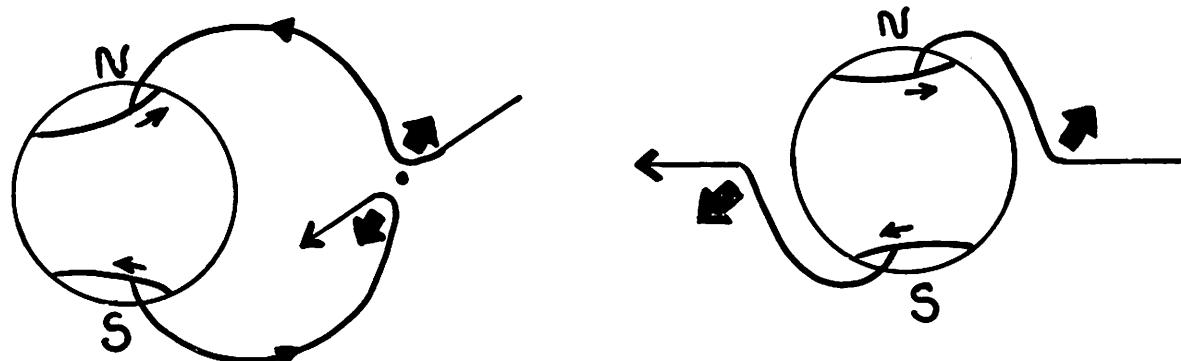
as explained by
Jørgensen et al. (1972)
Atkinson (1972)
Cowley (1981)

IMF $B_y > 0$



Westward } flow in { Northern } hemisphere
Eastward }
 } Southern

IMF $B_y < 0$



Eastward } flow in { Northern } hemisphere
Westward }
 } Southern

$A_1 B_1 > 0$

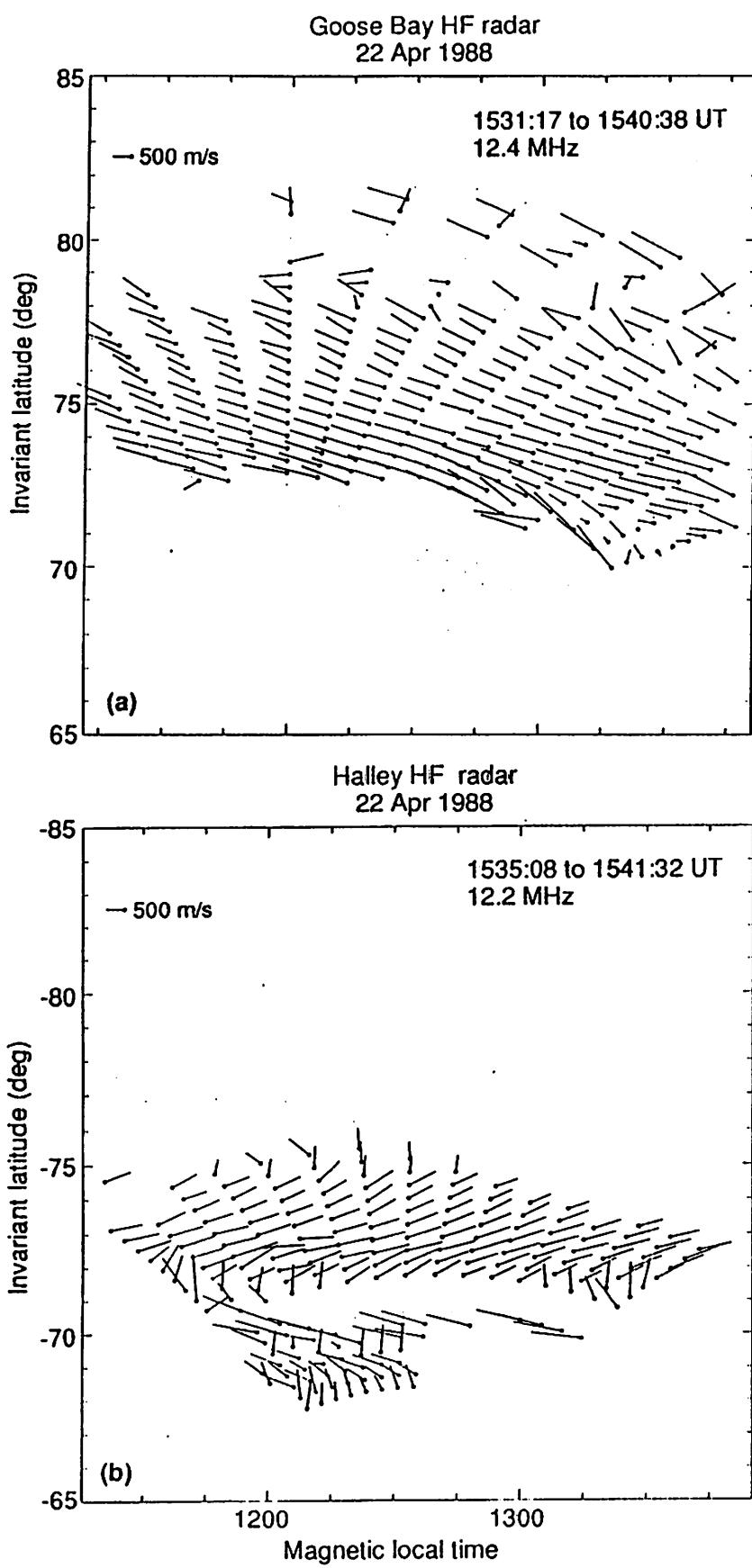


Fig. 6

$B_x B_y < 0$

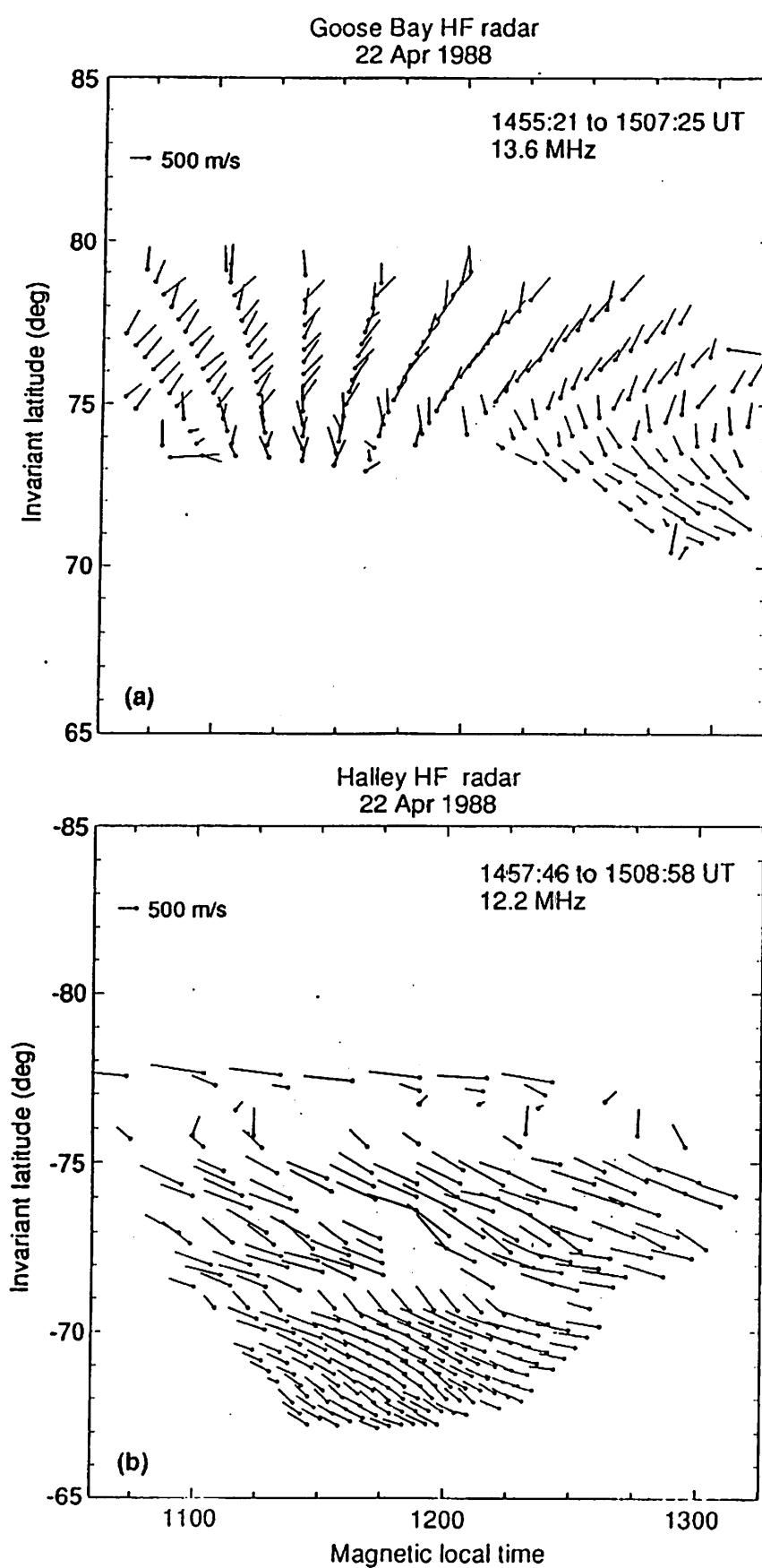


Fig. 7

event C
N. hemisphere
 $+ \rightarrow -$

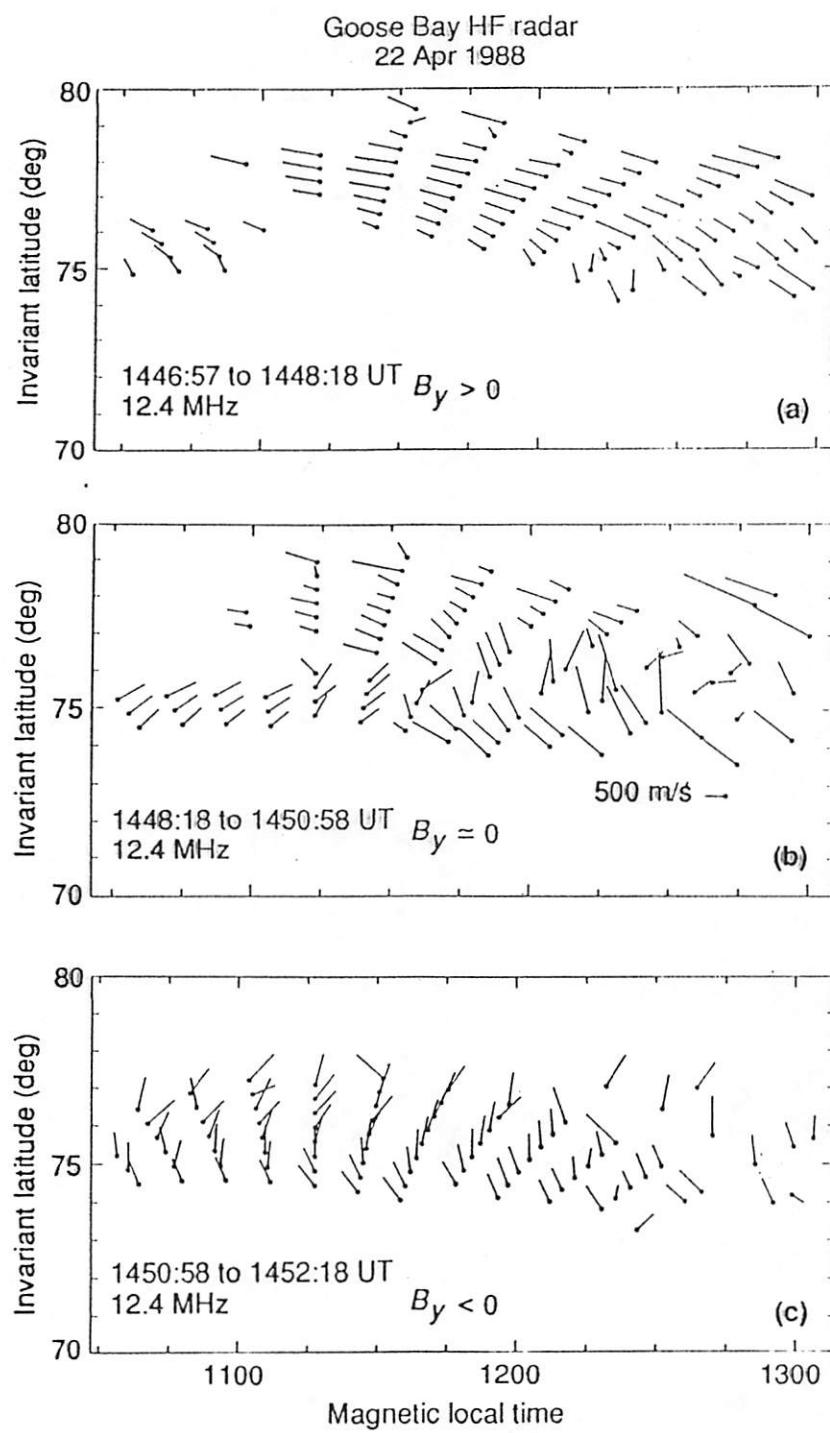
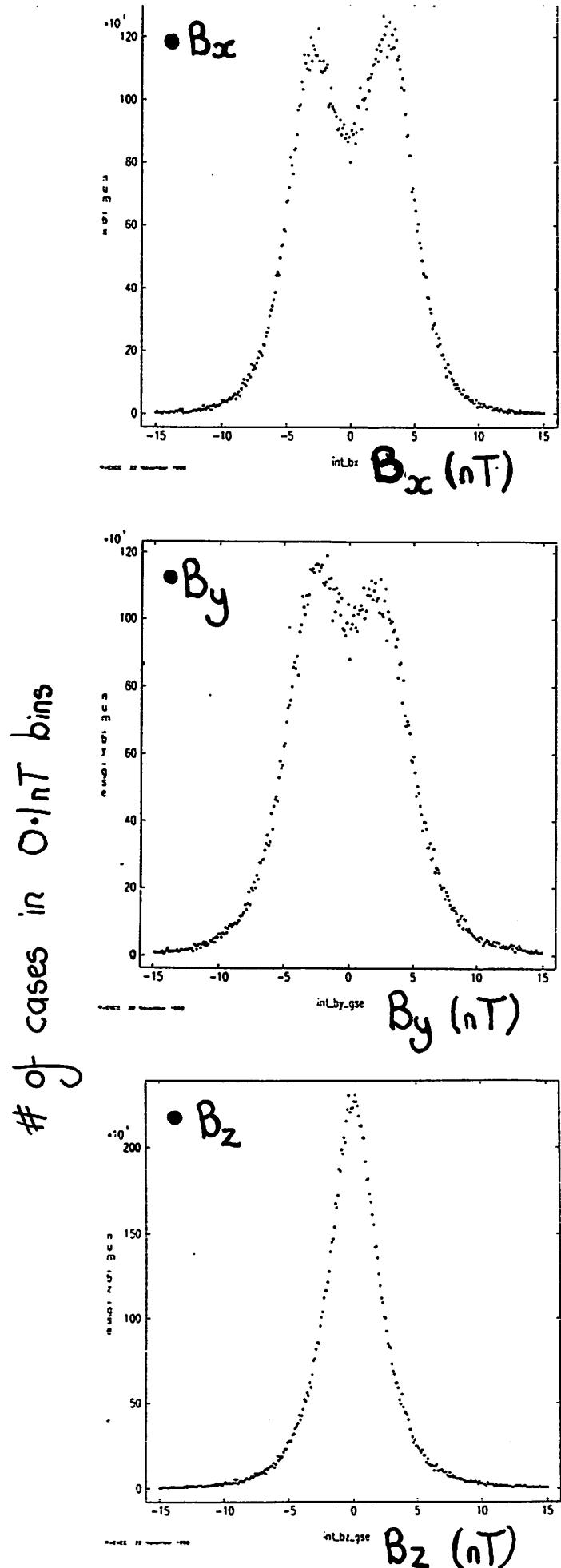
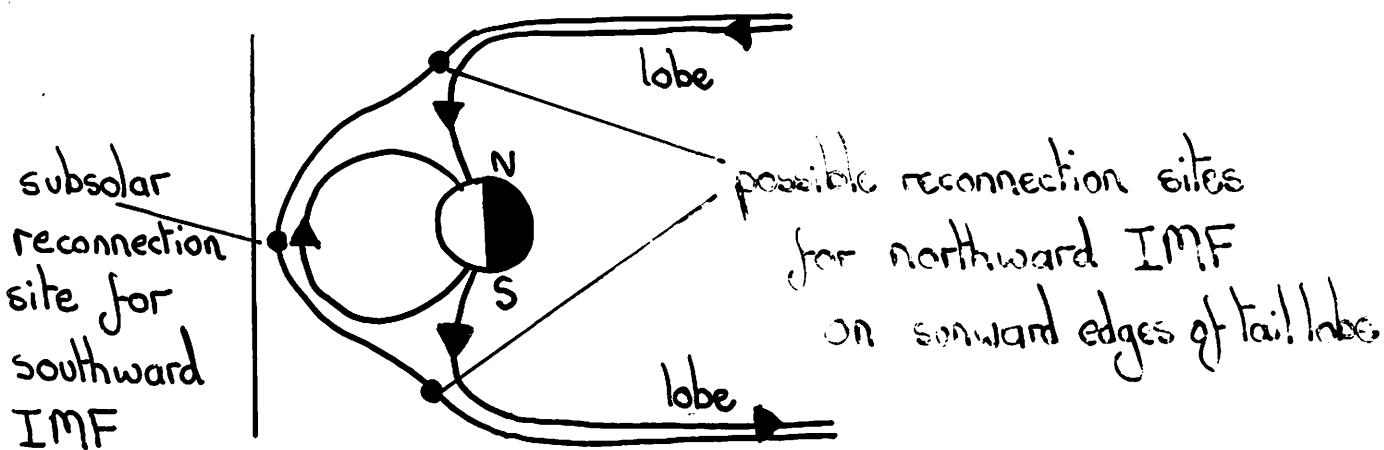


Fig. 9



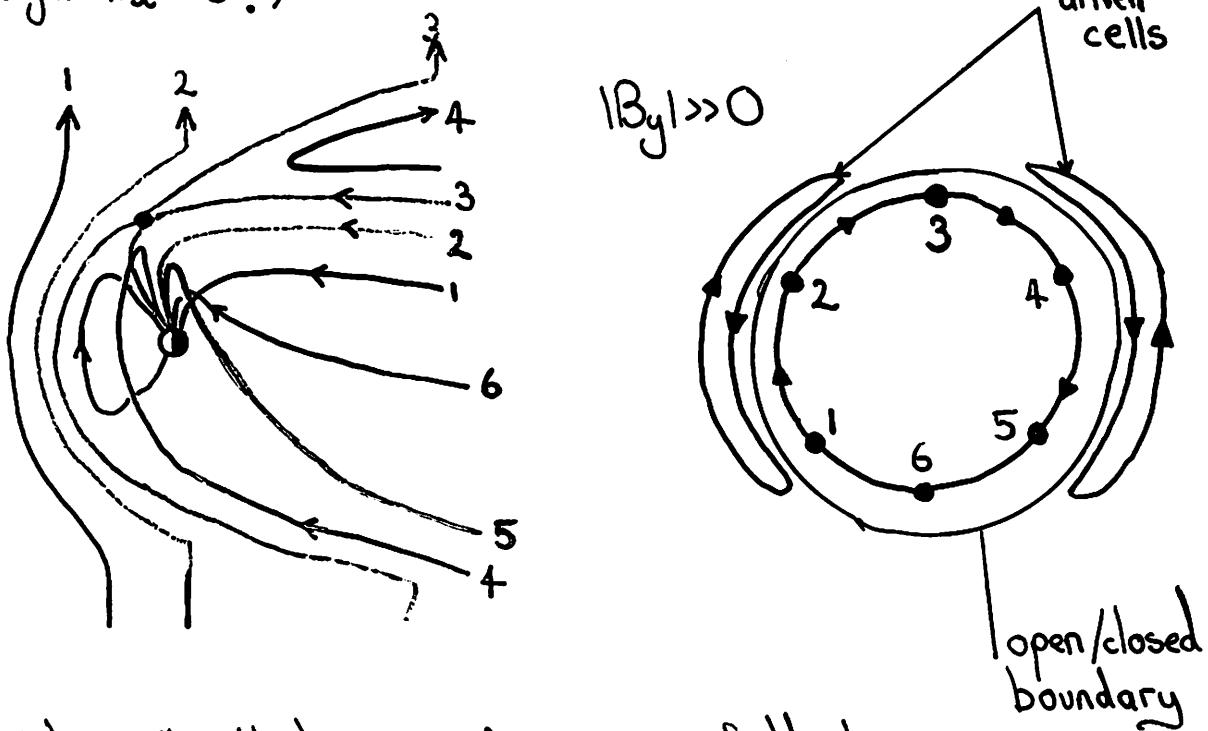
24 years' IMF data (hourly averages)

CONVECTION FOR NORTHWARD IMF



Introduces "lobe cells" into polar cap convection,
(NB. I assume open lobes never disappear completely)

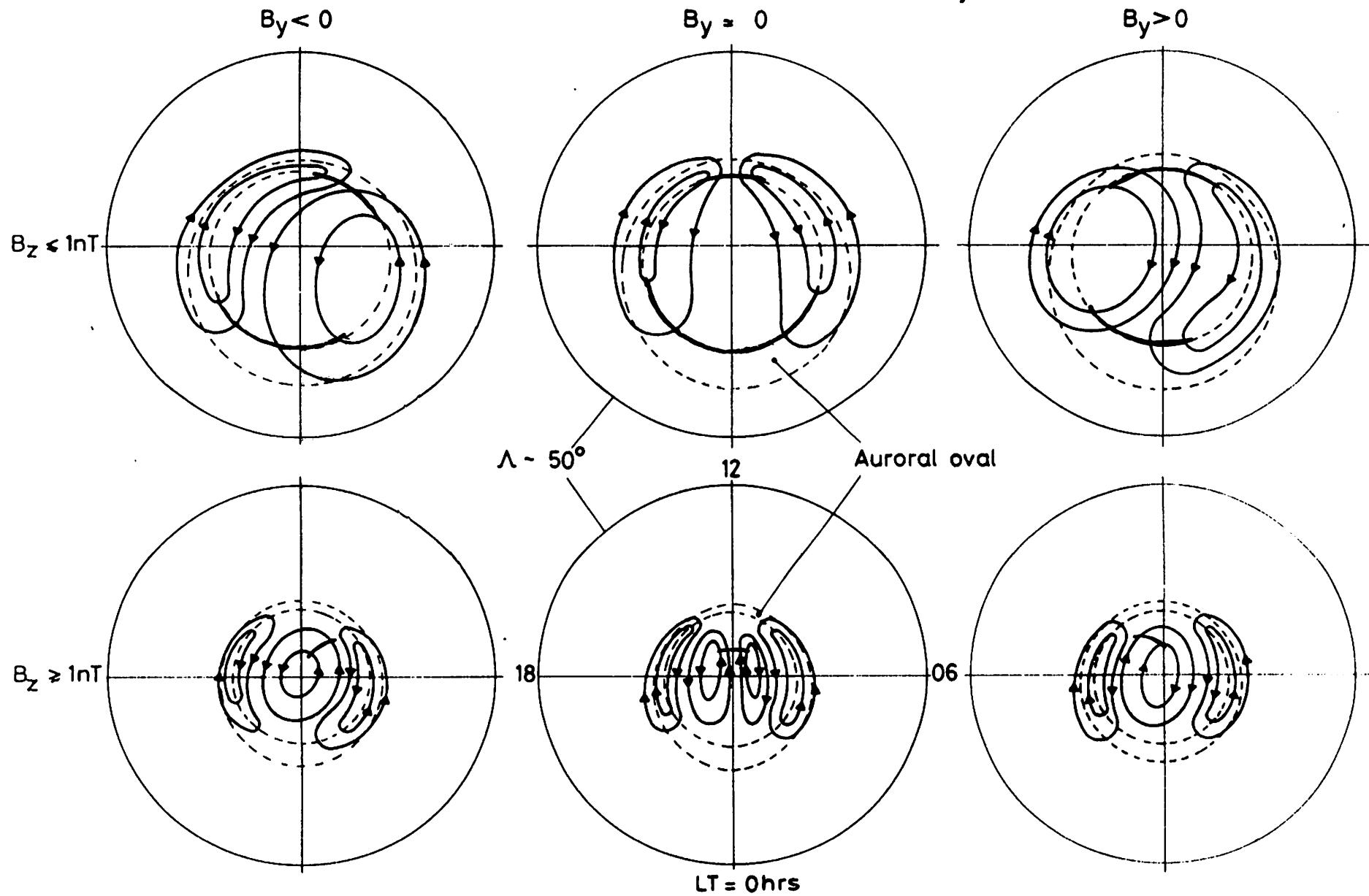
If reconnection occurs at one or other lobe, possibly
depending on IMF B_x (?) (Northern hemisphere
favoured for $B_x < 0$?)



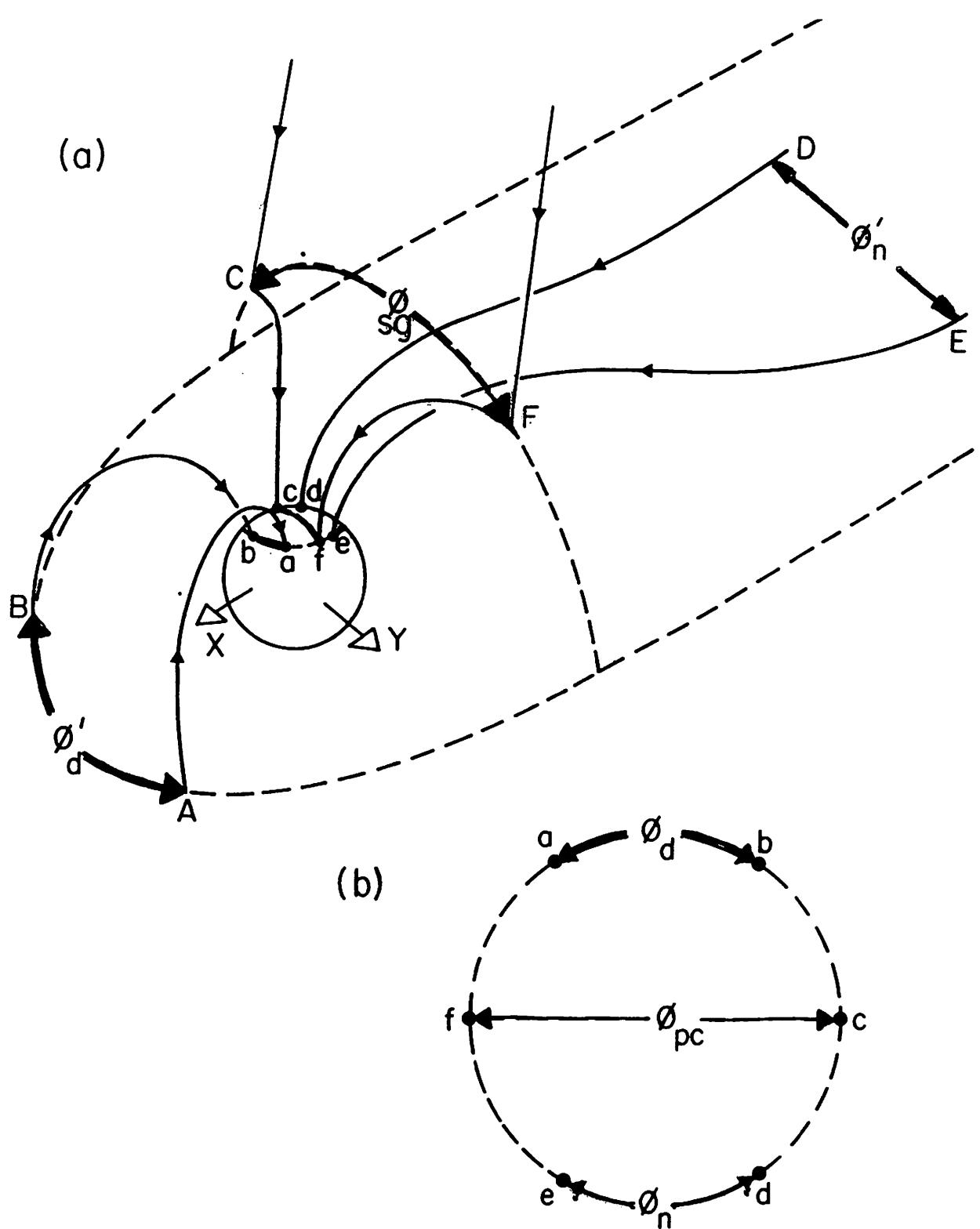
This "stirring" effect re-configures open field lines
and causes circulation on open field lines.

Inferred convection patterns in the Northern hemisphere

\odot $B_z = \text{Northward}$
 $\leftarrow B_y = \text{Dawn-to-dusk}$ } Component
 of IMF



All steady-state patterns because no account taken of IMF or magnetospheric history



Only in steady-state:

$$\phi_{pc} = \phi_{sg} = \phi_d = \phi'_d = \phi_n = \phi'_n$$

Generally, apply Faraday's Law, e.g. to cCFf

$$\phi_{pc} - \phi_{sg} = \frac{d}{dt} \int_{cCFf} \bar{B} \cdot d\bar{A} \neq 0 \quad (E_{||} \text{ negligible})$$

Polar Cap Expansion / Contraction

shows steady state does not always (indeed often) apply

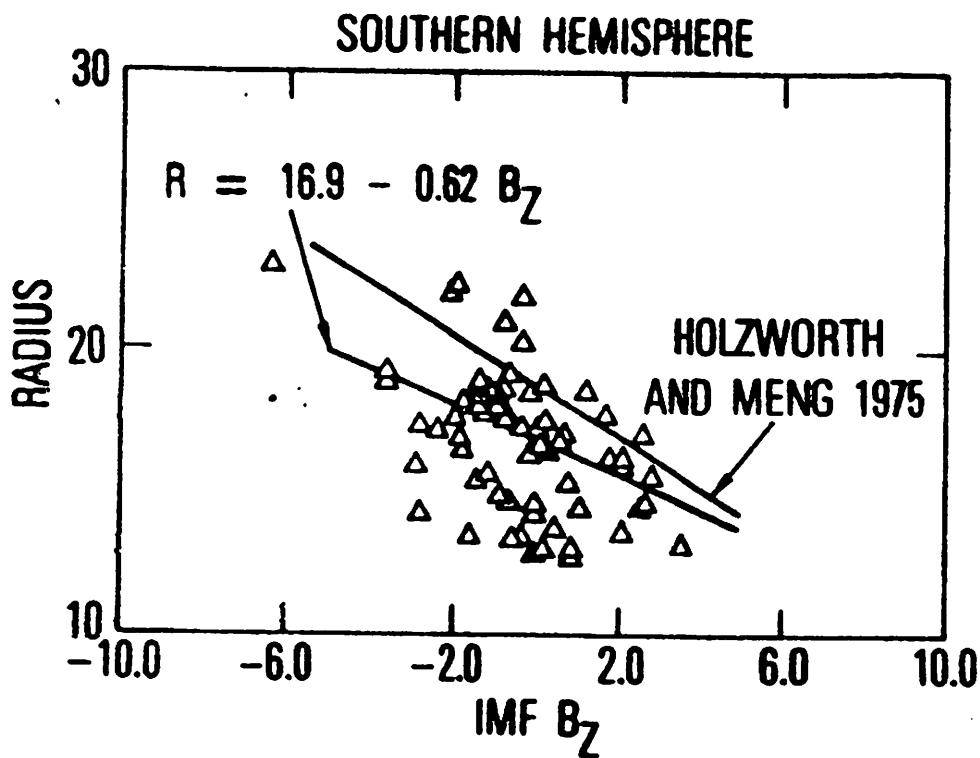


FIG. 3. COMPARISON OF CIRCLE RADIUS R FROM THIS STUDY AND PREVIOUS STUDY.

$$\text{Polar Cap Flux, } F = B_i \pi R^2$$

$$R = 16.9 - 0.62 B_Z \text{ (deg.)}$$

$$\text{gives } F = (571.2 - 46 B_Z + 0.76 B_Z^2) \times 10^6 \text{ Wb}$$

$(B_Z \text{ in nT})$

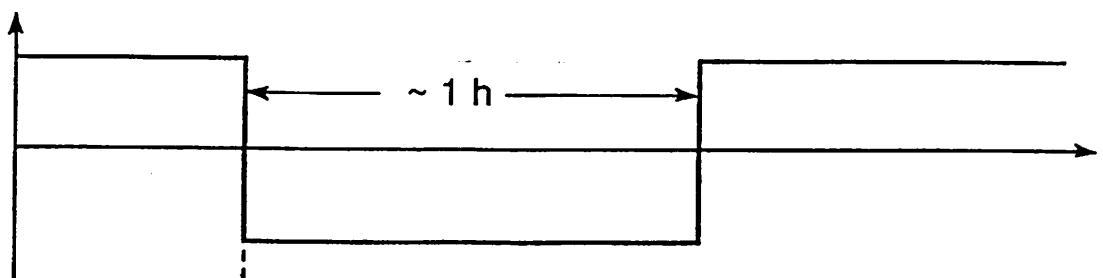
$$B_Z = +5 \text{nT}, \quad F \approx 3.6 \times 10^8 \text{ Wb}$$

$$B_Z = -5 \text{nT}, \quad F \approx 8.2 \times 10^8 \text{ Wb}$$

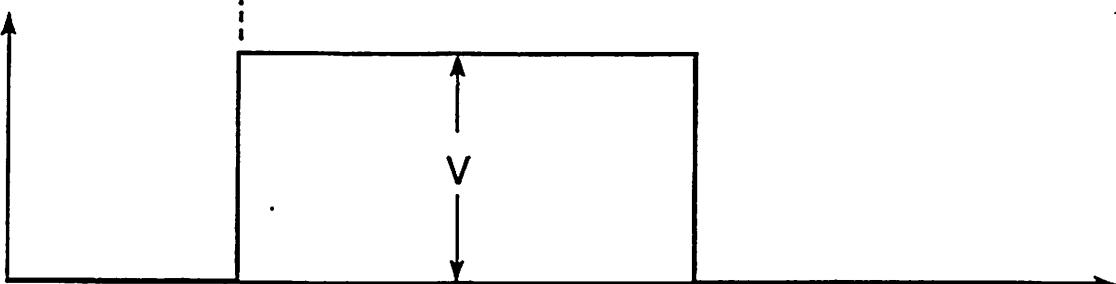
$$\Delta F / \Delta t \sim 130 \text{kV} = |\phi_d - \phi_n| \quad (\text{for } \Delta t = 1 \text{ hr.})$$

Effect of propagation delay between dayside and nightside
 X-line ($\gtrsim 30$ min.) :-

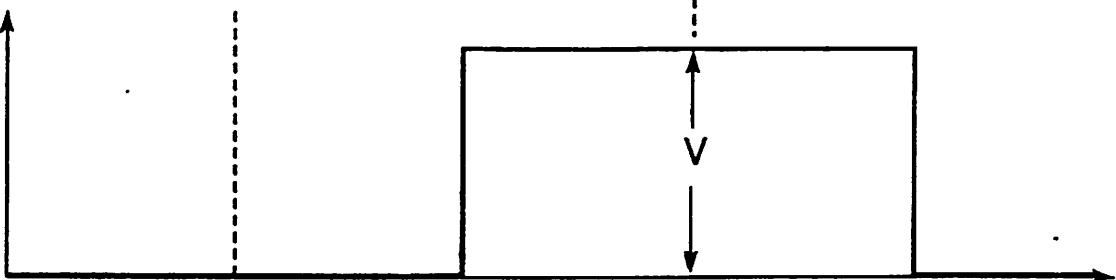
a Subsolar B_z



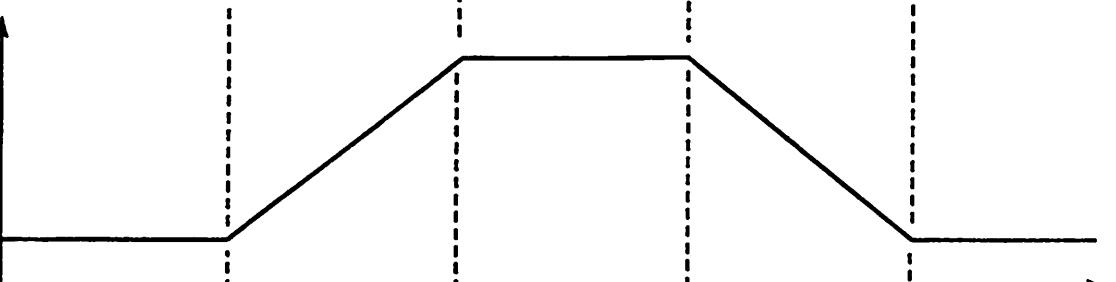
b Dayside NL Voltage



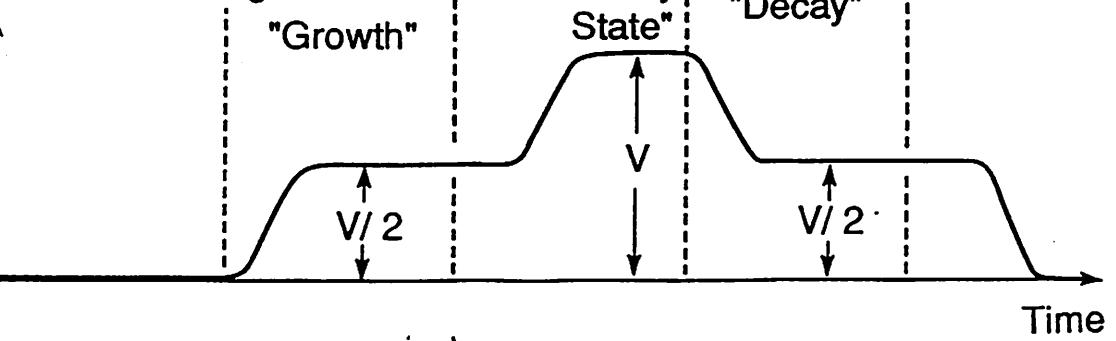
c Nightside NL Voltage



d Open Flux (Area PC)



e Central PC Voltage

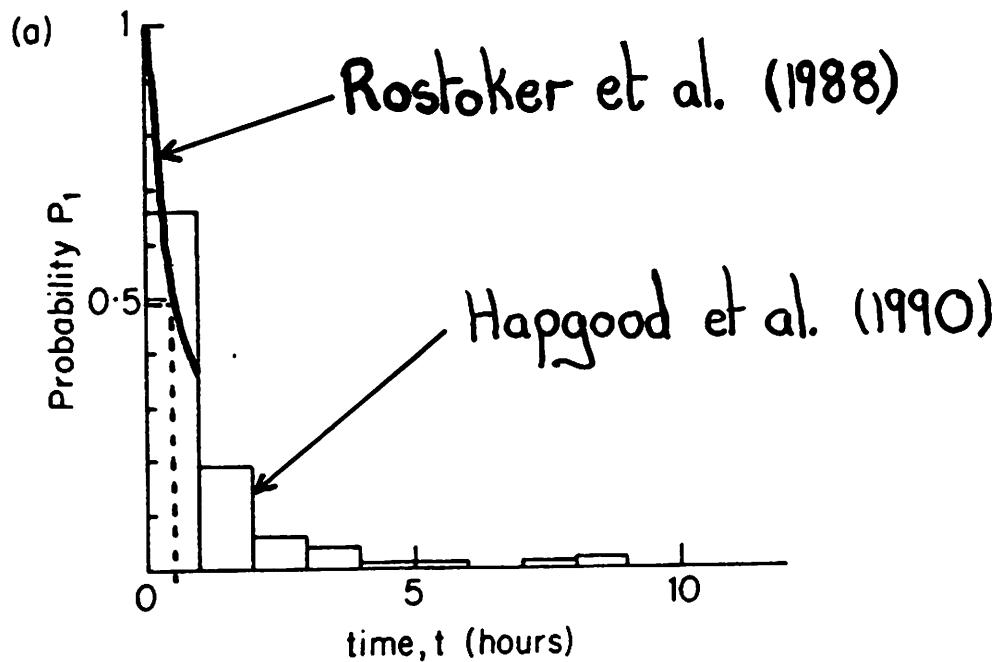


Time

in fact unlikely that $\phi_d' = \phi_n'$, even after
 ϕ_n' has responded \therefore this steady state unlikely

Stability of Polarity of IMF B_z

6



P_i is the probability that the IMF B_z component remains constant for a time t

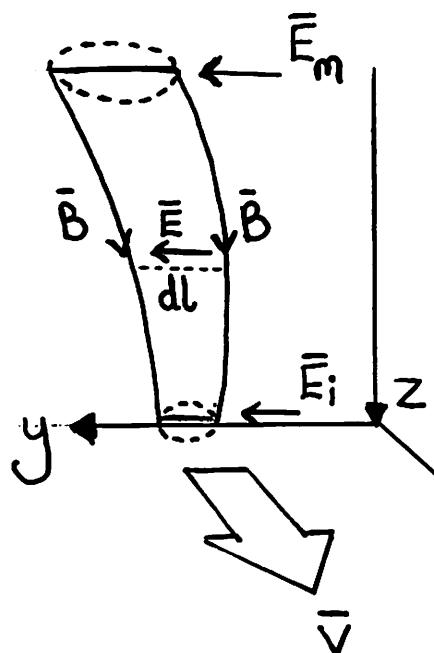
$$P_i \approx 0.5 \text{ for } t = 30\text{min.}$$

Cowley and Lockwood point out that information about ϕ_d' cannot reach DE for $t^* \gtrsim 30\text{min}$. Probability that ϕ_d' has changed polarity by then, $P_i^* \gtrsim 0.5$

Hence even if $\phi_n'(t^*) = \phi_d'(0)$ as soon as information reaches DE, $\phi_n'(t^*) \neq \phi_d'(t^*)$ i.e. NOT STEADY STATE

Electric field mapping

FLUX TUBE from magnetopause (m) to ionosphere (i)



$$\text{flux, } F = B \cdot \pi r^2 / 4$$

is constant

$\therefore \frac{dl}{\sqrt{B}}$ is constant w.r.t. s
(distance along flux tube)

Ideal MHD so $E_{||} = 0$

Faraday's Law
yields:

$$\oint \bar{E} \cdot d\bar{l} = E_m dl_m - E_i dl_i = \int_i^m \frac{\partial B_x}{\partial t} \cdot dl \cdot ds$$

$$\frac{E_m}{\sqrt{B_m}} - \frac{E_i}{\sqrt{B_i}} = \int_i^m \left(\frac{\partial B_x}{\partial t} \right) \frac{1}{\sqrt{B}} ds$$

In STEADY-STATE $(\partial B_x / \partial t) = 0$

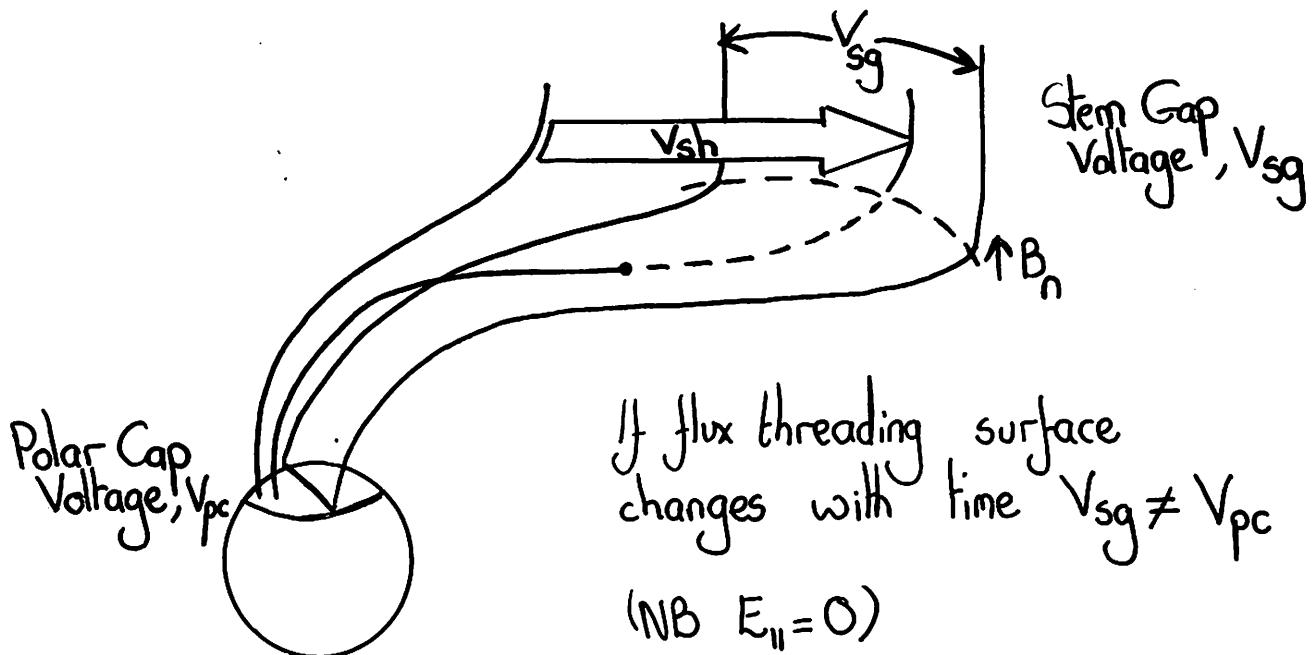
$(\therefore \text{curl } \bar{E} = 0, \text{ often called curl-free approximation})$

$$\frac{E_m}{\sqrt{B_m}} = \frac{E_i}{\sqrt{B_i}}$$

Often used but only applies to steady-state limit.

Non - Steady Effects

and mapping the solar wind electric field down open field lines.



$$\text{If } V_{sh} = 450 \text{ km s}^{-1}, B_n = 5 \text{ nT} \quad E_m = v_{sh} B_n = 2.25 \text{ mV m}^{-1}$$

$$(\text{For a Stern Gap width of } 7 R_E, V_{sg} = 2.25 \times 10^{-3} \times 7 \times 6370 \times 10^3 \approx 100 \text{ kV})$$

[In steady-state]

$$V_{sg} = V_{pc}$$

$$E_m / \sqrt{B_m} = E_i / \sqrt{B_i}$$

$$\text{for } B_m \approx 50 \text{nT}; B_i \approx 5 \times 10^{-5} \text{T}$$

$$E_i = 70 \text{ mV m}^{-1} \quad v_i = E_i / B_i \approx 1.4 \text{ km s}^{-1}$$

$$\frac{E_m}{\sqrt{B_m}} - \frac{E_i}{\sqrt{B_i}} = \int_i^m \left(\frac{\partial B_x}{\partial t} \right) \frac{1}{\sqrt{B}} ds$$

Let us consider $(\partial B_x / \partial t)$ needed to make $E_i = 0$ for this $E_m = 2.25 \text{ mV m}^{-1}$ and $B_m = 50 \text{nT}$, by assuming $(\partial B_x / \partial t)$ is independent of s

Can then use a model field (Tsyganenko T87 model) to compute $\int_i^m \frac{1}{\sqrt{B}} ds$

For Stern Gap at $x = -15R_E$, $\int_i^m B^{-1/2} ds \approx 10^{12} \text{ mT}^{-1/2}$

$$E_i = 0 \text{ if } \frac{\partial B_x}{\partial t} = \frac{2.25 \times 10^{-3} / (50 \times 10^{-9})}{10^{12}}^{1/2} = 10^{-11} \text{ Ts}^{-1}$$

$$= 0.01 \text{ nT s}^{-1}$$

Substorm growth phase typically lasts $\sim 45 \text{ min}$
 $= 2700 \text{ s}$.

during which $\Delta B_x \sim \left(\frac{\partial B_x}{\partial t} \right) \Delta t$
 $\sim 27 \text{nT}$

Field changes of this magnitude are observed during substorm growth phases (eg. Fairfield and Ness, 1970; Fritz, 1984) implying Stern Gap voltage of 100 kV need not cause any ionospheric flow, throughout a growth phase lasting 45 min. (instead, magnetic energy is stored in the tail)

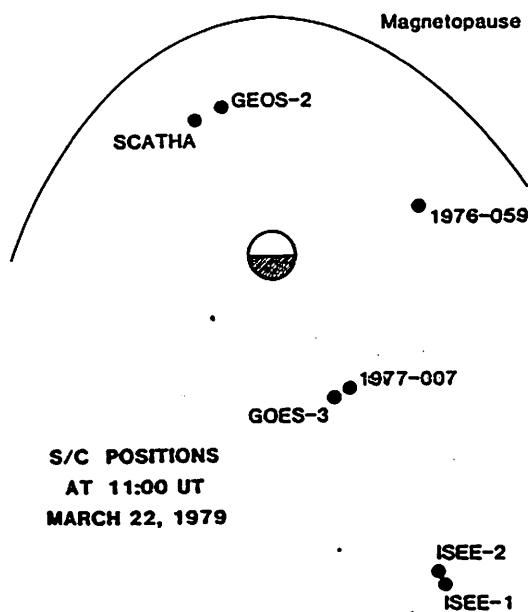


Figure 1. Location of seven spacecraft during the CDAW-6 substorm of 11 UT on March 22, 1979. Note the radial alignment of four satellites along the 0200 LT meridian-with ISEE-1 at a GSE position of $(-13.6 R_g, -7.2 R_g, 0.6 R_g)$ and ISEE-2 at $(-12.1 R_g, -7.1 R_g, 0.3 R_g)$. The relative positions of these satellites changed very little during the event discussed here.

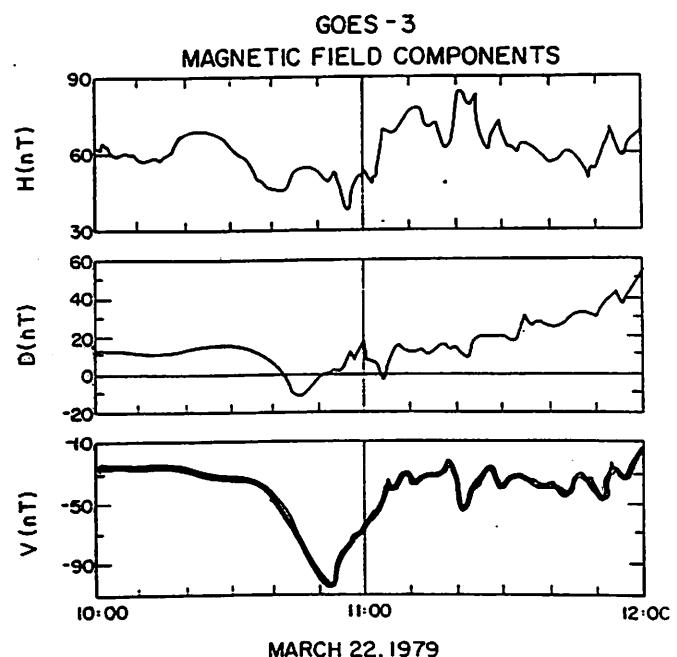


Figure 3. Magnetic field measurements made by the magnetometer on satellite GOES-3 during the CDAW-6 substorm of 11 UT on March 22, 1979. See text for definition of V, D, and H components.

FRITZ ET AL. 205

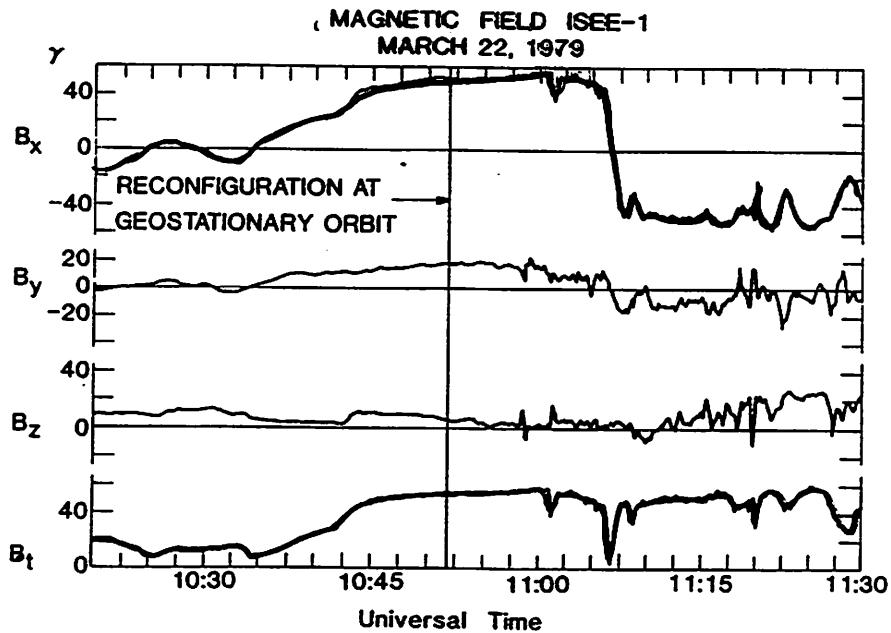
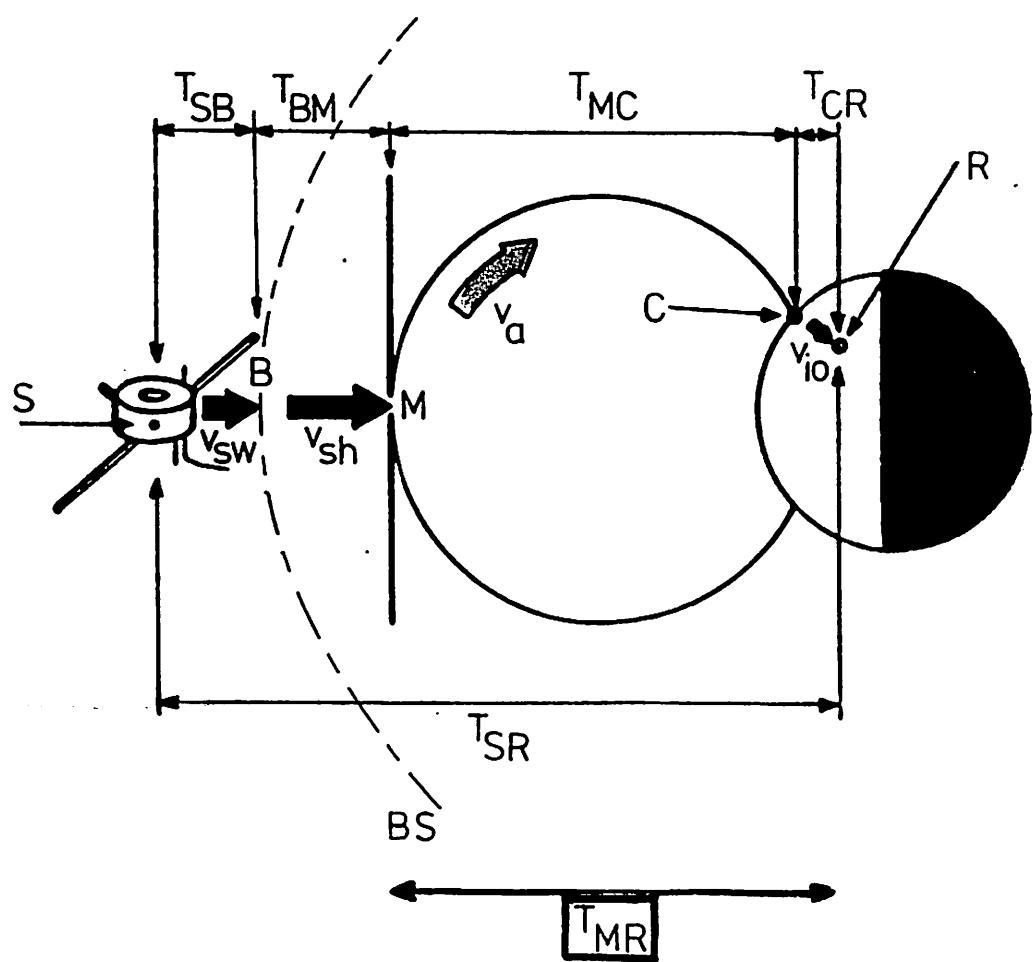


Figure 5. Magnetic field measurements in GSE coordinates made by the magnetometer on satellite ISEE-1 during the CDAW-6 substorm of 11 UT on March 22, 1979.

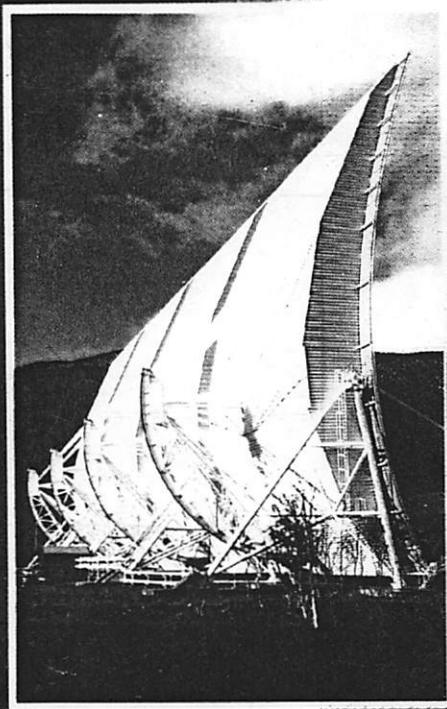
Fritz et al. (1984)



Lockwood et al., 1986



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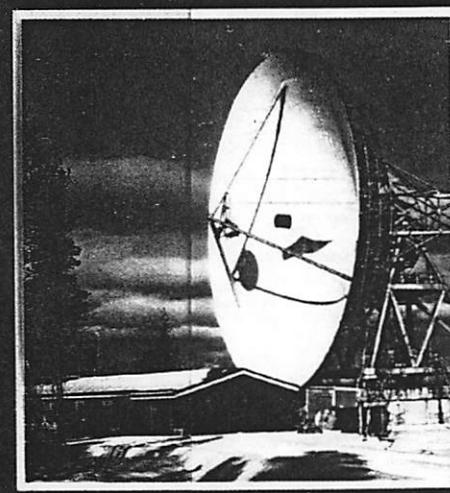


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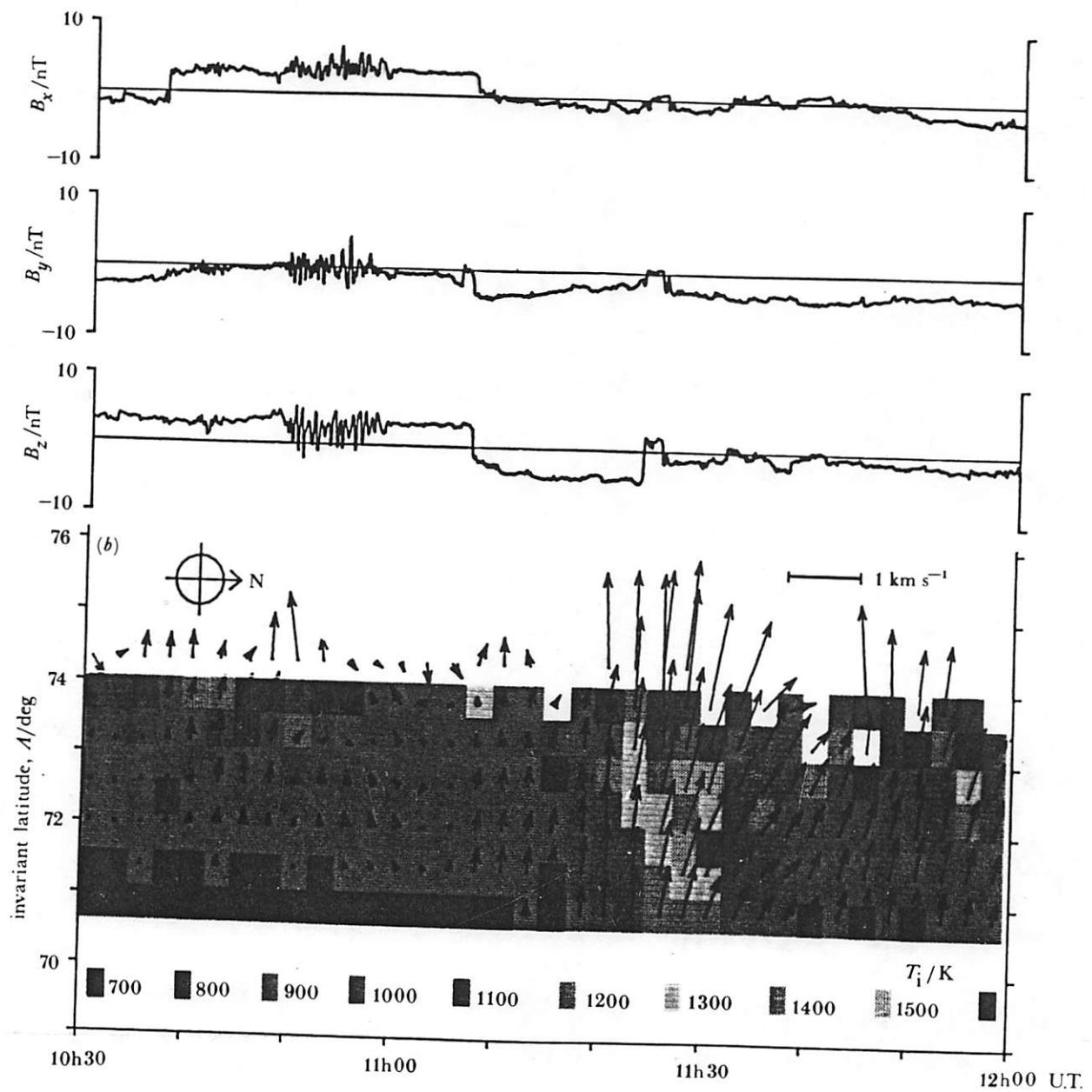


FIGURE 2. Example of an observed ionospheric response to a southward turning of the IMF (27 October 1984).
 (a) The sunward, downward and northward components of the IMF in GSM coordinates (B_x , B_y and B_z , respectively), observed by AMPTE-UKS when located immediately sunward of the Earth's bow shock. (b) The simultaneous observations by the EISCAT radar. The flow vectors have been rotated through 90° to avoid congestion of the plot; hence northward flow is shown by vectors directed to the right of the figure and westward flow by vectors directed upward (i.e. the vectors are in the direction of the electric field). The vectors are superimposed on a colour plot of two-minute averages of the ion temperature, T_i . (From Willis *et al.* 1986.)

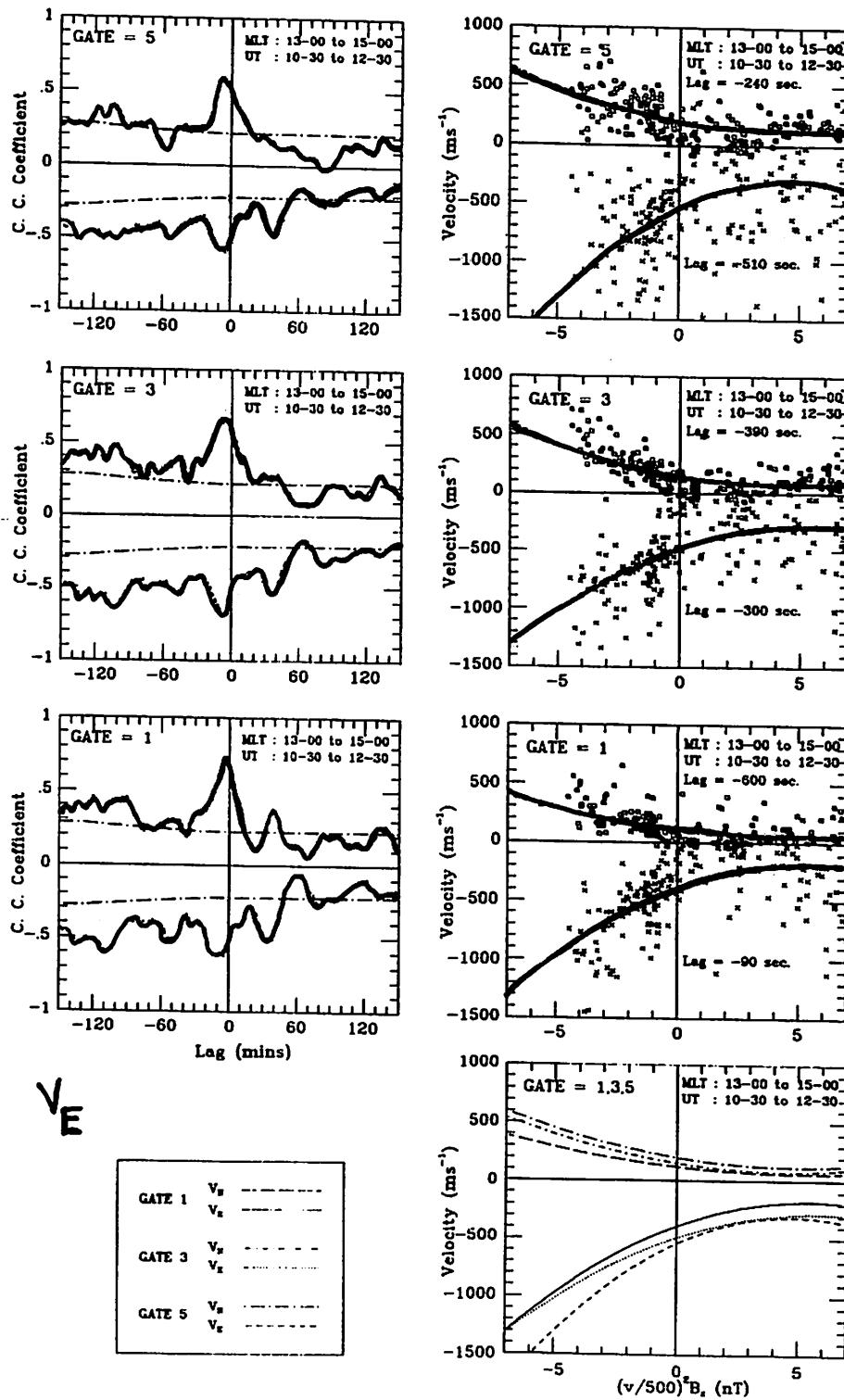
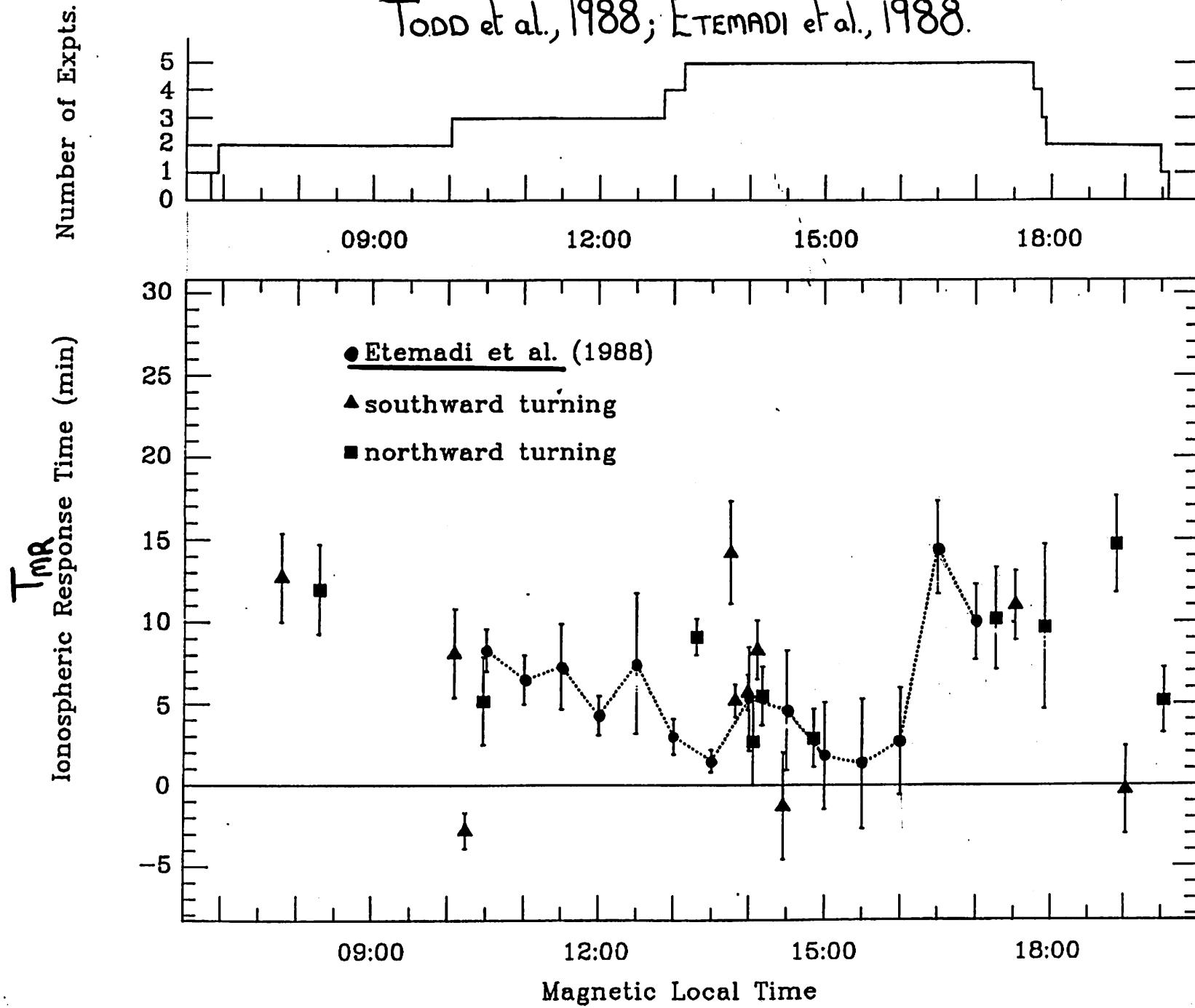
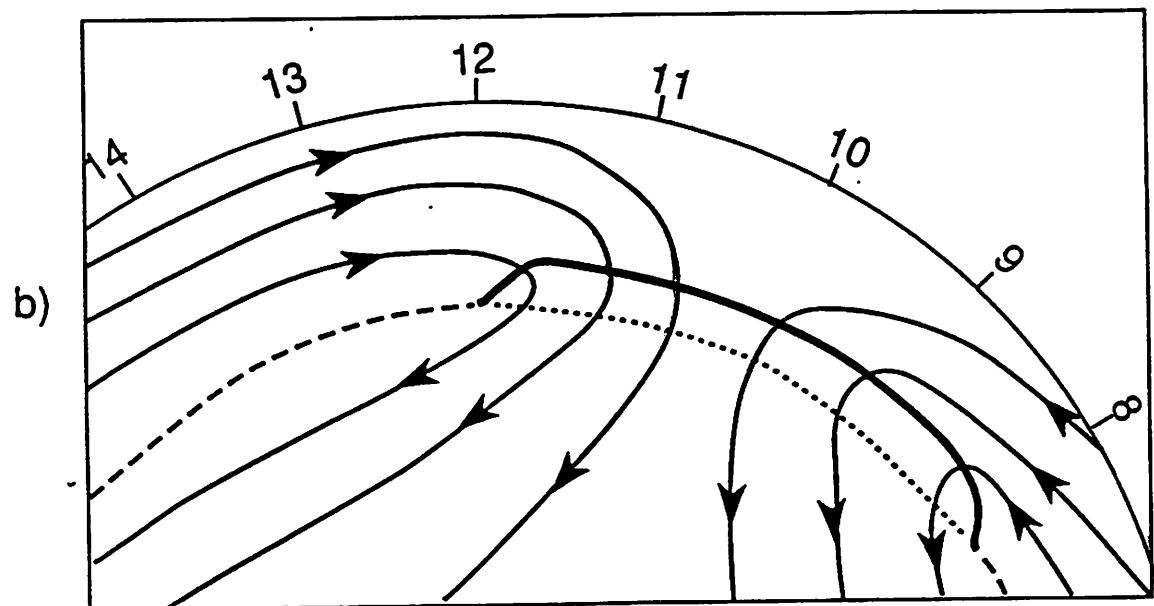
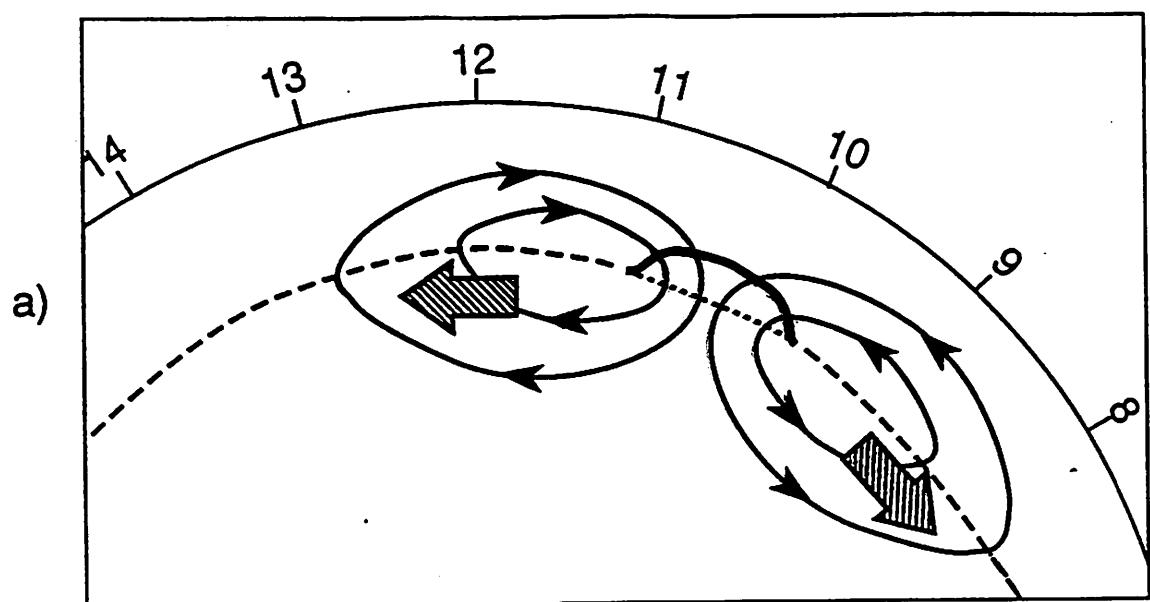


FIG. 4c. AS FOR FIG. 4a EXCEPT FOR THE LOCAL TIME INTERVAL FROM 13:00 TO 15:00 M.L.T.





----- Polar Cap Boundary Phase Motion
 Old Polar Cap Boundary ————— Reconnection X-line
 mapped to ionosphere

Phase velocity initially $\sim 10 \text{ km s}^{-1}$, falls with MLT from $\sim 12 \text{ hrs}$.

Directly Observed by Lockwood et al. (1986)
 Saunders et al. (1992)

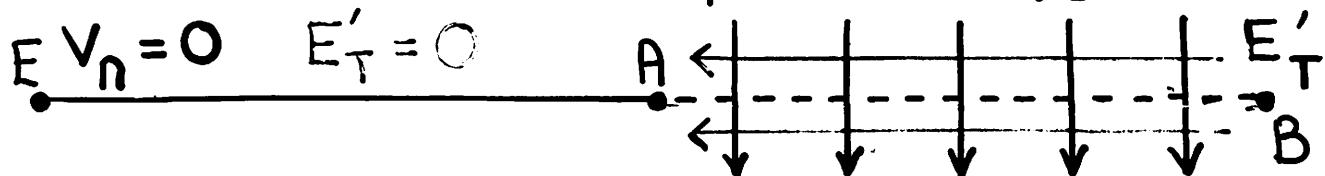
"Adiabatic" segments of open /closed boundary

(meaning "not flowing across") - Introduced by Siscoe and Huang (1985).

Open /closed boundary, in its OWN REST FRAME

$$\bar{v} = (\bar{E} \times \bar{B}) / B^2$$

$$\otimes \vec{B}_i$$



$$E'_T = v_n B_i$$

Adiabatic segment, EA

$$\text{Voltage, } \phi'_{EA} = \int_E^A E'_T dl = 0$$

Merging gap, AB

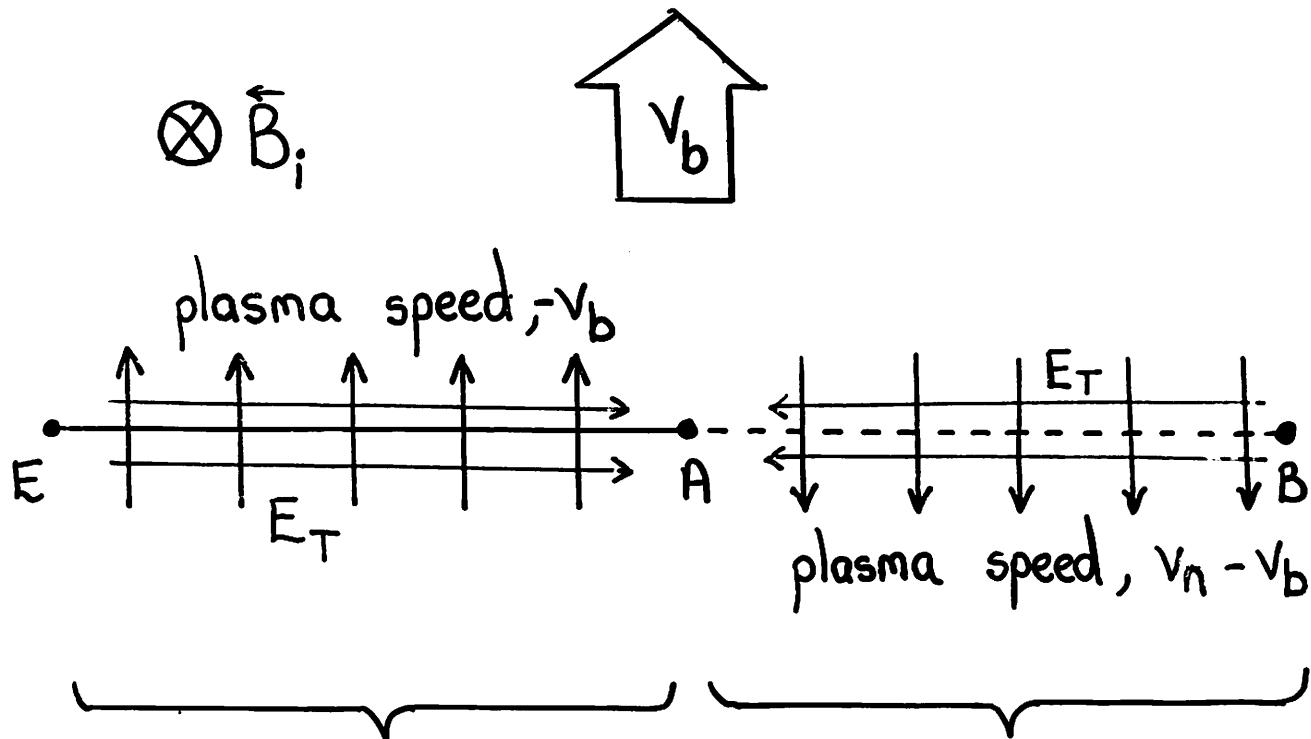
(Ionospheric projection of reconnection X-line)

$$\text{Voltage, } \phi'_{AB} = \int_A^B E'_T dl \neq 0$$

By Faraday's Law, ϕ' is a flux transfer rate

ϕ'_{AB} = voltage across corresponding reconnection X-line

Moving Open /Closed Boundary with speed v_b in EARTH'S FRAME (because polar cap expands/contracts)

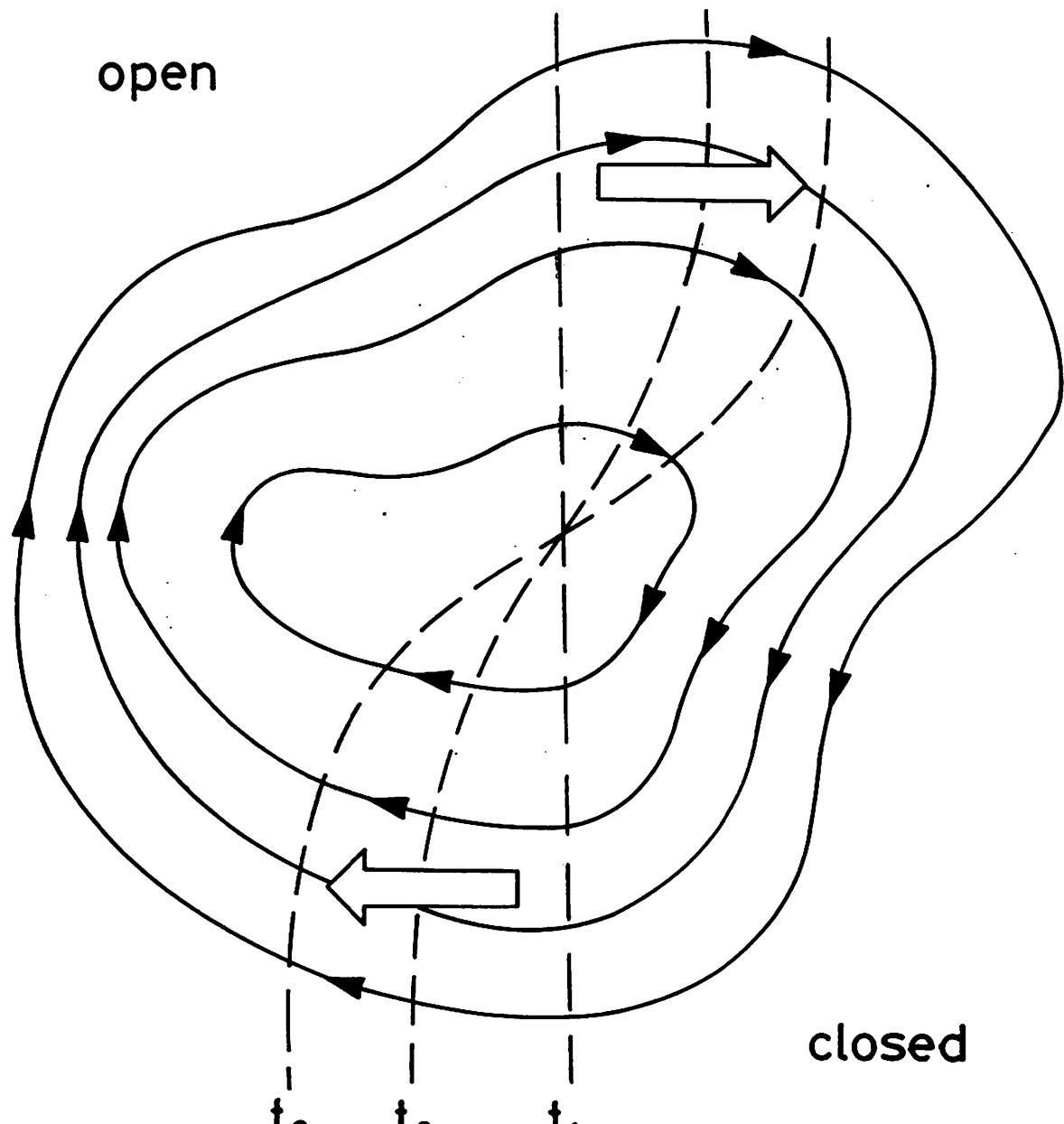


$$E_T = -v_b B_i$$

$$E_T = (v_n - v_b) B_i$$

$$\text{Voltage, } \phi_{EA} = \int_E^A -v_b B_i dl$$

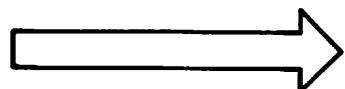
$$\begin{aligned} \text{Voltage, } \phi_{AB} &= \int_A^B (v_n - v_b) B_i dl \\ &= \phi'_{AB} - \int_A^B v_b B_i dl \end{aligned}$$



flow streamline



adiabatic open/closed boundary



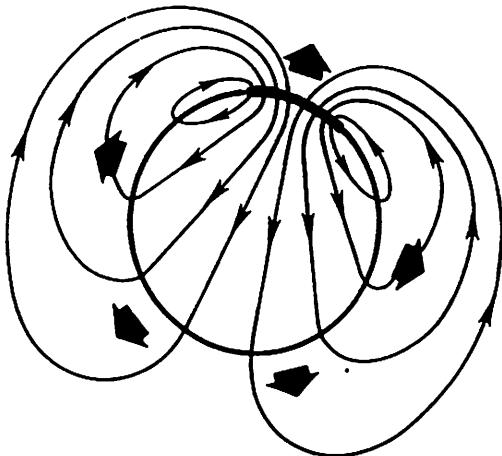
boundary motion

Lockwood and Cowley (1992)

Faraday's Law applied to polar cap:

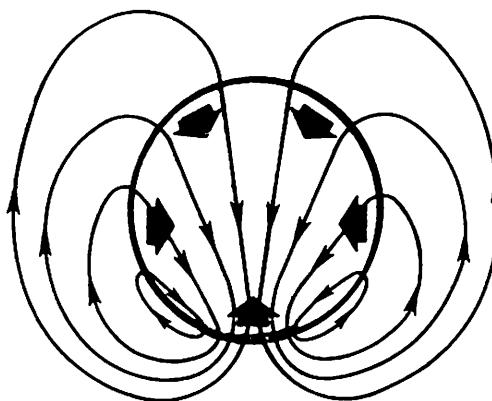
$$\oint_{pc} \mathbf{E} \cdot d\mathbf{l} = \partial(B_i A_{pc}) / \partial t = B_i \partial A_{pc} / \partial t = \Phi_d - \Phi_n$$

Expanding Polar Cap

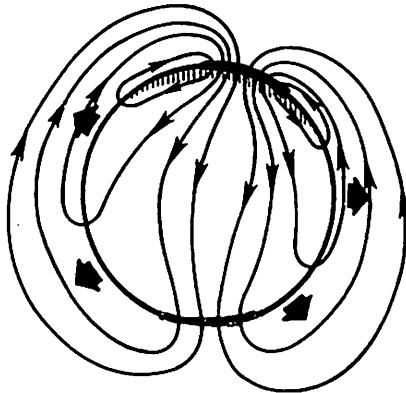


(a) $\Phi_d = 8\Delta$; $\Phi_n = 0$

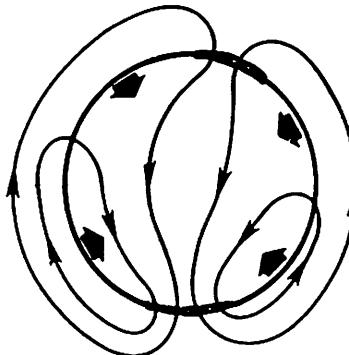
Contracting Polar Cap



(b) $\Phi_d = 0$; $\Phi_n = 8\Delta$



(c) $\Phi_d = 8\Delta$; $\Phi_n = 4\Delta$



(d) $\Phi_d = 2\Delta$; $\Phi_n = 4\Delta$

→ flow equipotential (Δ kV apart)

--- ionospheric projection of reconnection neutral line (merging gap)

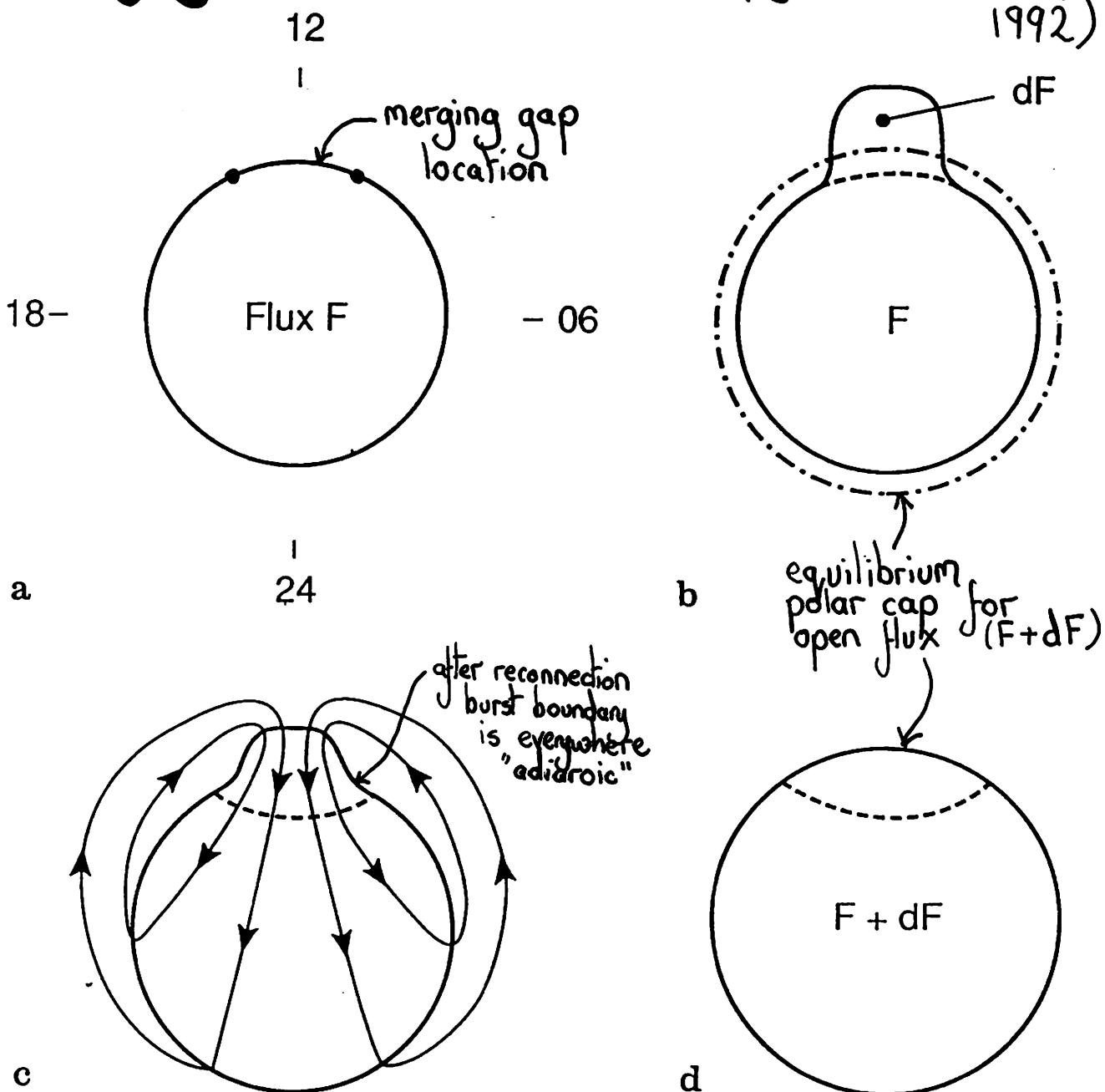
— adiabatic polar cap boundary

➡ boundary motion

Siscoe and Huang (1985)
 Lockwood and Cowley (1988)
 Lockwood et al. (1990)
 Russell (1972)

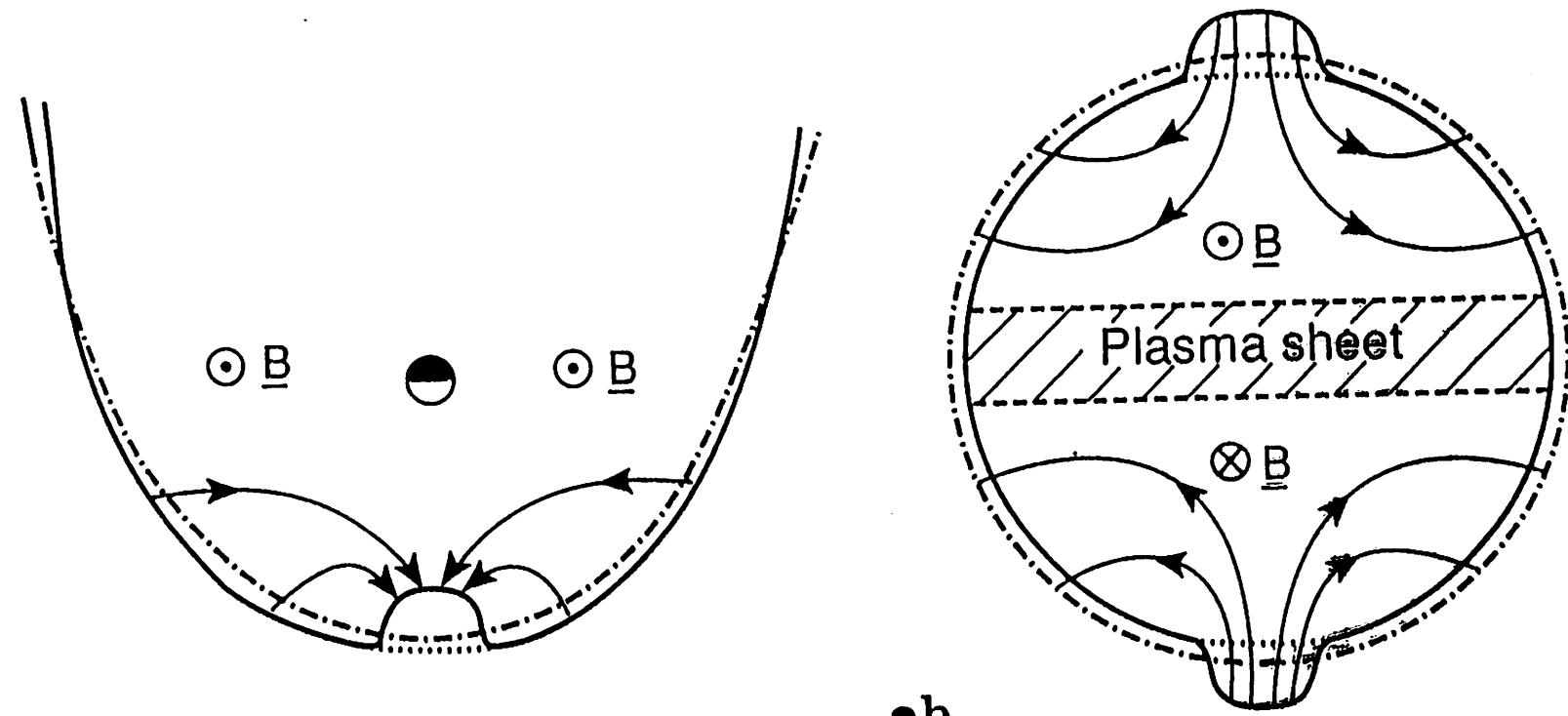
Zero-flow equilibrium configurations

(Cowley and Lockwood, 1992)



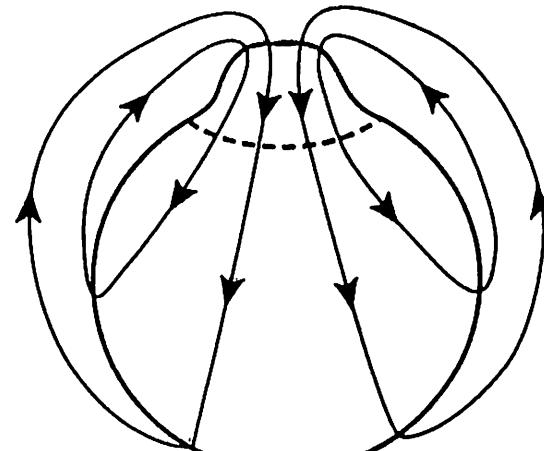
NB:

- equilibrium configurations only assumed circular here for illustrative purposes
- in reality, flux dF will not be opened instantaneously
- Lockwood et al. (1990) estimate $dF/F \lesssim 0.03$

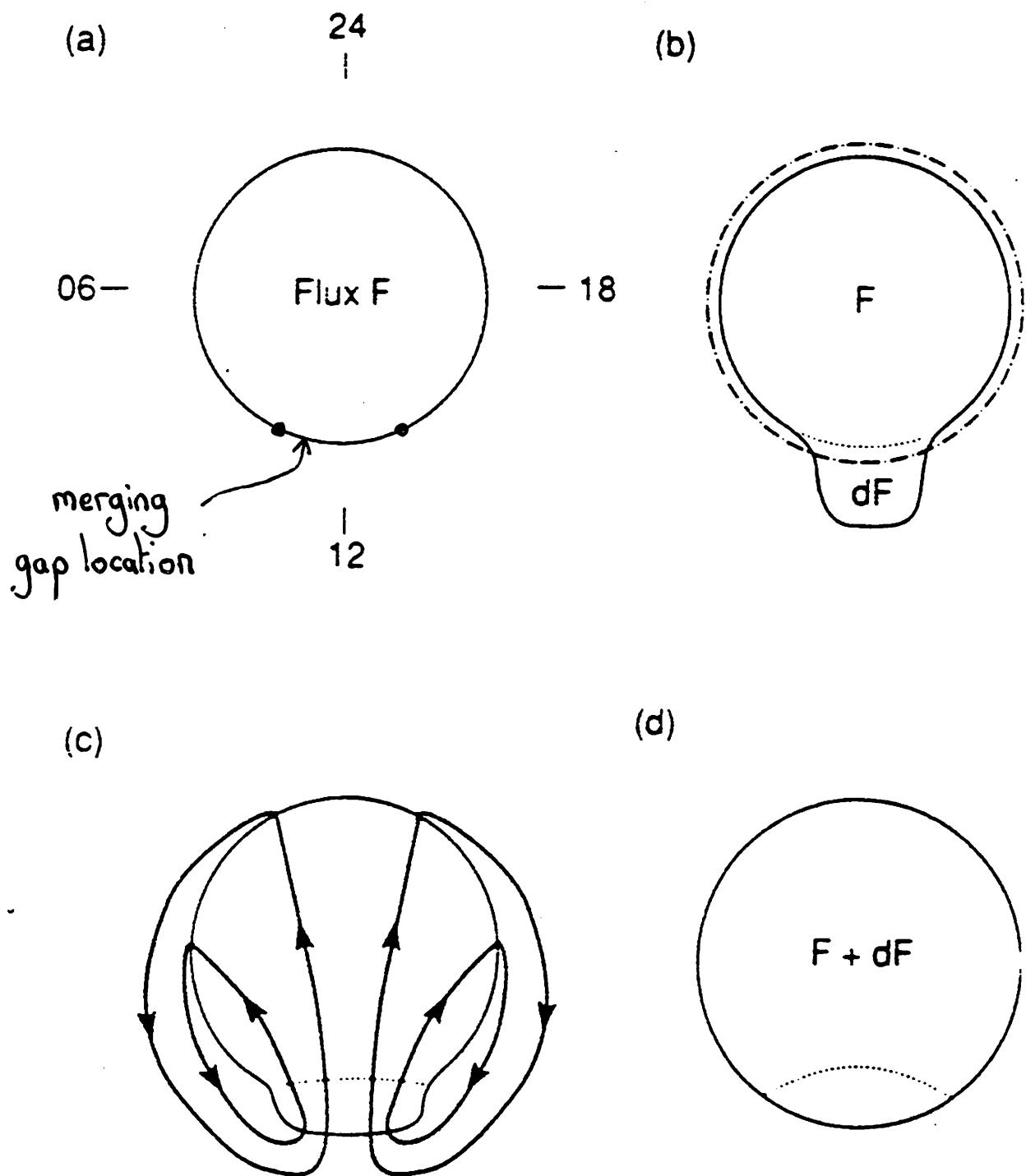


a
In the equatorial magnetosphere

b
In the geomagnetic tail



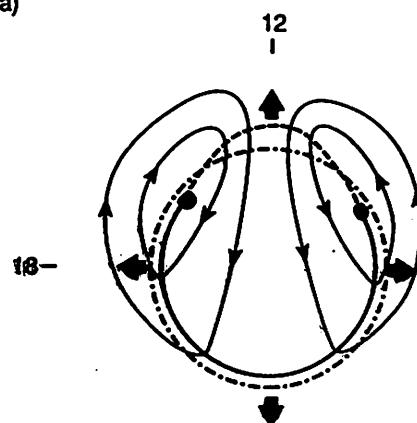
c
In the outer ionosphere



Zero-flow equilibria

(Introduced by Cowley and Lockwood, 1992)

(a)



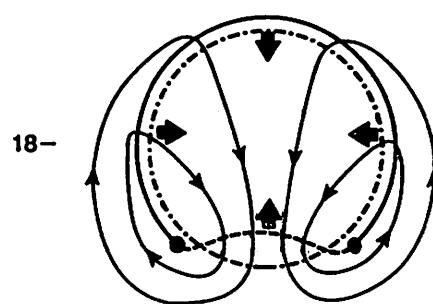
Continuous
Dayside
Reconnection

$$\phi_d > 0$$

$$\phi_n = 0$$

(b)

12
—
24



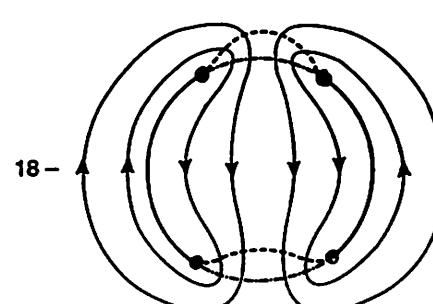
Continuous
Nightside
Reconnection

$$\phi_d = 0$$

$$\phi_n > 0$$

(c)

12
—
24



Steady State
 $\phi_d = \phi_n$

N.B. Average over a long
(several days) time scale and

$$\frac{dA_{pc}}{dt} = 0, \text{ i.e. } \phi_d = \phi_n$$

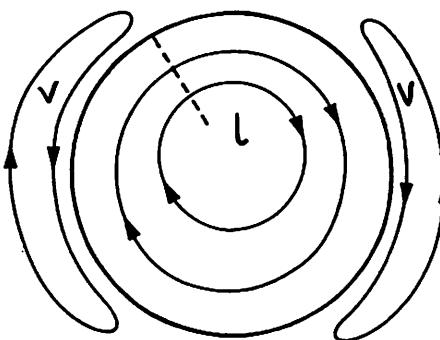
∴ Long term averages are
steady-state models.

∴ } ends of { dayside
 | nightside } merging gaps

Northward IMF Substorms

Steady-State
with $\phi_d' = \phi_n = 0$

(a)

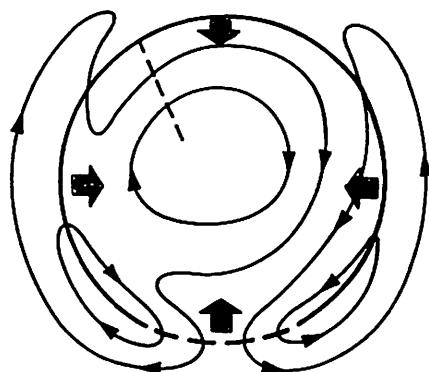


lobe circulation
cell (l) due to
reconfiguration of
already open
field lines.

v are viscously
driven cells.

$\phi_d' = 0$
 $\phi_n' > 0$
 $\phi_v > 0$

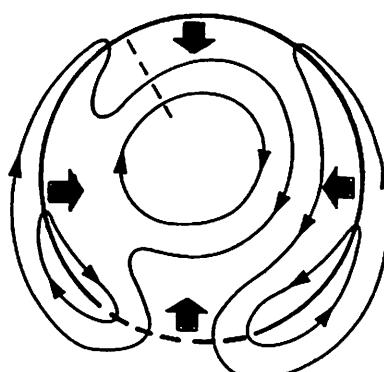
(b)



Boundary moves
so lobe and viscous
cells are no longer
distinct

$\phi_d' = 0$
 $\phi_n' > 0$
 $\phi_v = 0$

(c)



Similar pattern
if no viscous
interaction

- adiabatic open / closed boundary
- - - merging gap
- flow streamline
- ➡ boundary motion

Lockwood and Cowley (1992)

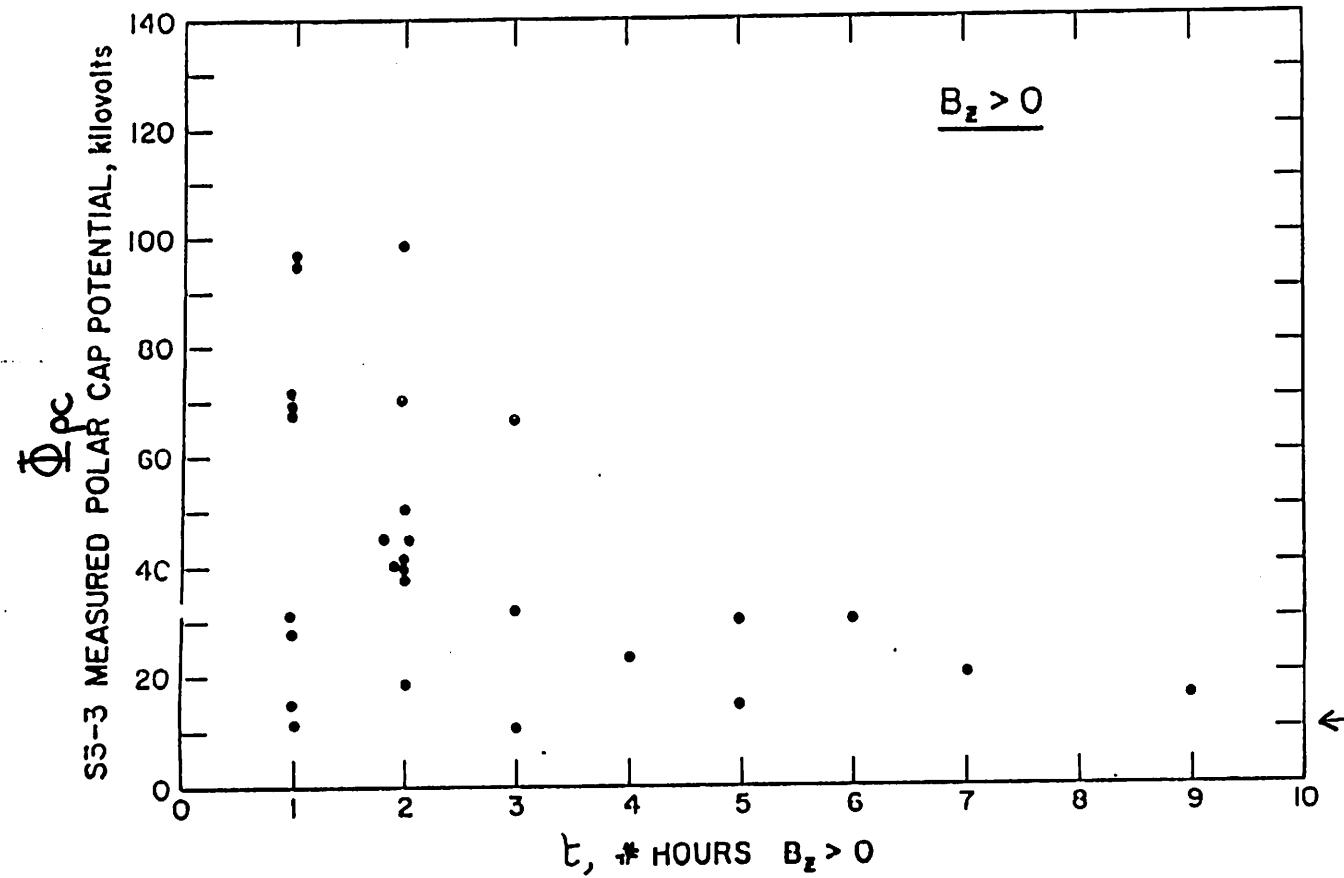


Fig. 6. The cross polar cap potential during periods when B_z^{IMF} was northward versus the number of hours previous to the measurement of the polar cap potential that B_z^{IMF} was northward.

Initial range of Φ_{pc} large, because of range of Φ_n

After ~9 hrs Φ_n is small $\Phi_{pc} \sim \Phi_v$

$\leftarrow \Phi_v \sim 10\text{kV}?$

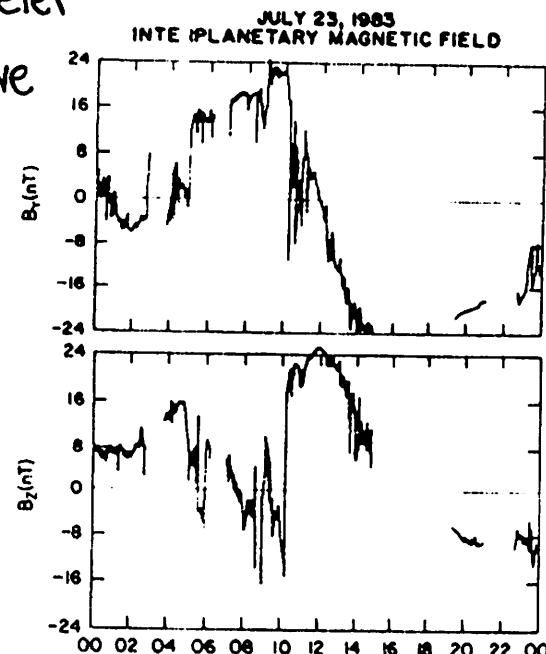
$$\phi_d = 0 \quad \text{for } t \geq 0$$

range of Δ and Φ_n , but $\Phi_{pc} \sim \Phi_n/2$ for $t \geq 0$

Φ_n decays over $t \sim 10\text{hrs}$

I.S. radar data
+ magnetometer

Observations give
"global convection
snapshots" by
the AMIE
technique:-



1. IMP-8 IMF data (GSM coordinates) for 23 July 1983.

↑
Northward IMF turning
10:15 UT at IMF-8
 \sim 10:25 UT at magnetopause

Knipp et al.: Ionospheric Convection Response

Knipp et al.
(Geophys. Res. Lett., 1991)

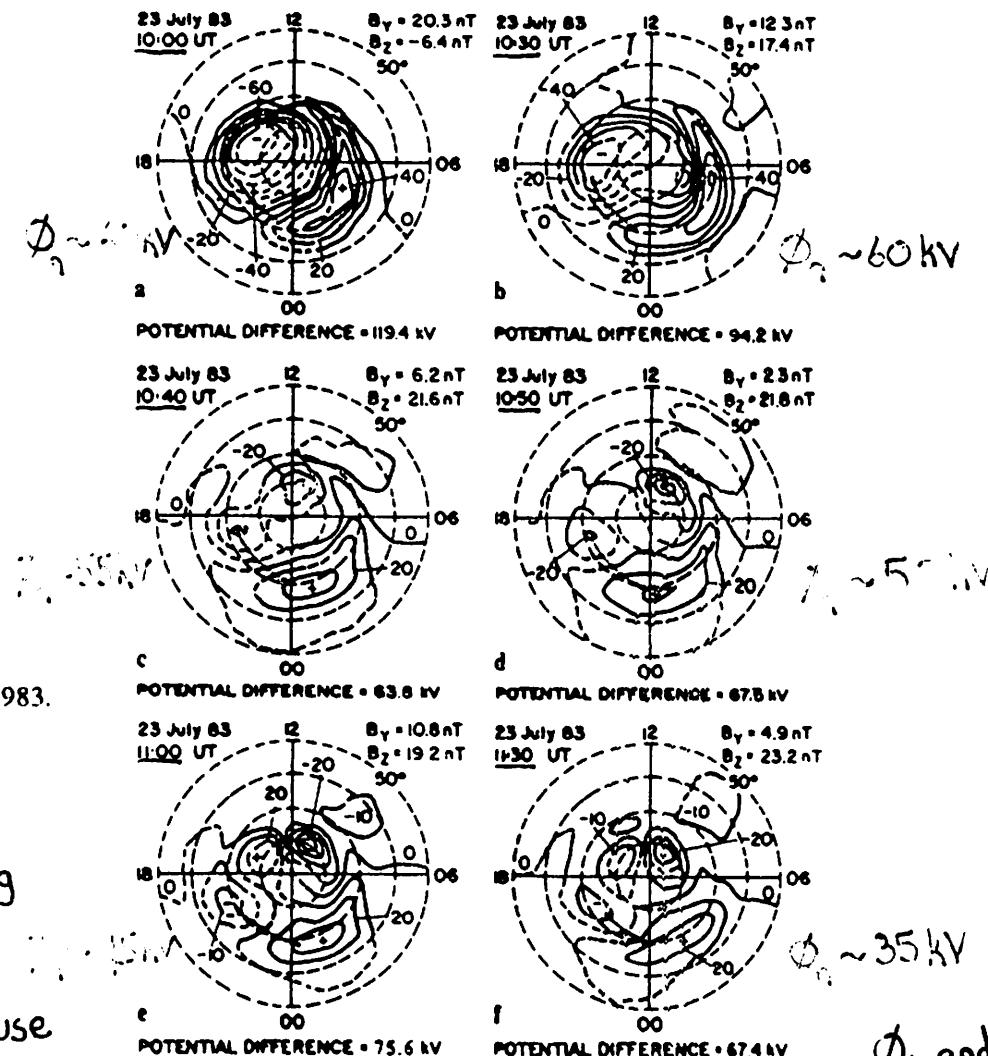
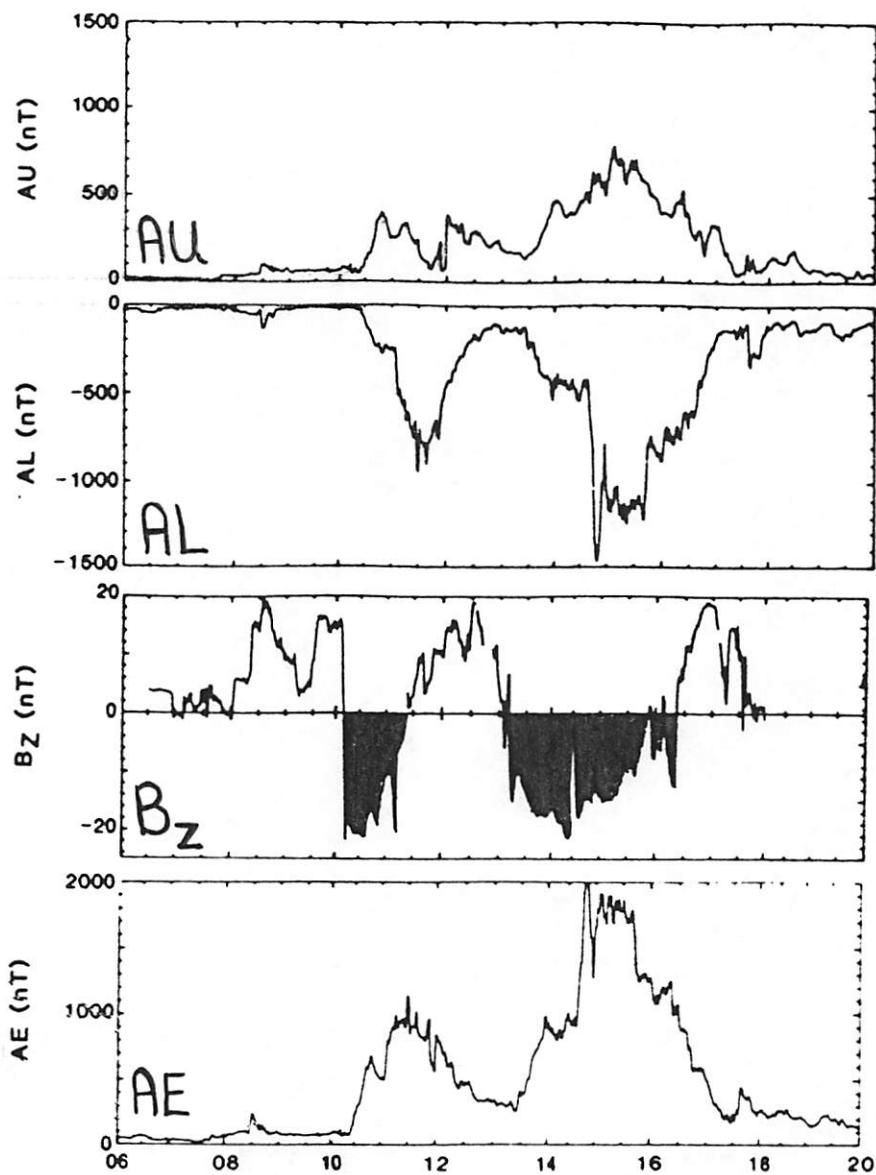


Fig. 3. Fitted electric potential patterns for selected times between 1000 UT and 1130 UT on 23 July 1983. The contour interval is 10 kV. The IMF values are shown at the upper right. In regions where the uncertainty in the large-scale electric field is 50% or greater the contours are dashed.

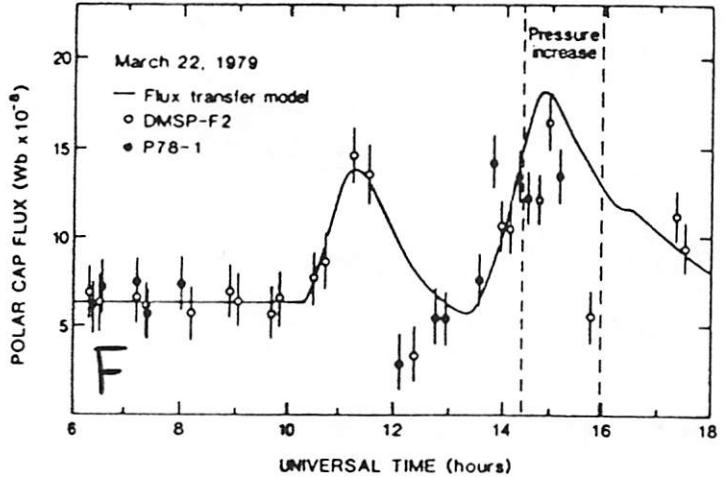
ϕ_d and dayside flow:
decay in \sim 10 min
 ϕ_n and nightside flow:
decay in several hours

March 22, 1979 Substorms
(CDAW 6)

McPherron
and Manka
(1985)



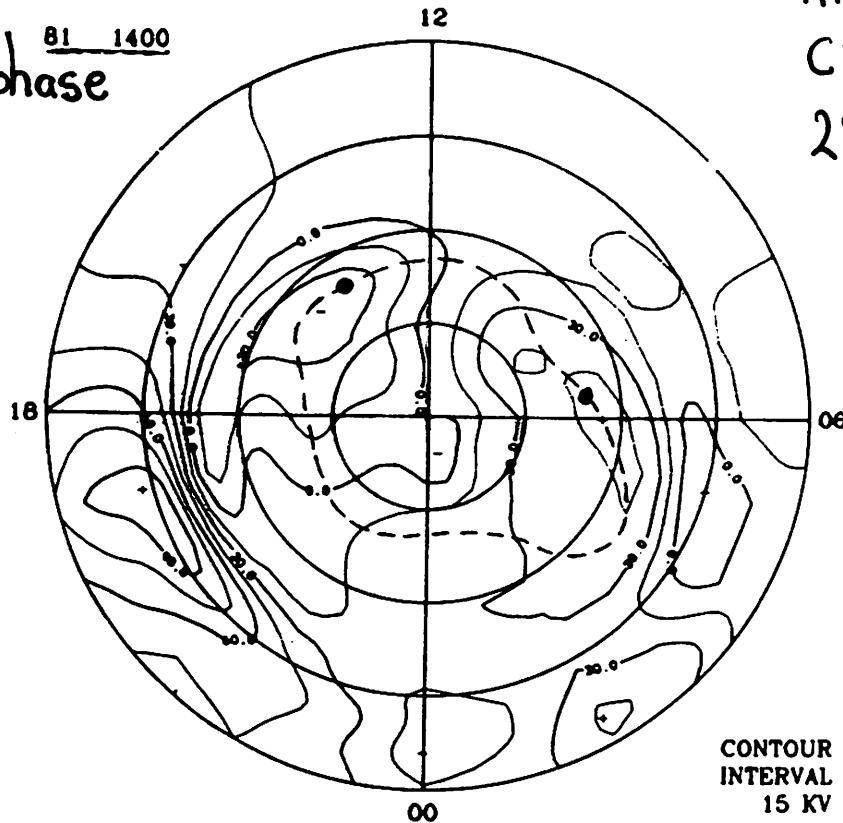
Holzer
et al. (1986)



c.f. Kamide and Baumjohann (1985)

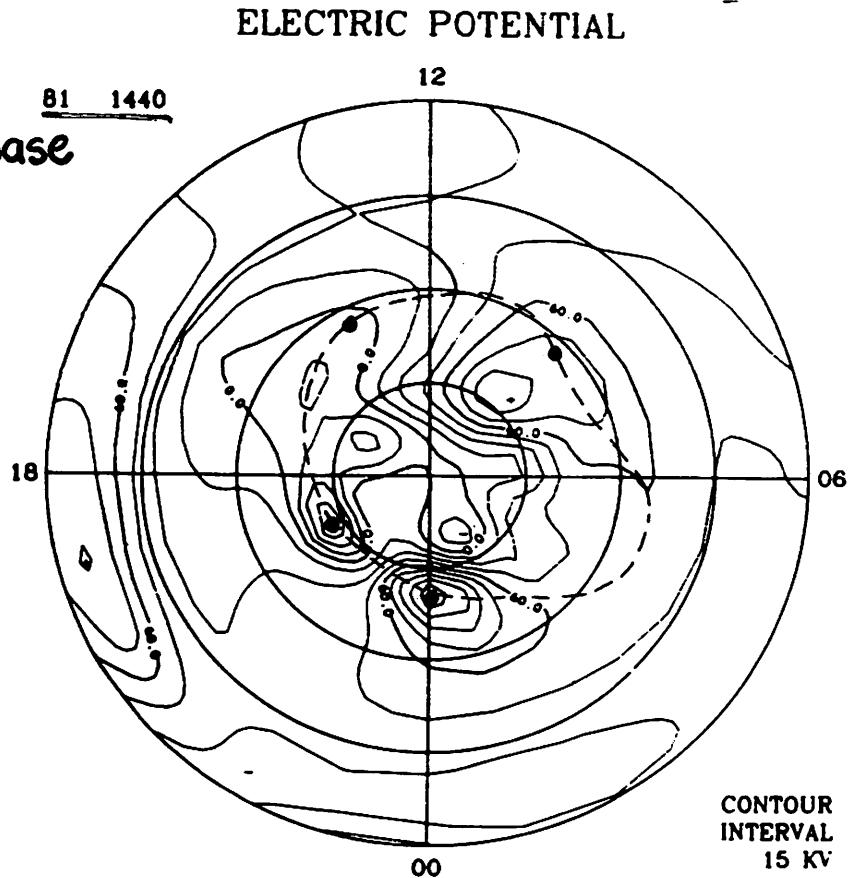
ELECTRIC POTENTIAL

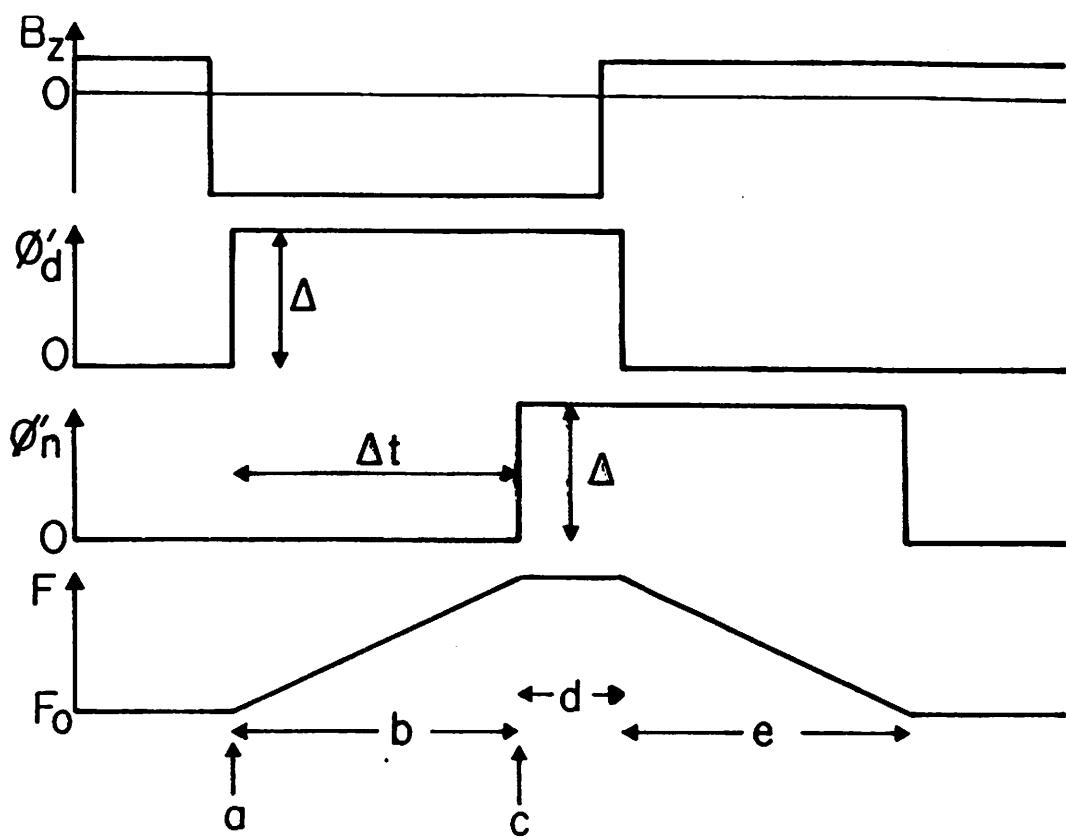
Growth phase



KRM method
CDAW-6
22 March 1979

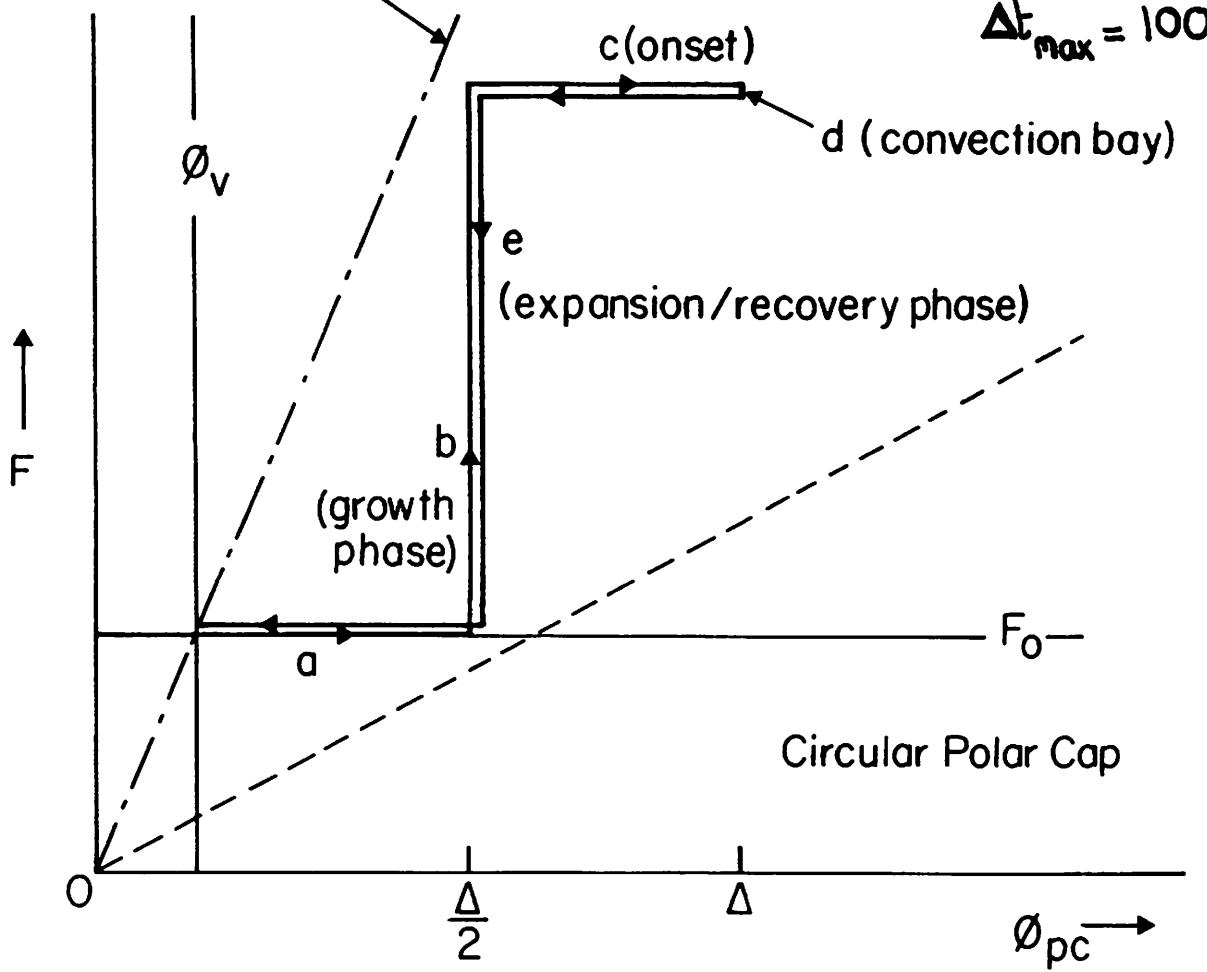
Expansion phase

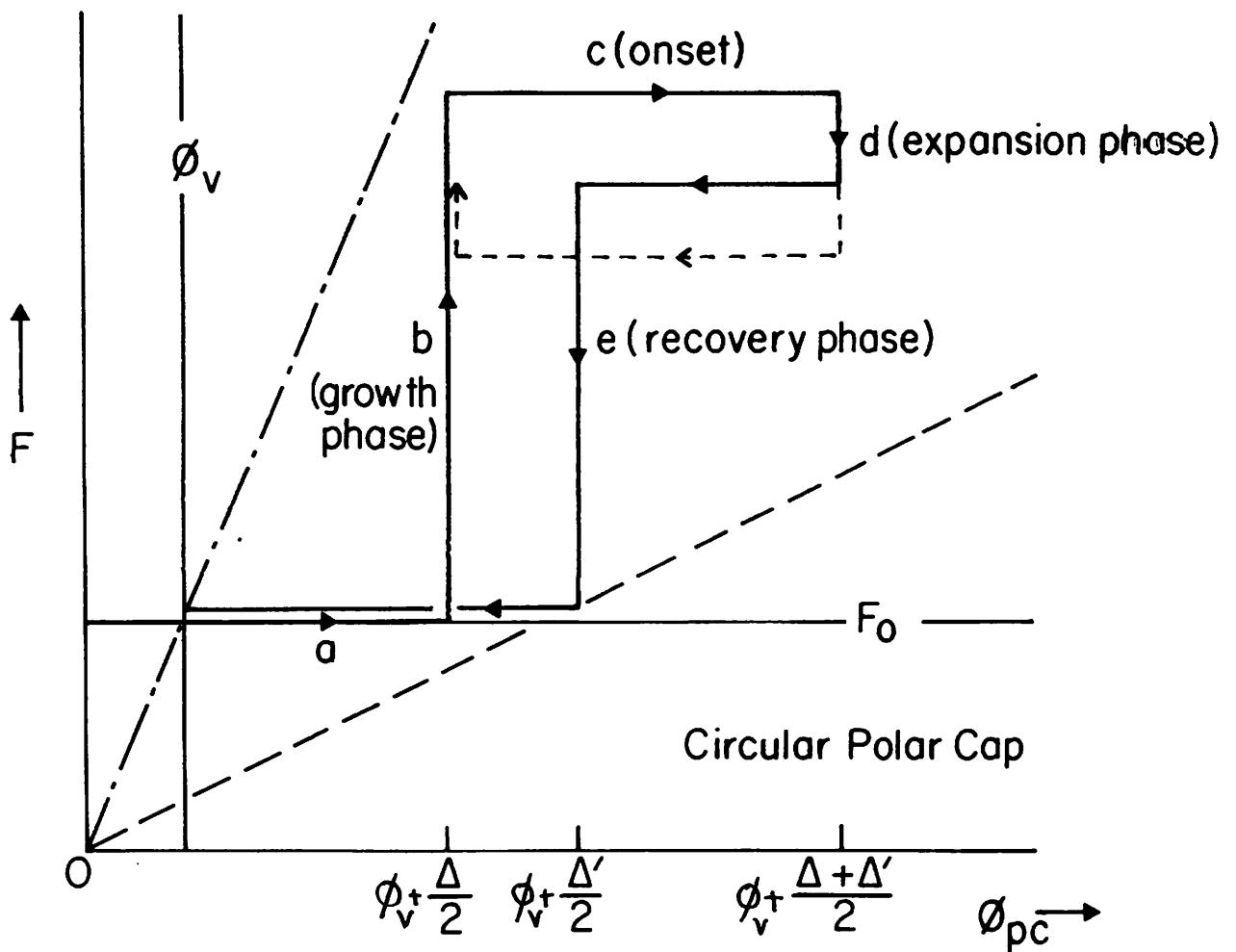
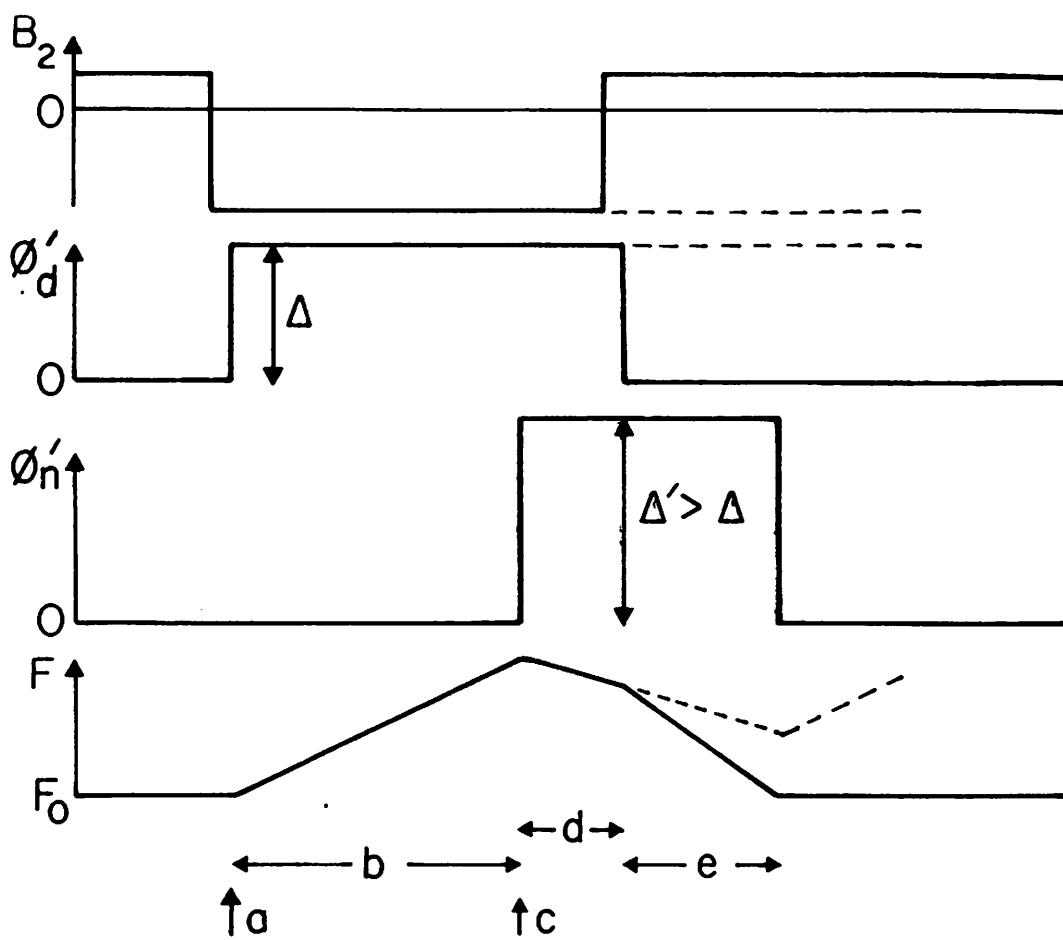




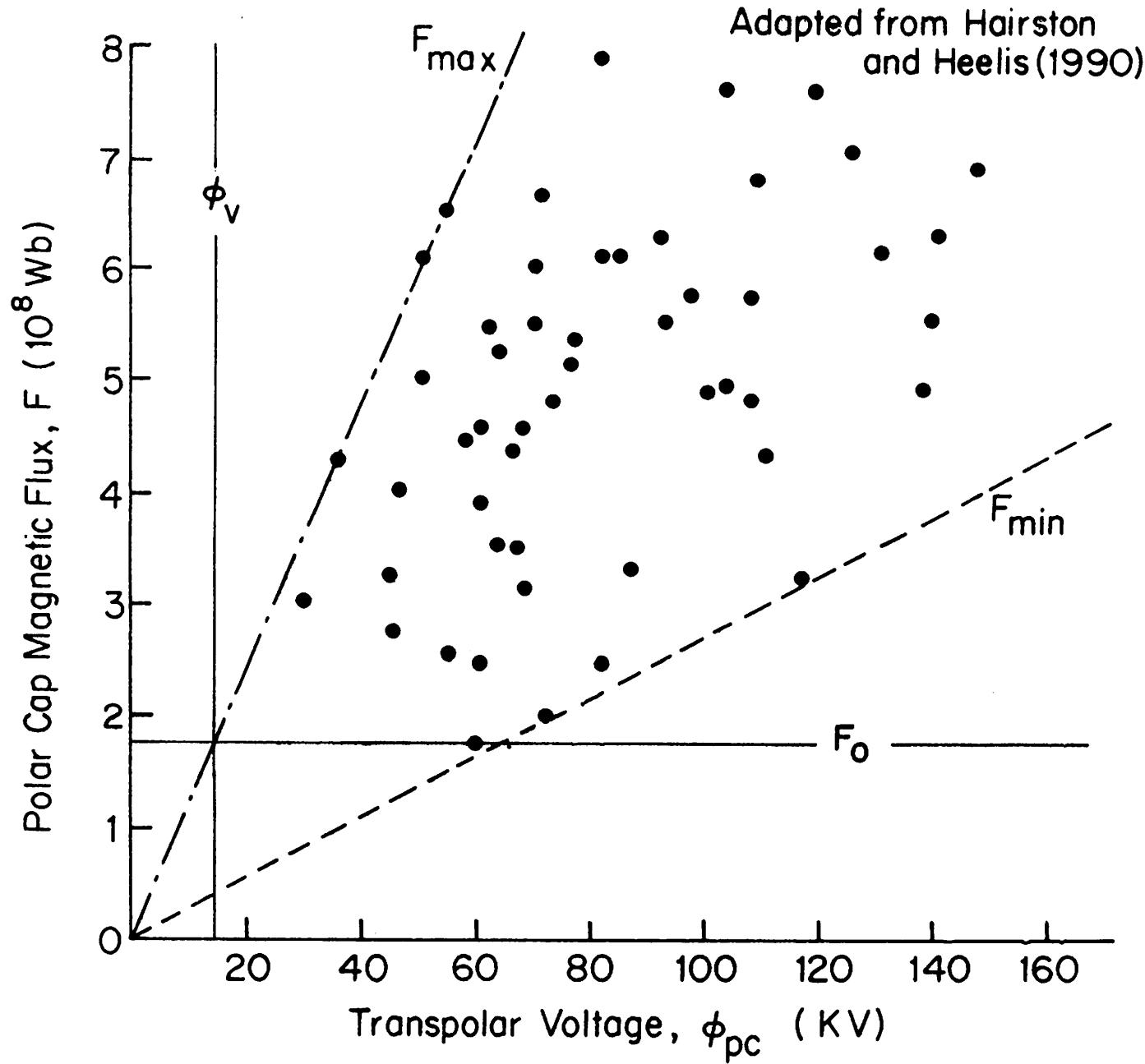
$$F_{\max} = F_0 + 2\Delta t_{\max}(\emptyset_{pc} - \emptyset_v) = F_0 + 1.2 \times 10^4 (\emptyset_{pc} - \emptyset_v)$$

$$\Delta t_{\max} = 100 \text{ min}$$





Lockwood and Cowley (1992)



Summary

- Comparison of ionospheric transpolar voltage with voltage across LLBL shows convection is predominantly driven by reconnection
- Explains B_y and B_z effects on convection
- Consideration of length of magnetotail and variability of IMF B_z shows steady-state is exceptional, in anything other than an average sense
- ALL empirical models of convection and MOST conceptual ones assume steady state
- Electric fields DO NOT map from the magnetopause to the ionosphere in non-steady conditions
- In non-steady cases flow streamlines cross non-reconnecting (but moving) open/closed boundary
- Average convection is the sum of 2 flow patterns driven by dayside and nightside reconnection
- Much viscous-like interaction is in fact residual nightside reconnection
- Transpolar voltage and polar cap size vary in a regular way during the substorm cycle