<u>Horizontal and Vertical Wave Parameters of</u> <u>Thermospheric Gravity Waves, and</u> <u>Relationship to Neutral Winds</u>

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High-resolution, single-beam measurements of gravity waves (GWs) have been made at the Arecibo Observatory (AO) for decades. Features include:

- Vertical wavelength of the spectrum increases exponentially with altitude (Djuth et al, 1997, 2004)
- GWs are omnipresent, although typically appear as discreet wave packets (Djuth et al, 2004; Livneh et al, 2007)

Consequence of single-beam: no horizontal information (including propagation direction). Therefore, GW sources are unknown, although speculations have been made (Djuth et al, 2010).

CEDAR Hine workshop, 1000-1200 Tuesday, June 25, 2013

Given the wave packet nature of the GWs in the thermosphere, a dissipative dispersion relation was derived which allowed for the determination of the parameters of a GW as a function of time explicitly (Vadas and Fritts, 2005). No steady state approximation was made. This dispersion relation was incorporated into a 3D ray trace program (e.g., Vadas, 2007; Vadas and Fritts, 2009).

White noise and deep convective plume spectra of GWs were raytraced into the thermosphere. The results were found to agree well with the AO and MU observations (Vadas, 2007)



Are the GWs in the thermosphere over AO steady state (e.g., full wave model of Waltersheid and Hickey, 2011), or discreet wave packets?

Major difference between theories:

GW wave packet theory: For an individual GW, λ_z increases rapidly in z until the altitude where the momentum flux u_H 'w' is maximum= z_{diss} . Above z_{diss} , λ_z decreases with altitude in a zero-wind environment unless T increases, in which case λ_z increases slowly in z. (e.g., Vadas and Nicolls, JGR, 2012).

Steady state theory: For an individual GW, λ_z increases exponentially above z_{diss} (Waltersheid and Hickey, 2011).

300 250 <u>ل</u> 200 ک 150 100 60 80 100 n 20 λ_{r} (km) 300 250 <u>ل</u> کے 200 150 100 60 80 1001 201 40 Vadas and Nicolls, JGR, 2012 λ_{τ} (km)

140

100

350

5

10

 λ_{-} (km)

15

<u>ل</u> 120

 $mn \gtrsim \lambda_x = 400 \text{ km}, \tau_r = 32 \text{ mn}$

#40

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λ_x=200

24

20

New multi-day dual beam observations are helping to answer these questions

- fixed vertical beam
- off-vertical beam which rotated 90 degrees every few hours



23-25 July, 2009, the AO

TID amplitudes increase rapidly with increasing GW period from τ ~5 to 40 min

For larger periods, the TID amplitude increases slowly



Nicolls et al, JGR, submitted

GW spectral properties change substantially with altitude. Can be explained by GW dissipative theory.



Black: Wind rotates with altitude in mesosphere, with amplitude of 100 m/s Nicolls et al, JGR, submitted



(if possible) their sources

<u>What about individual GWs?</u>

Follow the phase lines up in altitude (Vadas and Nicolls, 2009; Djuth et al, 2010)

Convert electron density perturbations to ion velocity perturbations via a simplified single-ion electron continuity equation with no chemistry and diffusion

$$V_{ion}' \simeq \frac{\omega_r}{\sin I} \frac{\delta N_e}{N_e^0} \left[\frac{-i}{N_e^0} \frac{dN_e^0}{dz} + k_{H||} \cot I + \left(m - i \left\{ \frac{1}{2H} + \frac{1}{g} \frac{dg}{dz} \right\} \right) \right]^{-1}$$

I= local dip angle, D= declination angle of magnetic field, $\delta N_e = TID$ amplitude, $N_e^{0} = background electron density,$ $k_{H}^{-} = horizontal wavenumber, m=vertical wavenumber, <math>\omega_r = wave$ frequency, $\psi = wave propagation direction, k_{H||} = k_{H} \cos(\psi - D),$ $V'_{ion} \alpha g(z) \exp(z/2H)$. Here, we set dg/dz=0 Nicolls et al, JGR, submitted



Therefore, AO thermospheric GWs propagate as wave packets, not steady state wave fields

